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造山带橄榄岩中锆石的成因及其地质意义

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摘要: 造山带橄榄岩记录了板块俯冲、碰撞、折返等复杂过程信息, 可分为壳源和幔源两种类型。造山带橄榄岩(特别是幔源类型)中锆石极为罕见, 锆石内部具有橄榄岩的特征矿物或组合包裹体说明这些锆石可以生长于地幔中。造山带橄榄岩在经历板块汇聚(例如超高压变质等)复杂作用过程中, 经历了不同时期的熔/流体的交代作用, 对橄榄岩的矿物组合和元素组成都能产生重要影响。橄榄岩中锆石作为典型的交代作用产物, 它的形成受控于熔/流体的化学组成、来源属性以及形成物理化学环境等。幔源型造山带橄榄岩中锆石的形成过程可能包括:(1) 锆石结晶能力强, 在地幔环境变化中 Zr 优先与其他地幔硅酸盐矿物中的 Si 结合, 形成锆石;(2) 原始锆石的溶解和含 Zr 矿物相(如石榴石等)的破坏或晶间熔体析出, 在亚固相线条件下形成锆石;(3) 再循环地壳物质来源的熔/流体, 交代地幔楔并结晶形成锆石。因此利用锆石可以揭示特定岩石圈域的演化历史, 有助于深刻理解大陆克拉通及其边缘过程。

关键词: 造山带橄榄岩; 锆石成因; 结晶能力; Zr 元素饱和; 地幔交代作用; 岩石学。

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Genesis and Geological Significance of Zircons in Orogenic Peridotite

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Abstract: Orogenic peridotites record the complex geological processes of continental subduction, collision and exhumation, and they can be divided into two types: crustal origin and mantle origin. Zircons are rarely found in orogenic peridotites. The in-situ zircons and zircon inclusions with mantle-rock mineral assemblages indicate that the zircons can grow within the peridotites. During the complex processes of plate aggregation (such as UHP metamorphism), the orogenic peridotites experienced the melt/fluid interaction in different periods, which have an important impact on the mineral and elemental compositions of peridotites. Zircon is one of the typical metasomatic minerals of orogenic peridotites. The zircon formation is controlled by the composition of the melt/fluid, source properties, and the formation of a physical and chemical environment. The mantle-derived zircons from orogenic peridotites have three origins: (1) Zircon has strong crystallization ability, and Zr prefers to combine with Si from silicate minerals in mantle rocks to form zircons. (2) Metamorphic destruction of Zr-bearing mineral phases and precipitation from intergranular melts generated can nucleate zircons under sub-solidus conditions. (3) The melt/fluid from the recycled crust can metasomatize the mantle wedge and form zircons. Thus, zircons can be used to unravel the history of specific lithospheric domains and thus contribute to our understanding of the evolution of continental cratons and their margins.

Key words: orogenic peridotite; origin of zircon; crystallization capacity; Zr saturation; mantle metasomatism; petrology.

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地质体橄榄岩是造山带的重要组成部分,对其深入研究有利于有效揭示大陆深俯冲、碰撞、折返和块体之间相互作用等复杂过程的丰富信息(Katayama *et al.*, 2003; Zhang *et al.*, 2005; Zheng *et al.*, 2006).锆石分析对追踪地质事件有着重要的地位,通过对其年龄、微量元素和 Hf-O 同位素组成的测定,并结合形态和内部结构等综合分析可以为深入了解壳—幔化学交换作用提供关键记录.研究表明,地幔橄榄岩中的锆石被认为是变质重结晶和熔/流体交代作用的产物(Liati *et al.*, 2004; Song *et al.*, 2005; Zhang *et al.*, 2008; Zheng *et al.*, 2008),反映了地幔经历过由来自软流圈的熔/流体(Griffin *et al.*, 2004; Zheng *et al.*, 2006, 2014)或来自俯冲板块的熔/流体(Katayama *et al.*, 2003; Smith and Griffin, 2005; Liou *et al.*, 2007; Zhao *et al.*, 2008)的交代作用.本文回顾了造山带橄榄岩中锆石的包裹体、微量元素组成、U-Pb 年龄和 Hf 同位素等方面的研究成果,并全面分析了锆石在造山带橄榄岩中出现的可能性,旨在为研究深部俯冲带岩石圈的地球动力学演化提供新的认识.

1 地质体橄榄岩

造山带橄榄岩是指出露于造山带的具有变质结构的地质体(massif)橄榄岩岩块,有别于火山岩中携带的捕虏体(xenolith)橄榄岩.事实上,地质体橄榄岩不仅包括造山带中独立出露的橄榄岩体,也包含蛇绿岩套中的地幔橄榄岩组成部分.在过去的几十年里,人们已经发现了诸多造山带中发育石榴石相橄榄岩,包括闻名中外的阿尔卑斯造山带、大别—苏鲁造山带和俄罗斯的 Sharyzhalgai(Ota *et al.*, 2004)、瑞典的 Scandian 和柴达木北缘(Song *et al.*, 2004, 2005).从图 1 可知,造山带石榴橄榄岩大多出露于亚欧大陆,且多呈层状、透镜状,与榴辉岩、麻粒岩和片麻岩共生.造山带橄榄岩的一个显著特点是它们的年龄多与造山带变质峰期的年龄一致,其中瑞典的 Scandian 和印度尼西亚的 Sulawesi 造山带橄榄岩年龄稍晚于峰期变质年龄(Helmers *et al.*, 1990; Kadarusman, 2000; Brueckner and van Roermond, 2004).起初,Medaris and Carswell(1990)和 Yang *et al.*(1993)认为造山带变质橄榄岩是直接来自上地幔的物质.但是近十几年来,通过越来越多的研究发现,这些造山带橄榄岩多数是地幔楔物质随俯冲板块下插发生高压甚至超高压变质作用的

产物(Brueckner and Medaris, 2000; Spengler *et al.*, 2009).不难发现,对造山带石榴石橄榄岩的研究,有助于深入了解和探讨大陆俯冲的深度、俯冲和折返的动力学机制以及俯冲带与上覆地幔楔之间相互作用等方面的重要科学问题.自 20 世纪 90 年代以来,随着对大陆深俯冲作用研究的不断深入,造山带橄榄岩成为地球科学研究的前沿和热点之一.

