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吉林临江八道沟二长花岗岩地球化学、 Hf 同位素特征及其成因

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摘要: 八道沟具有复杂的构造环境, 且侏罗纪花岗岩在该地区分布较为有限, 因此十分缺乏对区域构造演化环境判断的证据. 对八道沟二长花岗岩进行了地球化学、锆石 U-Pb 年代学及 Hf 同位素研究, 结果显示: 岩石样品中的锆石主要为长柱状, 自形程度较好, 发育较清晰的韵律生长环带结构, 具岩浆锆石特征, 该岩石形成于早侏罗世 (176 ± 1 Ma); 岩石 SiO_2 含量为 68.93%~70.02%, 贫 MgO (1.27%~1.45%), CaO (1.80%~2.78%), A/CNK 值为 0.92~1.02, 轻稀土元素明显富集, 铕异常不明显 (0.96~1.03), 属于准铝质高钾钙碱性系列, 显示为 I 型花岗岩特征; 岩石 $\epsilon_{\text{Hf}}(t)$ 值为 $-20.21 \sim -24.08$, 二阶段模式年龄 $t_{\text{DM2}}(\text{Hf})$ 为 2 498~2 740 Ma, 且富集大离子亲石元素, 亏损高场强元素 Nb、Ta、P、Ti, $\text{Mg}^\#$ 值介于 45~47, 由此推断岩石源于新太古代下地壳熔融. 结合区域同时代岩浆岩展布, 认为八道沟地区二长花岗岩的形成与古太平洋板块向欧亚大陆俯冲有关.

关键词: 二长花岗岩; 锆石 U-Pb 定年; Hf 同位素; 早侏罗世; 八道沟; 地球化学; 地质年代学.

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Geochemistry, Hf Isotopes and Petrogenesis of Badaogou Monzonitic Granites from Linjiang, Jilin Province

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Abstract: The tectonic setting is complex while Jurassic granite is relatively limited in the Badaogou area, which results in a lack of theoretical basis for its regional tectonic evolution studies. This paper presents a synthesis of geochemical and zircon U-Pb chronological and Hf isotopic study on the Badaogou monzonitic granites. The zircons of rock samples which have distinct rhythmic growth zoning are long cylindrical and euhedral well-formed, implying the features of magmatic origin. Dating results indicate that the rocks were formed in the Early Jurassic (176 ± 1 Ma). The rocks which obviously are rich in LREE and have un conspicuous europium anomalies (0.96–1.03) with regard to high SiO_2 (68.93%–70.02%), low MgO (1.27%–1.45%), CaO (1.80%–2.78%) as well as A/CNK (0.92–1.02) belong to sub-aluminous and high-K calc-alkaline granites, being of the characteristics of I-type granites. The $\epsilon_{\text{Hf}}(t)$ values vary from -20.21 to -24.08 , and the two-stage model age (t_{DM2}) ranges from 2 498 to 2 740 Ma. These samples are enriched in LILE and depleted in HFSE (such as Nb, Ta, P, Ti), with $\text{Mg}^\#$ values of 45–47. Analyses show that the rocks came from the melting of neo-Archean lower crust. Combining with the background of

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regional contemporary magmatic distribution, it is concluded that the formation of monzonitic granites from Badaogou could be related to the subduction of the Paleo-Pacific plate under the Eurasian continent.

Key words: monzonitic granites; zircon U-Pb dating; Hf isotope; Early Jurassic; Badaogou; geochemistry; geochronology.

八道沟位于吉林省临江市东南部,大地构造上处于华北板块东北缘,隶属于辽吉古元古代造山带,其南部为狼林地块(李雪梅,2009)(图 1).研究区所在的中朝板块自太古宙起经历了一系列大规模的构造运动,区内构造环境复杂,一直是备受争议的热点地区.自中生代开始,该地区主要受北部的古亚洲洋构造域及东部的环太平洋构造域影响.前人普遍认为在早一中三叠世,长春-延吉缝合线最终闭合,古亚洲洋彻底消失(Wu *et al.*, 2011; 彭玉鲸等, 2012),随后进入环太平洋构造体系.作为俯冲作用产物之一的花岗质岩浆,在八道沟所在的吉南地区分布广泛,但现有研究结果显示侏罗纪花岗岩分布较为有限,目前仅见于荒沟山地区的草山、梨树沟、老秃顶子及遥林岩体(孙德有等, 2005; 秦亚等,

2013),明显缺乏对区域构造演化环境判断的充分证据.此外,吉林省地质调查研究院在研究区北部发现了较具规模的八道沟金矿,孙永杰等(2007)研究认为区内花岗岩体的构造蚀变带是找矿重要标志之一.因此,本文对八道沟地区出露的花岗岩进行了详细的年代学、岩石地球化学及 Hf 同位素研究,确定了其形成时代,从而揭示该区侏罗纪构造背景,并为八道沟金矿提供一些研究资料.