1.1 地质体橄榄岩分类

根据宏观、微观的岩石组合、变质变形特征,俯冲—碰撞造山带大致可以分为 A 型—阿尔卑斯型(Ernst and Liou, 1995; Maruyama *et al.*, 1996)和 B 型—环太平洋型(Ernst, 2001; Liou *et al.*, 2004)或者大陆型(continental-type)和大洋型(oceanic-type)(Song *et al.*, 2005).简言之,环太平洋型或大洋型是指洋—洋/洋—陆俯冲,岩石类型主要是大洋岩石圈残留的蛇绿岩带以及岛弧岩浆岩,如环太平洋带.而阿尔卑斯型或者大陆型是指陆—陆俯冲,岩石组合主要为长英质片麻岩、石榴辉石岩、石榴橄榄岩,如大别—苏鲁造山带、特提斯—喜马拉雅造山带、阿尔卑斯造山带.但是,同一个造山带,由于发展历史和阶段不同,也会有两种类型都存在的情况,如柴北缘从大洋俯冲到大陆俯冲的完整记录(Song *et al.*, 2005).

1.1.1 大洋型

大洋型俯冲带中的橄榄岩相对于大陆型具有比较简单的岩石(组合)类型.由于在洋中脊、弧后扩张和俯冲带的蚀变作用下,大部分橄榄岩发生强烈的蛇纹石化,只有少量个体较大的岩块保存着新鲜的地幔橄榄岩.在野外构造上,典型的剖面常可见与上覆的堆晶杂岩和枕状玄武岩以及远洋沉积物(如硅质岩)等紧密共生而构成较为完整的蛇绿岩套(即所谓的“三位一体”),自然界中更多见到的是呈独立的构造岩块分布于造山带中的橄榄岩地质体.典型的蛇绿岩包括雅鲁藏布江缝合带中的部分橄榄岩体(如罗布莎—泽当, Xiong *et al.*, 2016; 徐向珍等, 2018)和北祁连山的玉石沟橄榄岩体(Su *et al.*, 1999; Song and Su, 2010)、西秦岭松树沟橄榄岩体(Cao *et al.*, 2016; Yu *et al.*, 2017)等.大洋型俯冲带岩石圈地幔残片多构成蛇绿混杂岩,而大洋壳的低温高压变质作用多会形成蓝片岩和反映冷俯冲过程的榴辉岩.另外,大洋岩石圈俯冲产物还包括岛弧岩浆岩以及复理石建造等.大洋俯冲地幔橄榄岩主要为大洋或弧后盆地扩张之下软流圈物质经部分熔融形成的方辉橄榄岩和少量二辉橄榄岩.

1.1.2 大陆型

人们曾将广泛出露于大陆造山带

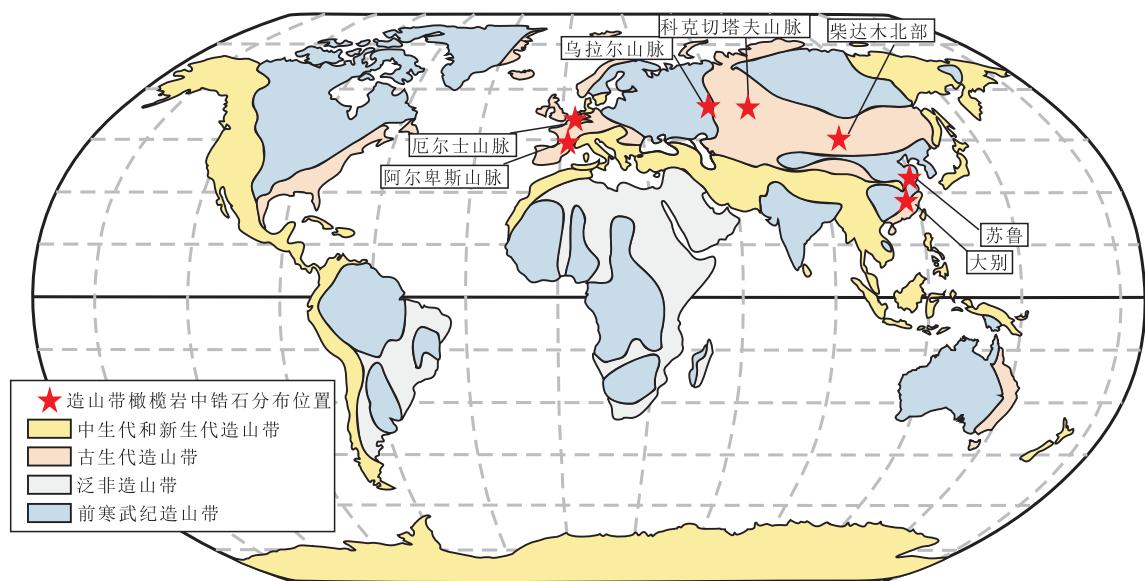


图1 造山带橄榄岩中锆石全球分布

Fig.1 Global distribution map of zircon in orogenic peridotite

据 Liou *et al.*(2009)改

中的橄榄岩称作“阿尔卑斯型”(Alpine-type)橄榄岩(Evans *et al.*, 1977).大陆型造山带橄榄岩可以根据岩石化学成分,进一步划分为 Fe-Ti 型和 Mg-Cr 型,或为 C 型(Crust-type)和 M 型(Mantle-type).

(1) Fe-Ti 型橄榄岩. Fe-Ti 型橄榄岩与 C 型橄榄岩类似,属于先期侵位于大陆深部地壳中的堆晶橄榄岩,在大陆俯冲带中随大陆深俯冲到地幔深度经受超高压变质作用后又折返回地壳的石榴橄榄岩(O’Hara *et al.*, 1971).它们在主量元素特征上与石榴辉石岩、榴辉岩等构成较为完整的超镁铁质—镁铁质岩浆序列.

(2) Mg-Cr 型橄榄岩. Mg-Cr 型橄榄岩与 M 型橄榄岩相对应,属于大陆深俯冲过程中所裹挟的上覆地幔楔底部的橄榄岩,它们与俯冲板片中的长英质和镁铁质岩石共同经历了超高压变质作用,并经历折返而成,属造山带中外来体.在陆—陆碰撞造山带,地幔楔主要来自相对古老的大陆岩石圈地幔;而在弧—陆碰撞造山带,地幔楔则主要是来自于相对年轻的弧下岩石圈地幔.因此,弧—陆碰撞过程也是造山带 M 型橄榄岩常见的构造环境之一.

大陆型俯冲碰撞带的岩石组合与大洋型明显不同,主要体现在:前者具有大陆地壳成分的长英质片麻岩及其呈大小不等的透镜体或构造块体产出的榴辉岩、大理岩、石榴辉石岩和石榴橄榄岩等.其中,作为地幔主体的橄榄岩,包括石榴二辉橄榄岩、石榴方辉橄榄岩、纯橄榄岩和石榴辉石岩等.我国的大别—苏

鲁超高压变质带、挪威的西片麻岩省、哈萨克斯坦的 Kokchetav 超高压地体都属于这种类型的典型地区(Nakajima, 1998; Vrijmoed *et al.*, 2013; Su *et al.*, 2016; Chen *et al.*, 2017a; Li *et al.*, 2018).

本文主要以造山带橄榄岩(大陆型俯冲带橄榄岩)为研究对象,来揭示造山带橄榄岩中的锆石成因和地质意义,从而反映俯冲带中复杂的壳—幔相互作用和变质变形历史.

1.2 造山带橄榄岩形成机制及构造背景

大陆俯冲碰撞造山带形成的 3 个主要动力学过程分别是:前期的洋壳俯冲、陆壳深俯冲和不同速率折返以及造山带抬升.起初,洋壳向陆壳逐渐俯冲,并且带动大陆板块边缘也从地表俯冲至地幔深度(约 100 km 甚至更深).在此过程中被带入俯冲带深部的岩体,包括原先侵位于俯冲板片深部的超镁铁—镁铁质岩石,也包括那些在上覆地幔楔底部的岩石,随深俯冲发生,压力的不断升高,经历一系列超高压变质和改造作用.后期由于大洋岩石圈的重力(多以榴辉岩等密度较大的高压岩石矿物为主)导致其与大陆地壳发生断离,而大陆地壳物质(包含密度较小的蛇纹石化橄榄岩和蛇纹岩)具有明显正浮力,使得俯冲带地幔深部的壳源岩石裹挟不同来源的地幔岩石发生折返过程(郑永飞, 2008).大陆深俯冲的动力学过程信息,可以从大陆边缘造山带中出露的含有柯石英、金刚石等超高压变质矿物的岩石中获取,而后期折返过程可以通过退变质矿物组成

与结构来反映,因此高压—超高压变质带是古板块汇聚边缘、大陆深俯冲碰撞的重要标志(Song *et al.*, 2007)。

石榴橄榄岩是大陆型俯冲碰撞带的标志性岩石类型,不出现在低温的大洋型俯冲带中。因此对其展开系统的形成过程研究,可以很好地还原大陆深俯冲的动力学演变及构造背景。Brueckner and Medaris(2000)汇总世界范围内的石榴子石相橄榄岩,按照化学特征将石榴橄榄岩划分为“壳源”和“幔源”后,再将幔源的橄榄岩形成的构造背景分为 3 类:(1)从不同深度地幔楔进入陆壳的俯冲进变质型橄榄岩,例如 Bohemian 和阿尔卑斯的 Arami 的 UHP 橄榄岩、大别—苏鲁造山带橄榄岩;(2)从超高温尖晶石相地幔冷却后侵位于地壳的超高温型橄榄岩;(3)古老大陆岩石圈地幔中已存在的石榴石相橄榄岩直接被带入并折返回陆壳的残余型橄榄岩。Zhang *et al.*(2000)和 Liou *et al.*(2007, 2009)认为幔源的石榴石相橄榄岩大部分来自俯冲带上盘地幔楔,少量来自俯冲盘的岩石圈地幔和在俯冲之前已经置位于俯冲陆壳内的地幔残片。

造山带石榴橄榄岩在解释俯冲大陆岩石圈地幔、对流地幔、地幔楔和俯冲带的性质和演化过程中起着重要的作用(Brueckner and Medaris, 1998; Medaris, 1999),虽然可以用传统常规的岩石学和地球化学技术来对其进行研究(Zhang *et al.*, 2000; Song *et al.*, 2004; Ye *et al.*, 2009),但这些橄榄岩保存的一些“化石”事件很难用这些手段发现,因此原位微区分析技术的运用就显得特别重要。

2 锆石在造山带橄榄岩中的出现

造山带橄榄岩是超高压碰撞造山带地体中常见组成部分之一,可分为幔源型(M 型)和壳源型(C 型)(Carswell *et al.*, 1983; Brueckner and Medaris, 2000; Zhang *et al.*, 2000)。俯冲大陆释放出的流体会使 M 型造山带橄榄岩更容易经历不同程度的交代作用,实现不同块体间的壳—幔相互作用(Zheng *et al.*, 2016)。因此,大陆俯冲带中的 M 型造山带橄榄岩为俯冲大陆岩石圈地幔的物质传递提供了直接的岩性记录(Zheng, 2012)。

造山带橄榄岩在变质作用过程中,由于熔/流体的加入,使其岩石学和地球化学特征发生明显改变,如其中微量元素丰度和配分型式的明显变化等。它们也普遍存在由于显性交代作用形成的矿物,而隐

性交代只是导致元素或同位素富集(Zheng, 2012)。在造山带橄榄岩中发现较多的含水矿物,为这些岩石受到过地壳来源的熔/流体交代作用提供了矿物学证据。一些典型实例包括:(1)橄榄石包裹利蛇纹石(liz),裂缝部分被叶蛇纹石(atg)充填(图 2a)。由于叶蛇纹石的稳定温压条件高于利蛇纹石,所以这种结构解释为利蛇纹石在后期的退变质过程中,流体沿裂缝渗入与橄榄石反应结晶的叶蛇纹石(Yang and Powell, 2008)。(2)早期橄榄石 Ol₁ 与陆壳深俯冲释放的富硅熔体反应形成 Opx₂,并包裹早期的 Opx₁ 与 Ol₁。推测该 Opx₂ 应继承 Ol₁ 的化学成分特点,而与 Opx₁ 的化学成分存在差异(图 2b, 2c),以及 Ol 中存在 Sp 和 Cpx 共生包裹体的现象(熊发辉等, 2018)。(3)早期橄榄石 Ol₁ 在超高压条件下,会向钛斜硅镁石转变(图 2d):Ol+含 Ti 熔/流体 + H₂O=Chu(樊祺诚等, 1996a, 1996b),说明具有富 Ti 的熔/流体在此期间发生交代作用。(4)在岩相学中可发现金云母(图 2e)、磷灰石等(图 2f),因此推测橄榄岩会受富集 K、P 等元素的熔/流体交代。造山带橄榄岩普遍符合先从上地幔就位于地壳层次深度发生蛇纹石化再经历高压变质作用的规律,记录着岩石圈从裂解到碰撞的演化历史。它们就位于地壳之后也不可避免地会与地壳发生物质交换,从而导致岩石学和地球化学特征发生变化。

传统上,橄榄岩中的锆石多被认为是在进入大陆地壳时由地壳混染所致,这是因为岩石的 Zr 和 Si 的元素含量及活动性较低,因此认为锆石无法从原始橄榄岩中结晶(Hermann *et al.*, 2006; Palme and O'Neill, 2007; Zheng, 2012)。在近几年的研究中,不少学者在全球超高压变质造山带的石榴橄榄岩中均发现了锆石(图 1),包括大别苏鲁造山带(Rumble *et al.*, 2002; Zhang *et al.*, 2005; Zheng *et al.*, 2006, 2008, 2014; Zhang *et al.*, 2011; Li *et al.*, 2016)、柴达木北部(Song *et al.*, 2005; Xiong *et al.*, 2011, 2014; Chen *et al.*, 2017b)、阿尔卑斯山脉(Hermann *et al.*, 2006)和厄尔士山脉(Liati and Gebauer, 2009)等。