1 地质背景

研究区内出露的地层包括下元古界老岭群花山组、新元古界青白口系钓鱼台组、寒武系及奥陶系,研究区北部出露有侏罗系且零星分布新近系船底山

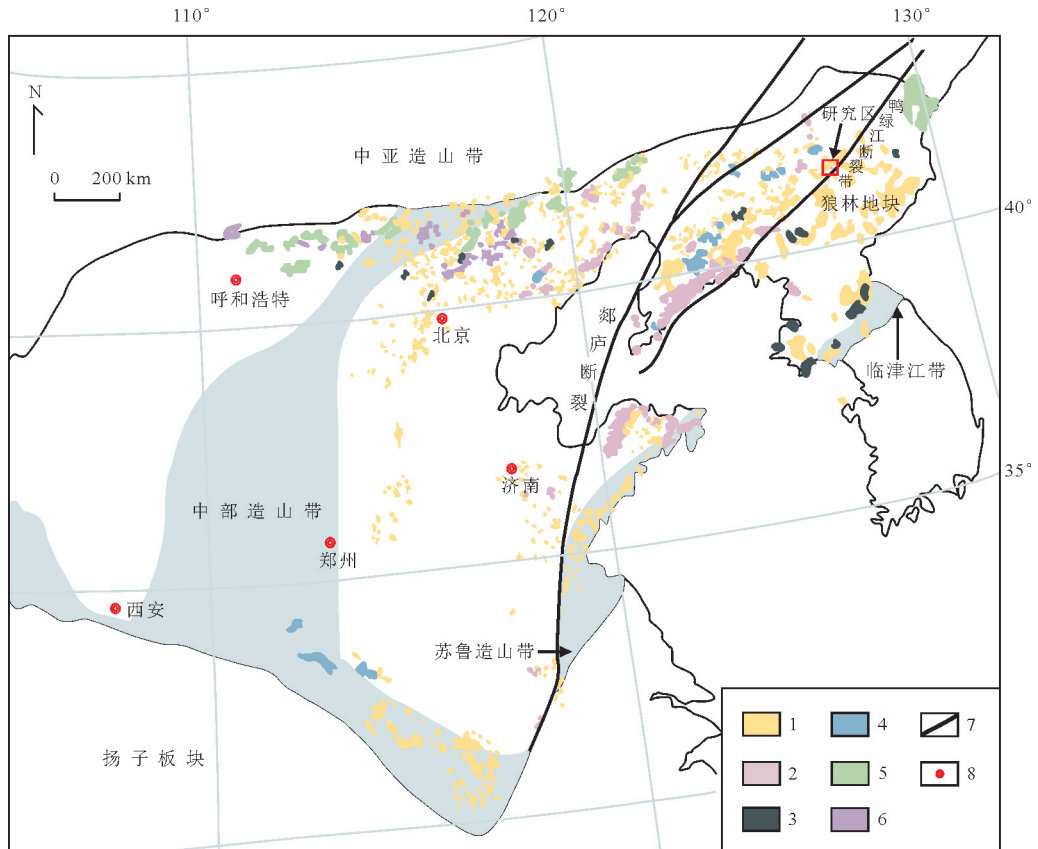


图 1 研究区大地构造位置及岩浆岩分布

Fig.1 Geotectonic location and magmatite distribution of the study area

1.早白垩世侵入岩, 2.侏罗纪花岗岩, 3.晚三叠世碱性岩, 4.晚三叠世闪长岩-花岗岩, 5.晚二叠-中三叠世侵入岩, 6.石炭纪-早二叠世侵入岩, 7.断裂带, 8.地名; 据 Xu(2001), Yang *et al.*(2003,2005), Zhao *et al.*(2006), Li *et al.*(2012), Zhang *et al.*(2014), Deng and Wang(2016), 刘金龙等(2016)