在这些造山带橄榄岩中分离出具有残留核部又有新生边部的锆石颗粒,其成因常受质疑;但在造山带橄榄岩的岩石薄片中也发现了原位锆石颗粒的存在(Zhang *et al.*, 2005; Zhang *et al.*, 2011; Xiong *et al.*, 2014; Chen *et al.*, 2017a),这些对该类岩石的成因解释非常重要。如图 3 所示,这些原位观测清楚地表明,锆石是造山带橄榄岩的固有矿物,而不是

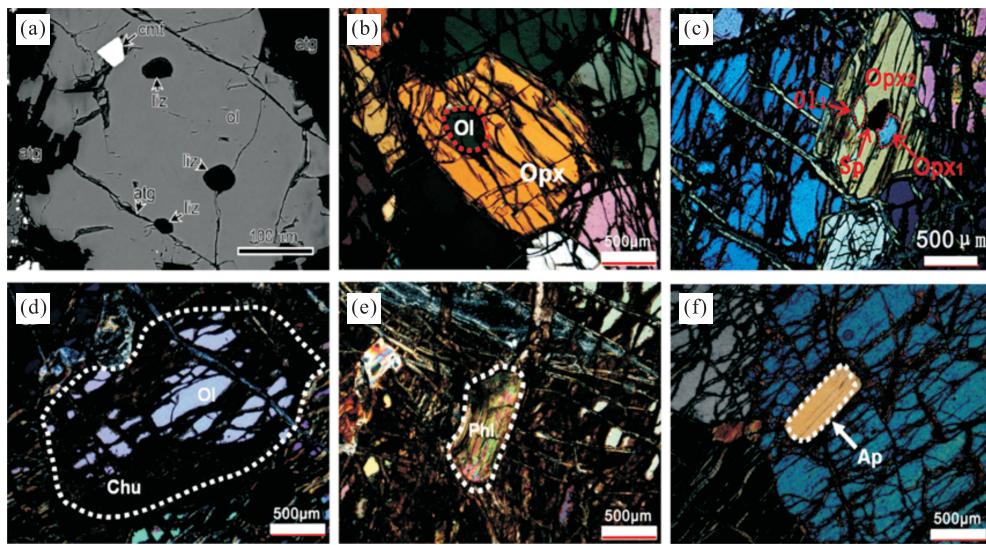


图2 造山带橄榄岩交代证据

Fig.2 Evidence of metasomatism in orogenic peridotite

a.柴北缘胜利口纯橄岩中橄榄石及其包裹体矿物背散射电子图像(杨建军等,2013);b.橄榄石被斜方辉石包裹;c.早期橄榄石(Ol₁)、尖晶石(Sp)和斜方辉石(Opx₁)被晚期斜方辉石(Opx₂)包裹;d~f.橄榄岩中硅镁石、金云母和磷灰石等交代矿物出现(b~f.出自本文采自大别山毛屋地区).liz.利蛇纹石;atg.叶蛇纹石;cmt.铬铁矿;Ol.橄榄石;Opx.斜方辉石;Sp.尖晶石;Chu.硅镁石;Phl.金云母;Ap.磷灰石

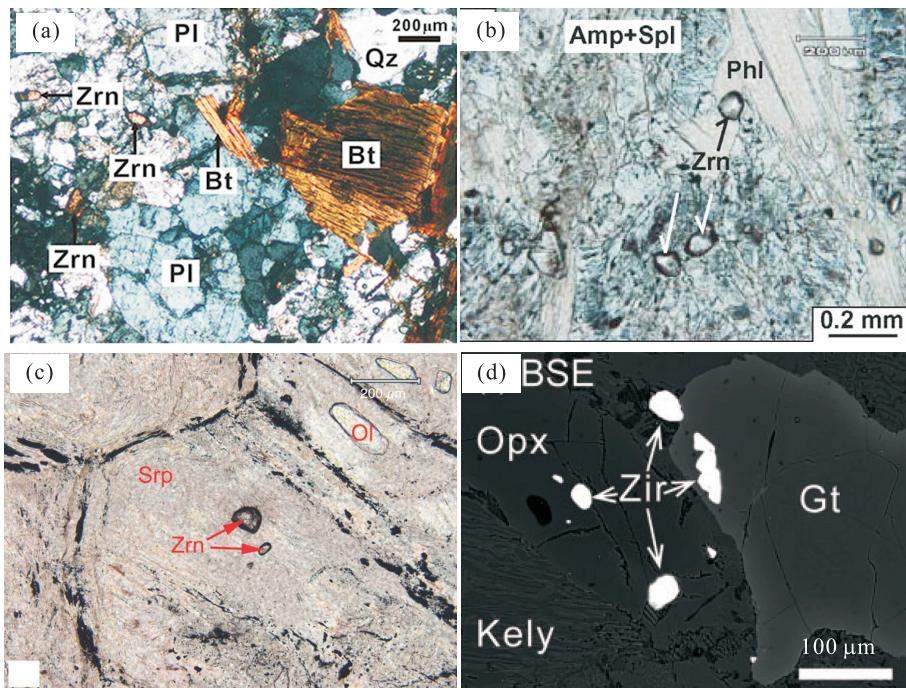


图3 造山带橄榄岩中的锆石在薄片中出现的证据

Fig.3 Evidence for the occurrence of zircons in thin sections of orogenic peridotite

a.苏鲁造山带滕家橄榄岩中锆石产生于石英和斜长石之间(Li et al., 2016);b.苏鲁造山带蒋庄橄榄岩中蛇纹石化橄榄石中包裹原位锆石(Zhang et al., 2005);c.苏鲁造山带芝麻坊异剥橄榄岩中原位锆石的出现(Zhang et al., 2011);d.柴达木北缘胜利口橄榄岩辉石岩脉中锆石包裹于石榴子石和斜方辉石中(Xiong et al., 2014).Zrn.锆石;Pl.斜长石;Qz.石英;Bt.黑云母;Srp.蛇纹石;Ol.橄榄石;Amp.角闪石;Phl.金云母;Opx.斜方辉石;Gt.石榴石;Kely.绿泥镁铝榴石

在挑选矿物过程中的实验混染,在造山带橄榄岩形成中,熔/流体相的存在是必要的,其可将原岩矿物

组合通过脱水作用转化为变质矿物组合.在橄榄岩锆石中熔/流体包裹体的存在表明其是在寄主橄榄

岩与熔/流体交代作用中生长而来的(Zhang *et al.*, 2005; Liati and Gebauer, 2009).因此,尽管造山带橄榄岩中锆石极为罕见,但为交代作用和记录地质事件提供了非常重要的矿物学证据(Gebauer, 1996; Bea *et al.*, 2001; Katayama *et al.*, 2003; Zheng *et al.*, 2008, 2014).

3 造山带橄榄岩中锆石的特征

3.1 锆石微量元素特征分类及意义

锆石的微量元素丰度及配分特征对原岩类型和结晶环境具有一定的指示意义.Belousova *et al.* (2002)利用多元统计的方法所构筑的 CART 树型图提供了一种将母岩类型与锆石微量元素特征联系起来的快速方法,可以在 75% 或更高的置信度范围内识别岩石类型.从超镁铁质岩到镁铁质岩再到花岗岩,锆石微量元素丰度普遍增加.例如,金伯利岩锆石中稀土元素(REE)的平均含量一般小于 50×10^{-6} ,碳酸岩和钾镁煌斑岩为 $600 \times 10^{-6} \sim 700 \times 10^{-6}$,镁铁质岩石中可达 2000×10^{-6} ,在花岗岩和伟晶岩中可以达到百分数量级.金伯利岩和碳酸盐岩锆石具有平坦的 REE 分布,Yb_N/Sm_N 比值在 3~30 之间,在伟晶岩中该比值通常大于 100,Th/U 比值一般在 0.1~1.0 之间.利用 CART 树形图将造山带橄榄岩中锆石的微量元素进行物质来源判别,对理解锆石成因有一定的辅助作用.例如,在柴达木北缘胜利口橄榄岩锆石分析的 14 个点中,有 10 颗锆石的微量元素经过判别发现其与金伯利岩具有相同的来源,其余 4 颗锆石与正长岩相似(Xiong *et al.*, 2011).而苏鲁造山带的许沟、仰口、胡家岭造山带橄榄岩中锆石有 23 颗锆石与玄武岩有相似成因,3 颗锆石与辉绿岩的成因相同(Zheng *et al.*, 2014).这些明显不同于橄榄岩的微量元素特征,可能指示着交代熔体的性质.

3.2 造山带橄榄岩锆石的分类和组成特征

造山带橄榄岩中锆石从年龄上可分为老锆石(主要为残余岩浆锆石和重结晶锆石的核部)和新生锆石.老锆石的形态一般呈自形到半自形状,保留部分锆石晶形,其内部发育有振荡环带或残留锆石核(Liati and Gebauer, 2009; Yang *et al.*, 2009a; Zheng *et al.*, 2014; Li *et al.*, 2016).超高压变质岩中原始锆石的变质重结晶若保留有残余的岩浆锆石核部,Th/U 比值主体分布在 0.1~1.0 之间,在成分组成上几乎继承了原岩锆石的特征,因此常表现出

U-Pb 年龄接近或低于原岩年龄的不一致现象,岩浆成因锆石的稀土元素分布具有典型的陡峭的MREE-HREE 模式,以及高的微量元素含量.

新生的交代成因锆石其形态多呈浑圆状,不具有完好的锆石晶形.其 U-Pb 年龄与超高压变质作用年龄相一致.阴极发光(CL)图像显示,锆石内部结构均一,不具有分带现象.这些交代成因的新生锆石显示出 Th 和 U 含量的变化,Th/U 比值从 <0.01 到 1,化学微量元素的组成各不相同,具有可变的 $\delta^{18}\text{O}$ 值,表明它们是由于地壳岩石进行脱水反应产生的不同成分流体交代作用而形成的.原始地幔矿物经历交代反应后具有交代组分印记(Scambelluri *et al.*, 2006, 2014),交代锆石的微量元素很好地记录了交代介质的性质,变化的微量元素组成主要受变质作用的温压条件以及脱水交代过程中特定矿物的溶解和保存条件所控制.新生长的锆石的氧同位素组成则主要受主岩组成的控制.

结合锆石的 U-Pb 年龄,获取准确 Hf 同位素初始比值,可以很好地示踪物质的来源,了解大陆地壳增长和演化(Amelin, 2000; Bodet and Schärer, 2000; 郑建平, 2005).Zheng *et al.*(2006)通过对苏鲁造山带芝麻坊(CCSD-PP1)橄榄岩中的锆石 Hf 同位素特征分析(ϵ_{Hf} 范围在 -16.3~-13.8 之间),认为该岩体是难熔的太古代地幔的残余,并经历了中元古时期的交代作用.Li *et al.*(2016)对橄榄岩新生锆石进行 Hf-O 同位素分析(ϵ_{Hf} 范围在 -10.7~-5.8 之间),指出苏鲁造山带滕家橄榄岩在扬子板块折返早期经历了来自俯冲大陆板块脱水产生的熔/流体的交代作用.因此,对于锆石 ϵ_{Hf} 为负值有不同的解释模式(Griffin *et al.*, 2000):(1)不同年龄和 Lu/Hf 比值的岩浆储库,经历重熔后产生的岩浆结晶出锆石;(2)源区具有洋岛玄武岩特征的岩浆侵入到岩石圈中经历多阶段演化后结晶形成锆石;(3)以亏损地幔为源区产生的岩浆与岩石圈中非放射产生的 Hf 元素储库反应后,结晶产生的锆石;(4)岩石圈中具有源区 $\epsilon_{\text{Hf}} < 0$ 的特征的岩浆结晶产生的锆石(Nowell and Pearson, 1998; Nowell *et al.*, 1998a, 1998b).

锆石同位素特征可以提供交代流体组成和来源信息.大陆岩石圈地幔的 $\delta^{18}\text{O}$ 值范围:橄榄石在 $4.8\text{\textperthousand} \sim 5.5\text{\textperthousand}$,石榴石在 $5.4\text{\textperthousand} \sim 5.8\text{\textperthousand}$,斜辉石在 $5.2\text{\textperthousand} \sim 5.9\text{\textperthousand}$ 和单斜辉石在 $5.6\text{\textperthousand} \sim 6.0\text{\textperthousand}$ (Mattey *et al.*, 1994; Chazot *et al.*, 1997).在苏鲁造山带滕家纯橄榄岩的锆石具有负的 $\delta^{18}\text{O}$ 值,是由于超高压

火成岩原岩结晶于低 $\delta^{18}\text{O}$ 岩浆并受到了高温热液蚀变释放出负 $\delta^{18}\text{O}$ 流体交代橄榄岩而产生(Li et al., 2016)。大别造山带毛屋石榴辉石岩通过 Mg-O 同位素对碳酸盐交代源区给出制约(Shen et al., 2018)。毛屋石榴辉石岩 Mg 同位素组成($\delta^{26}\text{Mg} = -0.99\% \sim -0.65\%$)显著低于地幔 Mg 同位素含量($\delta^{26}\text{Mg} = -0.25\% \pm 0.07\%$)，而毛屋辉石岩氧同位素远高于地幔碳酸盐氧同位素，说明交代碳酸盐来源为循环的地表沉积碳酸盐。在柴达木北缘绿梁山造山带橄榄岩中锆石的 Hf-O 同位素组成记录了两种源区的交代作用，高 $\delta^{18}\text{O}$ 流体起源于深俯冲大陆地壳，低 $\delta^{18}\text{O}$ 流体来自大陆俯冲之前的俯冲洋壳脱水(Chen et al., 2017b)。