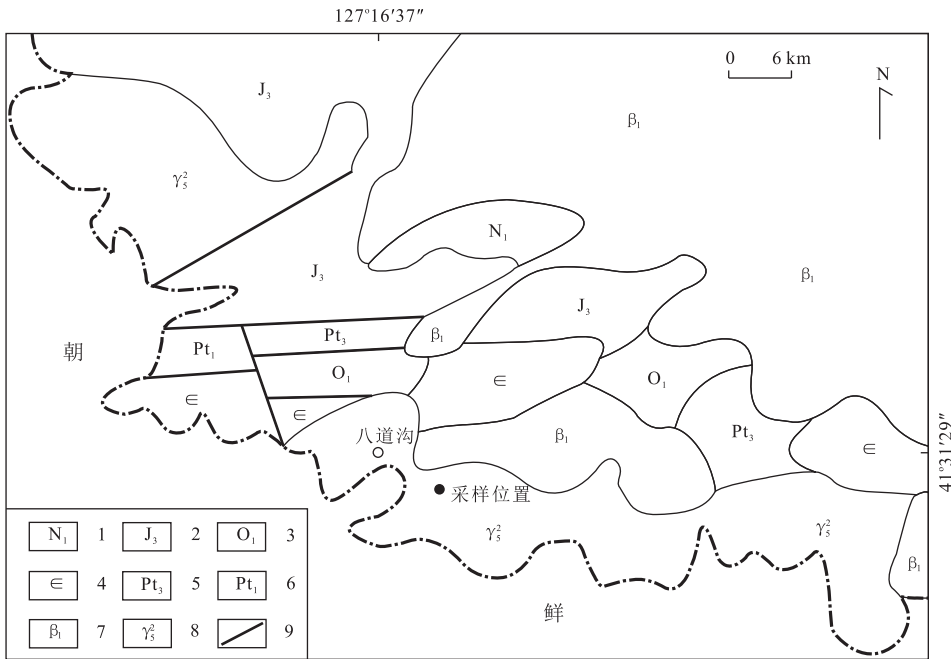


图 2 研究区地质简图

Fig.2 Geological map of the study area

- 1.新近系船底山组;2.侏罗系;3.奥陶系;4.寒武系;5.青白口系钓鱼台组;6.老岭群花山组;7.喜马拉雅期玄武岩;8.侏罗纪花岗岩;9.断裂及节理

组玄武岩(李东津等,1997)(图 2)。其中老岭群花山组分布在八道沟西北部,岩性为以千枚岩为主的泥质粉砂质浅变质岩系;青白口系钓鱼台组则出露在八道沟东部及北部,岩性为以石英砂岩为主的碎屑沉积建造;古生界寒武系及奥陶系交替出露在研究区北缘,前者为碳酸盐-碎屑岩沉积建造,后者则为以灰岩为主的碳酸盐沉积建造。

八道沟地区岩浆岩分布广泛,岩体出露面积较大且总体呈南东向展布。该岩体北部大部分被新生代喜马拉雅期玄武岩覆盖,其余部分与老岭群花山组、青白口系钓鱼台组、奥陶系及寒武系呈侵入接触,花岗岩与灰岩、大理岩接触处形成规模较小且形

态复杂的矽卡岩。区内构造发育,主要以近东西走向的节理及脆性断裂的形式出现,东西向构造形成较早,多被后期北东、北西向断裂错开。北西、北东及近南北向线性构造与东西向断裂带共同切割元古代及古生代地层,形成了一套角砾化、碎裂化岩石组合,同时也控制了该区矿脉在空间上的产出与分布(吉林省地质矿产局,1988;孙永杰等,2007)。

2 样品描述

岩石样品为二长花岗岩,具花岗结构、块状构造,主要矿物为斜长石、碱性长石、石英、角闪石及黑

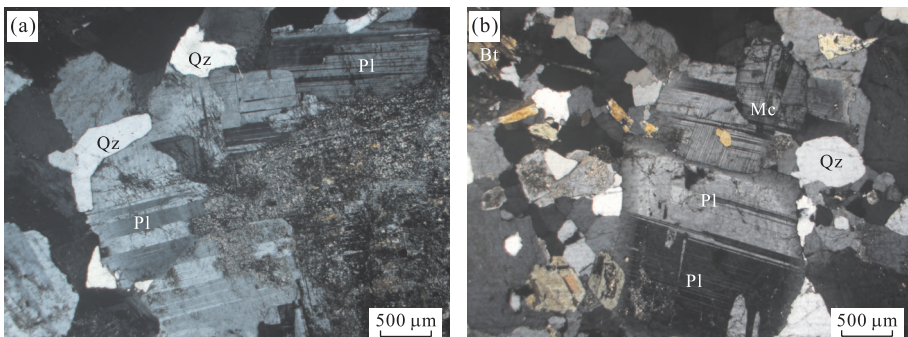


图 3 八道沟二长花岗岩显微照片

Fig.3 Microscopic photographs from Badaogou monzonitic granites

- Bt.黑云母;Mc.微斜长石;Pl.斜长石;Qz.石英

云母,副矿物主要为磁铁矿、锆石、绿帘石、榍石和磷灰石等。其中斜长石含量约为 35%,自形一半自形,粒度为 0.5~2.0 mm,发育聚片双晶,表面有弱绢云母化;碱性长石包括微斜长石、条纹长石,含量约为 30%,粒度为 1~2 mm;石英含量约为 25%,粒度为 0.2~1.0 mm,呈不规则粒状,局部交代长石;角闪石含量约为 5%,呈长柱状,长轴为 1~2 mm;黑云母含量小于 5%,呈片状,少量绿泥石化(图 3)。