3.3 造山带橄榄岩锆石中的包裹体

在不同的环境下生长的锆石可以俘获矿物、液体和熔体作为包裹体，为其形成条件和机制提供重要的记录(Hermann et al., 2001; Liu and Liou, 2011; Li et al., 2013)。它们一旦被俘获，由于锆石的稳定性，其组成几乎不会发生变化。Liu and Liou (2011)对大别—苏鲁造山带各变质岩中的锆石进行了详细的矿物包裹体研究发现：在俯冲过程中，变质锆石中的包裹体以石英、石榴石、绿泥石、金红石、钾长石、白云岩和磷灰石为主，在峰期榴辉岩相变质阶

段，包裹体以柯石英、石榴石、绿泥石、白云石、金红石、翡翠、蓝晶石、钛石、钾长石、文石、菱镁矿和磷灰石为主。在角闪岩相退变质阶段，包裹体仅由石英、斜长石、钠长石、角闪石、方解石和磷灰石等低压矿物组成。包裹体的变质组合不仅与变质条件有关，而且还受宿主岩石成分的控制。对它们的分析，结合锆石 U-Pb 年龄，可以为锆石寄主岩的 P-T-t 轨迹进一步提供重要的约束(Hermann et al., 2001; Liu and Liou, 2011)。

而橄榄岩中锆石是指示交代流体来源的重要证据。造山带橄榄岩中的锆石，它们经历了不同程度的交代作用，因而常有含水矿物，如绿云母、角闪石、钛斜硅镁石等(Katayama et al., 2003; Zhang et al., 2005, 2011; Hermann et al., 2006)。通常情况下，造山带橄榄岩锆石的包裹体组成与橄榄岩组成矿物一致，说明锆石产生于橄榄岩之中，是在地幔橄榄岩内经历了流体交代或变质结晶而来(图 4a~4c)。如在柴达木北缘胜利口橄榄岩和辉石岩的锆石中，发现了与地幔矿物一致的石榴石、斜方辉石和橄榄石等矿物包裹体组合，说明锆石是原始橄榄岩形成阶段结晶的产物(Xiong et al., 2011)。但也有不一样的情况，如在大别—苏鲁、阿尔卑斯和波希米亚等地造山带橄榄岩中新生的锆石中发现了磷灰石、辉绿石、角

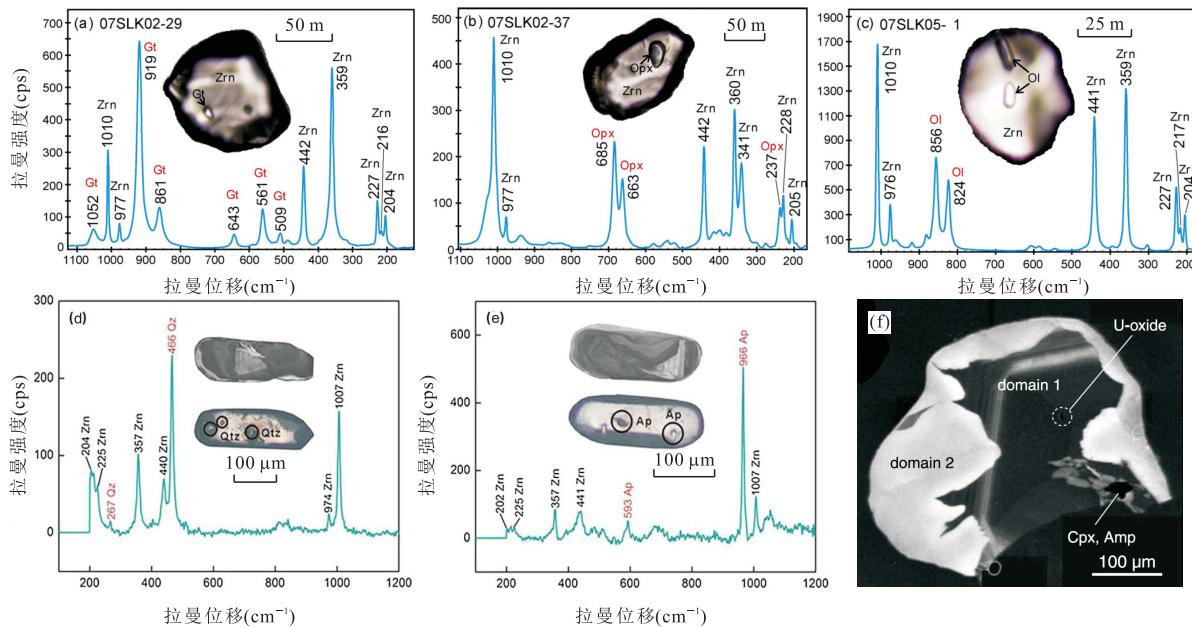


图 4 造山带橄榄岩锆石中的包裹体

Fig.4 Inclusions in zircons from orogenic peridotite

a~c.柴达木北缘胜利口橄榄岩锆石中的石榴石、斜方辉石和橄榄石包裹体(Xiong et al., 2011);d~e.苏鲁造山带滕家橄榄岩锆石中石英和磷灰石包裹体(Li et al., 2016);f.阿尔卑斯 Duria 橄榄岩锆石中单斜辉石、角闪石和铀氧化物包裹体(Hermann et al., 2006)。Zrn.锆石, Gt.石榴石, Opx.斜方辉石, Ol.橄榄石, Qtz.石英, Ap.磷灰石, Cpx.单斜辉石, Amp.角闪石, U-oxide.铀氧化物

闪石和氧化铀等矿物包裹体以及长石熔体包裹体(图 4d~4f)(Zhang *et al.*, 2005, 2011; Hermann *et al.*, 2006; Liati and Gebauer, 2009),这些非地幔矿物包裹体的存在表明它们的形成可能与地壳物质参与有关(Hermann *et al.*, 2006).总之,结合造山带橄榄岩中新生锆石形态和内部结构特征,以及微量元素变化和年龄特点,可以更好地约束交代流体的性质、成分以及交代作用发生的时间.

4 造山带橄榄岩中锆石的成因意义

对于造山带橄榄岩锆石的成因一直存在争议,对其成因机制及交代流体来源还没有定论.Zheng *et al.*(2014)认为造山带橄榄岩中的锆石与深俯冲再循环地壳物质交代作用有关,这些地壳交代物质是在俯冲通道的壳幔界面中发生变质脱水和部分熔融产生的.虽然在地幔储库中,Zr 元素含量微乎其微,活动性弱,理论上不具备形成锆石的条件,但是在上述诸多证据中,是可以证实锆石在造山带橄榄岩中是真实存在的.因此,如何打破前人的认识,并科学地解释锆石成因,对造山带橄榄岩的研究有着进一步推进作用.造山带橄榄岩锆石形成可能有 3 方面原因.

(1) 锆石有非常强的结晶能力.Zr 作为不相容元素,在地幔环境中锆石比其他地幔矿物(如石榴石等)具有更强的与 Zr 元素结合并结晶的能力.因此,在变质反应下 Zr 元素会优先与地幔中的 Si、Hf、Th、U 等元素结合形成锆石.当地幔岩石中全部的 Zr 元素吸纳 Si 形成锆石之后,剩余的其他元素才依次组合形成其他矿物.

(2) 原始锆石的溶解和含 Zr 相矿物的破坏或晶间熔体析出.一般来说,锆石的形成与生长需要它们的宿主岩石供应 Zr.这种元素可以从石榴石、黑云母、角闪石和钛铁矿等含 Zr 矿物的分解中得到(Zheng, 2012),也可以来自原始锆石的溶解(Vavra *et al.*, 1999; Ayers *et al.*, 2003).由于 Zr 是一种不相容元素,在大多数地质环境中通常是不稳定的,因此通常认为锆石会容易产生在岩石基质中有 Zr 元素释放的反应附近.而在地幔环境中,石榴石在超高压变质过程中的解体与熔融为橄榄岩中锆石的形成提供了条件.Degeling *et al.*(2001)曾在来自挪威西南部含石榴石混合岩中发现,许多小颗粒锆石生长在由于石榴石分解而形成的堇青石中.另外,Fraser *et al.*(1997)和 Bingen *et al.*(2001)分别报道了由于

石榴石和钛铁矿的分解而生长出了原位锆石.由于 Zr 在水溶液中的不易迁移性,通常不能通过流体有效转移,但是 Zr 作为不相容元素(Harrison and Watson, 1983; Watson, 1996; Harrison *et al.*, 2007),在熔体的变质反应中具有将反应物转化为具有 Zr 局部饱和的能力,促进锆石的生长.因此在贫锆岩石中,也可从熔体中生长出锆石,它们的 U-Pb 年龄可以给出熔融结晶的时间和对应处相应的地质事件.变质反应中石榴石的赋存情况也会对锆石生长过程的性质产生影响(Rubatto, 2002; Whitehouse and Platt, 2003).如果石榴石和锆石都作为变质反应的产物出现,由于重稀土元素(HREE)被优先划分为石榴石,锆石则反映出平滑的 HREE 特征.同样地,如果石榴石是反应的残余相,由于其具有高的 HREE 含量,产物锆石也会在 HREE 中表现出耗尽.相反,如果石榴石作为反应物参与反应,则产物锆石将富集 HREE.因此,在超高压变质反应过程中,含 Zr 相矿物的解体或晶间熔体析出,均为锆石的形成提供 Zr 元素,锆石在亚固相线条件下成核,形成锆石.另外,在晚期折返退变质作用和熔体冷却过程中,也可有锆石的生长.变质成因锆石伴随有脱水反应,变质反应有能力将反应物转化为固相产物使局部达到 Zr 饱和,促进变质锆石和石榴子石的生长并释放熔/流体.因此,锆石在变质之前的原岩性质和含 Zr 矿物在俯冲带变质期间的行为决定了造山带橄榄岩中锆石的生长和重结晶.

(3) 携带 Zr 和其他锆石元素的循环变质熔/流体交代.熔/流体相在俯冲带变质过程中对锆石生长和重结晶有重要的作用(Rubatto *et al.*, 1999; Rubatto and Hermann, 2003; Wu *et al.*, 2006; Zheng, 2009).交代成因的新生锆石通常归因于以下两种机制之一:① 地幔起源的交代熔/流体结晶而成(Grieco *et al.*, 2001; Zheng *et al.*, 2006).如苏鲁造山带的芝麻坊橄榄岩经历了早中生代(223.5 ± 7.5 Ma)的超高压变质作用,其中锆石记录的交代介质具有与金伯利岩和碳酸岩的微量元素相似的特征,表明橄榄岩锆石记录下了来自软流圈地幔的熔/流体在中元古代对橄榄岩进行了交代改造,并不是来自俯冲大陆组分(Zheng *et al.*, 2006).② 从深俯冲地壳板块脱水而来的交代流体结晶而成(Hermann *et al.*, 2006; Liati and Gebauer, 2009; Li *et al.*, 2016; Chen *et al.*, 2017a).地壳岩石中具有很高的 Zr 含量,因此,在大陆俯冲隧道的壳幔界面,地壳脱水产生的熔/流体发生交代作用,非常容易形成壳源

的变质锆石(Zheng, 2012).在这两种情况下,都强调交代熔/流体来源决定了造山带橄榄岩中锆石的意义.大量研究表明变质锆石一般具有相似的微量元素特征(Chen et al., 2010; Liu and Liou, 2011),这些通过变质反应产生的锆石,在流体作用下,微量元素组成可能均一化.此外,流体的存在也将大大促进原生锆石的重结晶作用.熔/流体交代程度不仅决定了新生锆石的生长,还决定了俯冲带变质作用中变质重结晶的程度.因此,碰撞造山带中橄榄岩的锆石研究可以为俯冲带的熔/流体作用提供重要的制约.

5 造山带橄榄岩中锆石的地质年代意义

造山带橄榄岩锆石的年龄分布范围广,记录了复杂的壳幔相互作用和地质演化历史.在超高压变质作用的不同阶段中(Hermann et al., 2006; Liati and Gebauer, 2009; Li et al., 2016; Chen et al., 2017a),锆石U-Pb年龄、Lu-Hf同位素以及氧同位素分析,是综合反演地质演化过程的重要手段.

我国北部柴北缘胜利口(或称为绿梁山)造山带石榴橄榄岩记录着早古生代(~ 440 — 420 Ma)大陆深俯冲过程(深度超过200 km),揭示了俯冲隧道中复杂的壳—幔相互作用(Song et al., 2004).最新的研究成果显示胜利口造山带橄榄岩,可能是具有太古宙起源的大陆岩石圈地幔楔(Xiong et al., 2011, 2014, 2016; Chen et al., 2017b),不同于Song et al.(2007)认为是阿拉斯加型超镁铁质堆晶杂岩体和Yang and Powell(2008)认为的早先置位于俯冲陆壳的大洋岩石圈残片等.通过Re-Os、Lu-Hf、Sm-Nd

和U-Pb同位素等分析,认为胜利口橄榄岩杂岩体可能来自中太古代克拉通型岩石圈地幔,并伴随Columbia和Rodinia超大陆裂解时期的硅酸盐熔体交代.在早古生代时期记录了一套完整的洋—陆俯冲到陆—陆俯冲的造山过程(Xiong et al., 2011, 2014, 2016).