3 分析方法

3.1 锆石 LA-ICP-MS 年代学

锆石单矿物挑选是由河北省廊坊市区域地质调查研究所实验室完成,利用标准重矿物分离技术,在双目镜下挑选出晶形较好且无明显裂痕的锆石单矿物颗粒,将不同矿物固定在环氧树脂表面,打磨抛光后露出锆石的表面,最后进行透射光、反射光及阴极发光(CL)图像的采集。锆石激光剥蚀等离子体质谱(LA-ICP-MS)U-Pb 同位素分析则在吉林大学东北亚矿产资源评价国土资源部重点实验室完成。激光剥蚀仪器使用德国相干公司(Coherent)COMPEX Pro 型 ArF 准分子激光器,质谱仪使用美国安捷伦公司 7500A 型四极杆等离子质谱。本次测试采用的激光剥蚀束斑直径为 32 μm ,能量密度为 10 J/cm²,剥蚀频率为 8 Hz。实验中使用高纯度氦气(流速为 600 mL/min)作为载气流,辅助气为氩气(气流量为 1.15 L/min)。采用国际标准锆石 91500(1 062 Ma)进行同位素比值的外部校正,标准锆石 PLE/GJ-1/Qing Hu 为监控盲样。锆石元素含量以国际标样 NIST610 为外标、Si 为内部标准计算,NIST612 和 NIST614 为监控盲样。样品同位素比值及元素含量计算由 ICPMSDataCal(Liu *et al.*, 2008, 2010)测定,年龄计算及绘制图件采用国际标准程序 Isoplot(Ludwig, 2003),普通铅校正则使用 Anderson(2002)给出的程序计算。

3.2 岩石地球化学

本次样品主、微量元素均在吉林大学测试实验中心完成。主量元素由 X 射线荧光光谱仪(PW1401/10)测定(GB/T14506.28-93),相对标准偏差为 2%~5%;微量及稀土元素则由美国安捷伦科技有限公司 Agilent 7500A 型耦合等离子体质谱仪测定(Z/T0223-2001),本次样品测试经国际标样 BHVO-2、BCR-2 和国家标样 GBW07103、GBW07104 监控,元素含量大于 10×10^{-6} 的分析精度小于 5%,元素含量小于 $10 \times$

10^{-6} 的分析精度小于 10%。

3.3 锆石 Hf 同位素测试

本次样品中锆石 Lu-Hf 同位素测定在中国地质科学院矿产资源研究所国土资源部成矿作用与资源评价重点实验室完成,实验所用仪器及系统分别采用 Finnigan Neptune 多接收电感耦合等离子质谱仪及 New Wave UP213 激光剥蚀系统,具体分析流程及仪器操作参见侯可军等(2007)。分析过程中采用锆石标样 GJ1 作为参考物,其¹⁷⁶Hf/¹⁷⁷Hf 加权平均值为 $0.282\ 015 \pm 0.000\ 028(2\sigma, n=10)$,与文献报道值(Elhrou *et al.*, 2006;侯可军等,2007)在误差范围内一致。

4 分析结果

4.1 锆石 LA-ICP-MS 年代学

岩石样品中的锆石主要为长柱状,自形程度较好,具有较清晰的韵律生长环带结构(图 4),且其 Th/U 比值为 0.22~0.62(表 1),具岩浆成因锆石特征(Weaver, 1991)。锆石分析点均落在谐和线上及其附近,²⁰⁶Pb/²³⁸U 的加权平均年龄为 176 ± 1 Ma, MSWD=0.92(图 5),代表花岗岩的侵位年龄属早侏罗世。

4.2 地球化学特征

4.2.1 主量元素 八道沟花岗岩主量元素分析结果见表 2,其中 SiO₂ 含量为 68.93%~70.02%,全碱含量(K₂O+Na₂O)为 7.69%~7.99%,MgO 含量为 1.27%~1.45%,CaO 含量为 1.80%~2.78%,Fe₂O₃^T 含量为 3.05%~3.32%,Mg[#] 值介于 45~47,铝饱和

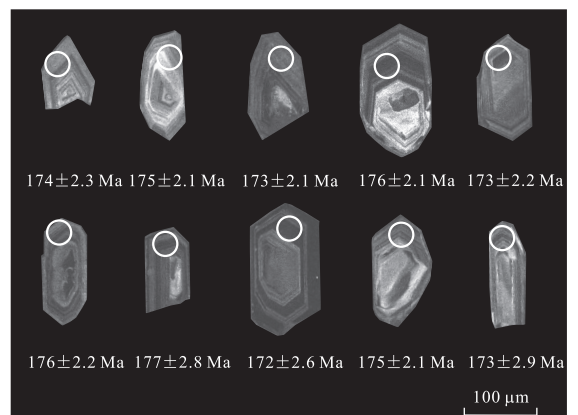


图 4 八道沟二长花岗岩部分锆石阴极发光(CL)图像
Fig. 4 Cathodoluminescence (CL) images of zircons from Badaogou monzonitic granites

表 1 八道沟二长花岗岩锆石 LA-ICP-MS U-Pb 同位素分析结果
Table 1 Zircon LA-ICP-MS U-Pb isotope analysis results of Badaogou monzonitic granites