大别—苏鲁造山带M型橄榄岩 $Mg^{\#}$ 值在92~90间,以过渡类型为主,说明华北克拉通在演化过程中其东南缘受到俯冲岩石圈板块所释放流体组分的交代改造和破坏作用(Zhang et al., 2002).大别—苏鲁造山带橄榄岩的锆石具有 ~ 3.1 Ga古老年龄,Hf同位素也具有一个 ≥ 3.2 Ga的组分,在华北克拉通南缘信阳火山岩中的橄榄岩包体也发现了相似的太古代年龄(Zheng et al., 2006).苏鲁造山带的许沟、仰口、胡家岭、赣榆,以及大别造山带的饶巴寨等地的M型橄榄岩经历了元古代交代改造,随后记录了一个早古生代强烈的构造热事件,U-Pb年龄 ~ 470 Ma,与华北克拉通内部早古生代蒙阴和复县金伯利岩喷发事件相一致(Lu et al., 1998; Yang et al., 2009a, 2009b; Li et al., 2011), ϵ_{Hf} 值在11.2~20.1之间.这一事件被解释为来自包括软流圈、岩石圈地幔和俯冲大陆板在内的多种来源的熔/流体交代作用的结果.大别—苏鲁造山带变质锆石的U-Pb年龄主要分布在三叠纪超高压变质时期(图5a),这与早中生代(~ 230 Ma)华北与扬子板块的碰撞事件相关.大别—苏鲁造山带橄榄岩的锆石记录了早中生代时期扬子板块向北俯冲/碰撞华北板块的过程(Zheng et al., 2005; Wu et al., 2006; Zheng, 2009; Xia et al., 2013).M型橄榄岩

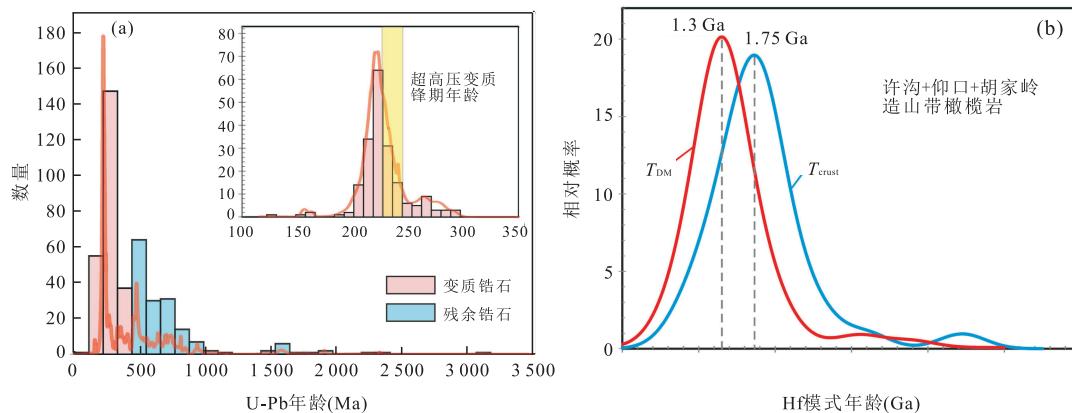


图5 大别—苏鲁造山带橄榄岩锆石U-Pb年龄和Hf模式年龄分布

Fig.5 U-Pb age and Hf model age distribution of peridotite in Dabie-Sulu orogenic belt

图a据Yang et al.(2009a)、Zhang et al.(2005)、Zhao et al.(2006)、Zheng et al.(2006, 2008, 2014)、Rumble et al.(2002)和Li et al.(2016);图b据Zheng et al.(2014)

交代成因的新生锆石 U-Pb 年龄范围在 205~244 Ma, 记录了扬子板块俯冲/碰撞到华北板块之下过程。这些 U-Pb 年龄大部分聚集在 212~227 Ma, 晚于 225~240 Ma 的超高压变质年龄, 表明它们是在早期深俯冲陆壳的交代作用下生长的。这些交代锆石的 Th、U 含量和 Th/U 比值在 <0.01~1 之间变化很大, 表明它们在地壳岩石下通过脱水反应发生交代作用而生长。矿物氧同位素也对交代过程中物质加入的来源起到识别作用 (Zheng *et al.*, 2003)。苏鲁造山带滕家纯橄榄岩中新生锆石具有负的 $\delta^{18}\text{O}$ 值, 为 $-11.3\text{\textperthousand} \sim 0.9\text{\textperthousand}$ (Li *et al.*, 2016), 与该区超高压片麻岩的氧同位素组成一致, 显示造山带橄榄岩中新生锆石受到壳源流体交代作用的可能性极大。苏鲁一大别造山带的 M 型橄榄岩中的锆石具有一致的 Hf 亏损地幔模式年龄 (T_{DM} 峰值为 1.30 Ga) 和 Hf 地壳模式年龄 (T_{crust} 峰值为 1.75 Ga)(图 5b), 表明华北克拉通东南缘下的大陆岩石圈地幔经历了强烈的中元古代软流圈交代作用, 类似于华北克拉通东部大陆岩石圈地幔性质 (Zheng *et al.*, 2014)。因此, 大别—苏鲁超高压变质造山带中橄榄岩锆石的 U-Pb 年龄、Hf 同位素系统和微量元素组成主要记录了华北克拉通南缘深部岩石圈 (≥ 3.2 Ga) 经历了元古代交代作用, 随后发生了强烈的早古生代 (~ 470 Ma) 构造热事件和早中生代 (~ 230 Ma) 的扬子板块向北俯冲、碰撞超高压变质 (Zheng *et al.*, 2014), 使得俯冲大陆岩石圈上覆的华北克拉通东南缘岩石圈地幔经历了多期次的交代与改造。

6 结语

造山带橄榄岩中的锆石是变质重结晶和熔/流体交代的产物。在大陆俯冲碰撞过程中, 原始锆石在高压—超高压条件下发生不同程度的改造, 这主要取决于其理化性质(锆石结晶能力, 锆石及其他富 Zr 矿物晶体中裂隙的存在与否)以及变质熔/流体的交代能力。造山带橄榄岩的锆石是记录地壳与地幔之间物质和能量转换的重要媒介, 它进一步制约了大陆俯冲带地幔楔组成演化、壳幔相互作用、流体—岩石相互作用及俯冲带地球动力学等重要信息, 是值得重视的研究对象, 特别是目前造山带橄榄岩中锆石 O 和 Li 同位素的分析结果很少, 应该也是未来值得重视的方向。

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