测点号	含量(10 ⁻⁶)		Th/U	同位素比值及误差				年龄(Ma)及误差				钛地质温度计算						
	Th	U		²⁰⁷ Pb/ ²⁰⁶ Pb	1σ	²⁰⁷ Pb/ ²³⁵ U	1σ	²⁰⁶ Pb/ ²³⁸ U	1σ	²⁰⁷ Pb/ ²³⁵ U	1σ	²⁰⁶ Pb/ ²³⁸ U	1σ	T(10 ⁻⁶)	最高温度(°C)	最低温度(°C)		
BDG-NI-1	441	851	0.52	0.055 2	0.004 4	0.207 0	0.016 4	0.027 2	0.000 5	433	180	191	14	173	3	15.8	789	776
BDG-NI-3	1 455	3 395	0.43	0.053 0	0.002 1	0.199 0	0.008 2	0.027 1	0.000 4	328	91	184	7	172	2	63.7	942	928
BDG-NI-5	1 306	2 307	0.57	0.052 6	0.002 6	0.203 6	0.010 2	0.028 0	0.000 4	322	119	188	9	178	3	12.0	763	751
BDG-NI-6	893	3 381	0.26	0.053 1	0.002 2	0.200 6	0.007 9	0.027 4	0.000 4	332	99	186	7	174	2	7.97	727	716
BDG-NI-7	1 545	4 679	0.33	0.053 9	0.002 0	0.203 1	0.007 5	0.027 2	0.000 4	369	85	188	6	173	2	8.41	732	720
BDG-NI-9	1 049	3 394	0.31	0.055 0	0.002 3	0.210 4	0.009 0	0.027 6	0.000 3	409	91	194	8	175	2	19.4	809	797
BDG-NI-10	1 012	1 906	0.53	0.053 1	0.002 6	0.204 0	0.011 0	0.027 7	0.000 5	332	113	189	9	176	3	8.40	732	720
BDG-NI-14	890	4 035	0.22	0.051 1	0.001 9	0.195 5	0.007 3	0.027 6	0.000 3	256	81	181	6	176	2	12.8	769	757
BDG-NI-15	1 108	4 354	0.25	0.051 4	0.001 8	0.199 8	0.007 3	0.028 1	0.000 4	261	81	185	6	178	2	15.8	789	777
BDG-NI-16	883	2 916	0.30	0.050 0	0.002 0	0.189 9	0.007 4	0.027 6	0.000 3	195	94	177	6	175	2	6.36	709	697
BDG-NI-17	1 103	3 553	0.31	0.050 9	0.001 9	0.191 1	0.007 2	0.027 2	0.000 3	235	89	178	6	173	2	15.6	788	776
BDG-NI-18	425	1 865	0.23	0.051 0	0.002 3	0.196 5	0.008 8	0.028 0	0.000 4	243	106	182	7	178	2	1.52	604	593
BDG-NI-19	973	2 169	0.45	0.053 9	0.002 8	0.208 3	0.011 2	0.028 1	0.000 5	369	120	192	9	179	3	7.62	723	712
BDG-NI-20	3 841	8 638	0.44	0.057 2	0.001 8	0.223 9	0.007 4	0.028 2	0.000 4	502	70	205	6	179	3	49.5	911	897
BDG-NI-22	853	2 938	0.29	0.048 3	0.001 9	0.184 6	0.007 3	0.027 6	0.000 4	117	94	172	6	176	2	17.7	800	787
BDG-NI-23	2 016	3 259	0.62	0.047 1	0.001 8	0.185 7	0.007 8	0.028 2	0.000 5	53.8	89	173	7	179	3	24.8	834	821
BDG-NI-24	1 147	4 061	0.28	0.051 1	0.002 5	0.194 9	0.008 4	0.027 9	0.000 5	256	111	181	7	177	3	13.2	772	760
BDG-NI-25	1 055	3 096	0.34	0.048 4	0.002 1	0.180 7	0.007 4	0.027 1	0.000 4	120	102	169	6	172	3	6.76	713	702

注:常用的锆石 Ti 地质温度计算公式为 $T(^{\circ}\text{C}) = (5.080 \pm 30) / [(6.01 \pm 0.03) - \lg(\text{Ti})] - 273$, 据 Waston *et al.* (2006).

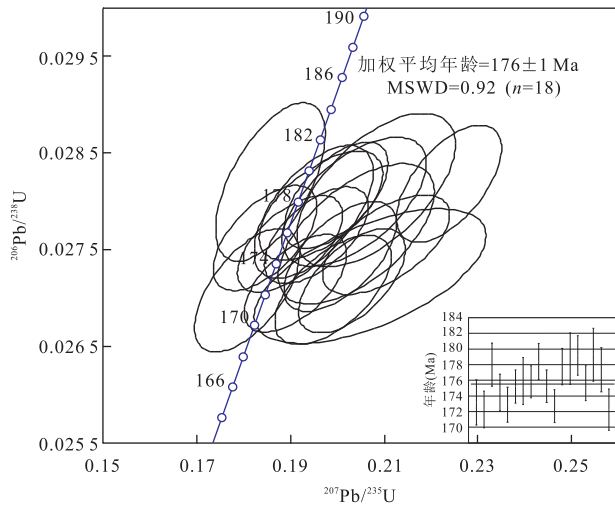


图 5 八道沟二长花岗岩 U-Pb 年龄谐和图

Fig.5 U-Pb concordia diagram for Badaogou monzonitic granites

表 2 八道沟二长花岗岩主量元素 (%)、微量元素 (10^{-6}) 和稀土元素 (10^{-6}) 含量

Table 2 Major elements (%), trace elements (10^{-6}) and REE (10^{-6}) contents of Badaogou monzonitic granites

样品	BDG-B1	BDG-B2	BDG-B3	BDG-B4	BDG-B5
SiO ₂	69.19	70.02	69.48	69.51	68.93
TiO ₂	0.31	0.31	0.30	0.28	0.30
Al ₂ O ₃	14.15	14.04	14.57	14.15	14.32
Fe ₂ O ₃ ^T	3.25	3.14	3.05	3.07	3.32
MnO	0.12	0.10	0.11	0.11	0.12
MgO	1.36	1.41	1.27	1.33	1.45
CaO	2.74	1.80	2.40	2.78	2.53
Na ₂ O	3.44	3.31	3.61	3.47	3.53
K ₂ O	4.25	4.68	4.32	4.28	4.41
P ₂ O ₅	0.10	0.09	0.10	0.09	0.09
LOI	0.91	0.82	0.63	0.50	0.89
Total	99.81	99.71	99.83	99.57	99.90
ALK	7.69	7.99	7.92	7.74	7.94
Na ₂ O/K ₂ O	0.81	0.71	0.84	0.81	0.80
A/CNK	0.93	1.02	0.97	0.92	0.94
Mg [#]	45	47	45	46	46
La	20.7	22.2	20.7	13.9	15.8
Ce	41.6	41.1	37.7	25.9	30.4
Pr	4.81	5.02	4.71	3.51	3.88
Nd	16.9	17.6	16.9	12.6	14.3
Sm	3.37	3.34	3.40	2.79	3.15
Eu	1.00	1.01	1.04	0.90	0.97
Gd	2.85	2.88	2.85	2.41	2.86
Tb	0.45	0.42	0.45	0.39	0.44
Dy	2.31	2.27	2.36	2.11	2.43
Ho	0.51	0.46	0.49	0.44	0.49
Er	1.41	1.34	1.41	1.29	1.42
Tm	0.23	0.21	0.22	0.20	0.22
Yb	1.51	1.37	1.41	1.35	1.53
Lu	0.25	0.21	0.22	0.22	0.24
Y	11.4	10.6	11.3	10.3	11.5

续表 2

样品	BDG-B1	BDG-B2	BDG-B3	BDG-B4	BDG-B5
ΣREE	97.9	99.5	93.8	68.0	78.1
LREE	88.4	90.4	84.4	59.6	68.5
HREE	9.51	9.15	9.43	8.41	9.63
LREE/HREE	9.29	9.87	8.95	7.09	7.11
(La/Yb) _N	9.82	11.7	10.5	7.38	7.41
δEu	0.96	0.97	0.99	1.03	0.96
δCe	0.99	0.92	0.90	0.89	0.93
Rb	137	148	149	86.3	110
Ba	550	643	611	488	530
Th	12.8	11.5	12.4	8.95	9.37
U	3.07	2.42	2.12	2.00	2.02
Nb	12.0	11.4	12.7	10.4	11.3
Ta	1.63	1.54	1.62	1.50	1.44
Sr	370	317	341	304	356
Zr	50.2	39.2	42.5	95.4	47.2
Hf	2.58	2.19	2.32	3.87	2.48
Nb/Ta	7.35	7.38	7.85	6.94	7.81
Rb/Sr	0.37	0.47	0.44	0.28	0.31
La/Nb	1.73	1.95	1.62	1.33	1.40
Ba/Nb	46.0	56.5	48.0	46.9	47.1

注: ALK=摩尔(Na₂O+K₂O), A/CNK=摩尔 Al₂O₃/(CaO+Na₂O+K₂O); Na₂O/K₂O 为质量分数比值; Mg[#] = 100 × (MgO/40.31)/(MgO/40.31+Fe₂O₃^T × 2/159.7).

指数 A/CNK 为 0.92~1.02. 岩石样品在 SiO₂-K₂O 图解上均落在高钾钙碱性系列(图 6a), 在 A/CNK-A/NK 图解中均落在准铝质范围内(图 6b).

4.2.2 微量元素 八道沟花岗岩稀土总量 ΣREE 为 78.1 × 10⁻⁶ ~ 99.5 × 10⁻⁶, 在稀土元素配分模式图上曲线基本一致(图 7a). (La/Yb)_N 为 7.38 ~ 11.7, LREE/HREE 为 7.09~9.87, 配分曲线为明显的右倾型, 强烈富集轻稀土元素(LREE)且分馏明显, 重稀土元素(HREE)相对亏损; δEu 为 0.96 ~ 1.03, 铕异常不明显. 图 7b 显示, 微量元素相对富集 K、Ba、Rb 等大离子亲石元素和 Th、U 等活泼的不相容元素, 亏损 Nb、Ta、P、Zr、Ti 等高场强元素.

4.3 锆石 Lu-Hf 同位素

八道沟花岗岩锆石的 ¹⁷⁶Hf/¹⁷⁷Hf 比值为 0.281 989~0.282 098, ε_{Hf}(t) 值为 -20.21 ~ -24.08, Hf 同位素单阶段模式年龄(t_{DM1})和二阶段模式年龄(t_{DM2})分别为 1 672~1 826 Ma 和 2 498~2 740 Ma(表 3). 此外, 由于测试点的 f_{Lu/Hf} 值为 0.94~0.96, 明显低于铁镁质地壳 f_{Lu/Hf} 值(-0.34)和硅铝质地壳 f_{Lu/Hf} 值(-0.72)(Vervoort et al., 1996; Amelin et al., 2000), 因此二阶段模式年龄更能反映其源区物质从亏损地幔抽取或在地壳中的平均存留年龄(刘跃等, 2014).

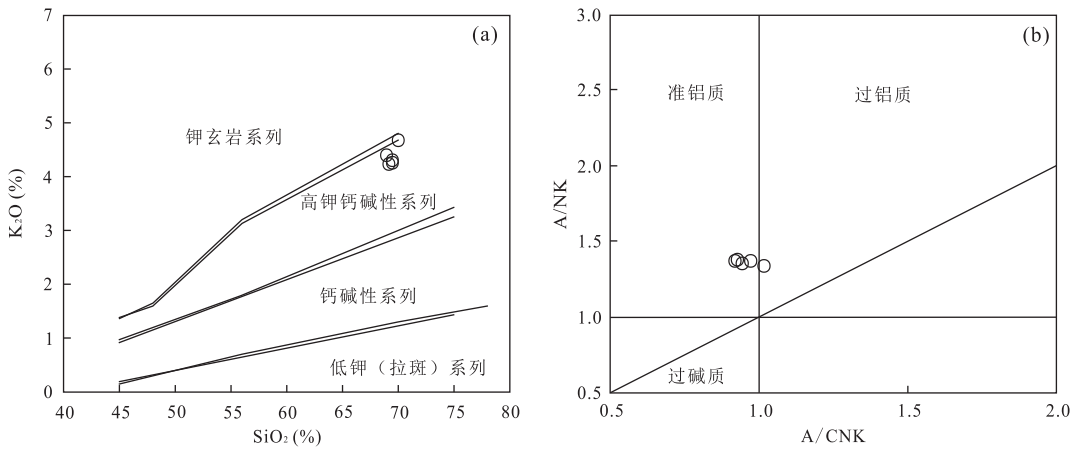


图 6 八道沟二长花岗岩 $\text{SiO}_2\text{-K}_2\text{O}$ 图解和 A/CNK-A/NK 图解

Fig.6 SiO_2 vs. K_2O diagram (a) and A/CNK vs. A/NK diagram (b) of Badaogou monzonitic granites

表 3 八道沟二长花岗岩锆石 Hf 同位素分析结果

Table 3 Zircon Hf isotopic compositions of Badaogou monzonitic granites

样品	t (Ma)	$^{176}\text{Yb}/^{177}\text{Hf}$	$^{176}\text{Lu}/^{177}\text{Hf}$	$^{176}\text{Hf}/^{177}\text{Hf}$	2σ	$\epsilon_{\text{Hf}}(0)$	$\epsilon_{\text{Hf}}(t)$	2σ	$t_{\text{DM1}}(\text{Hf})$	$t_{\text{DM2}}(\text{Hf})$	$f_{\text{Lu/Hf}}$
BDG-N1-1	176	0.040 955	0.001 710	0.282 023	0.000 014	-26.49	-22.83	0.50	1 764	2 662	-0.95
BDG-N1-2	176	0.041 711	0.001 770	0.282 044	0.000 015	-25.73	-22.08	0.54	1 736	2 615	-0.95
BDG-N1-3	176	0.037 639	0.001 569	0.282 039	0.000 013	-25.92	-22.25	0.45	1 735	2 626	-0.95
BDG-N1-4	176	0.031 808	0.001 396	0.282 010	0.000 014	-26.96	-23.27	0.49	1 768	2 690	-0.96
BDG-N1-5	176	0.032 890	0.001 437	0.282 018	0.000 017	-26.67	-22.98	0.59	1 758	2 671	-0.96
BDG-N1-6	176	0.034 119	0.001 424	0.282 059	0.000 016	-25.22	-21.53	0.56	1 700	2 581	-0.96
BDG-N1-7	176	0.038 575	0.001 623	0.282 036	0.000 013	-26.02	-22.35	0.45	1 741	2 632	-0.95
BDG-N1-8	176	0.030 998	0.001 291	0.282 036	0.000 016	-26.04	-22.34	0.58	1 727	2 631	-0.96
BDG-N1-9	176	0.048 996	0.001 986	0.281 989	0.000 017	-27.71	-24.08	0.61	1 826	2 740	-0.94
BDG-N1-10	176	0.041 459	0.001 751	0.282 089	0.000 015	-24.14	-20.49	0.54	1 672	2 515	-0.95
BDG-N1-11	176	0.047 229	0.001 990	0.282 050	0.000 015	-25.54	-21.92	0.52	1 739	2 605	-0.94
BDG-N1-12	176	0.028 037	0.001 170	0.282 006	0.000 016	-27.07	-23.35	0.57	1 762	2 695	-0.96
BDG-N1-14	176	0.052 130	0.002 129	0.282 098	0.000 016	-23.82	-20.21	0.56	1 676	2 498	-0.94
BDG-N1-15	176	0.051 916	0.002 079	0.282 056	0.000 014	-25.33	-21.72	0.51	1 735	2 592	-0.94

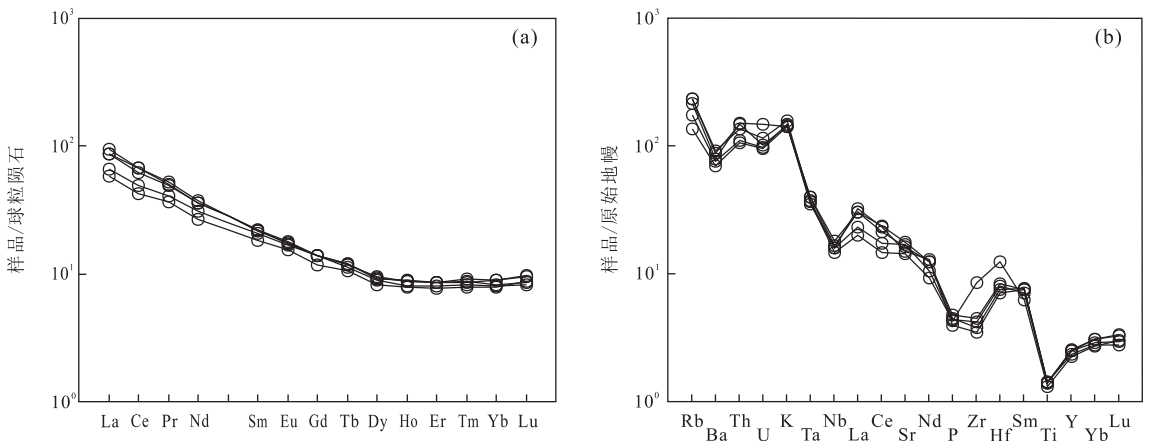


图 7 八道沟二长花岗岩稀土元素球粒陨石标准化配分模式(a)和微量元素原始地幔标准化蛛网图(b)

Fig.7 Chondrite-normalized REE pattern (a) and PM-normalized trace element spider diagram (b) of Badaogou monzonitic granites
球粒陨石标准化值据 Boynton(1984), 原始地幔标准化值据 Sun and McDonough (1989)

5 讨论

5.1 岩浆源区

八道沟花岗岩 SiO_2 含量为 68.93%~70.02%, Al_2O_3 含量为 14.04%~14.57%, 贫 MgO (1.27%~1.45%)、 MnO (1.11~1.12) 和 CaO (1.80%~2.78%), 全碱含量为 7.69%~7.99%, $\text{Na}_2\text{O}/\text{K}_2\text{O}$ 比值为 0.71~0.84, 指示其来源于陆壳(张旗等, 2008). 富集 Rb、K 等大离子亲石元素和轻稀土元素, 亏损高场强元素 Nb、Ta、P、Ti, 与下地壳岩石微量元素配分模式一致(Rudnick *et al.*, 1995). 岩石样品的 Rb/Sr 值为 0.28~0.47, 与地壳 Rb/Sr 标准值 0.35 接近(上地幔值为 0.034)(Taylor and McLennan, 1995); Nb/Ta 值为 6.94~7.85, 与地壳 Nb/Ta 标准值 8.30 接近(地幔平均值为 17.5)(Sun and McDonough, 1989); 此外 La/Nb 值为 1.33~1.95 (平均值 1.61), Ba/Nb 值为 46.0~56.5 (平均值 48.9), 两者均与地壳标准值相近(La/Nb 地壳、地幔标准值分别为 2.20、0.94; Ba/Nb 地壳、地幔标准值分别为 54.0、9.0)(Weaver, 1991); 结合 $\text{Mg}^\#$ 值介于 45~47, 暗示其岩浆可能源于下地壳部分熔融(Bea *et al.*, 2001). 岩石中 Cr、Ni、Co、Sc 和 V 含量分别为 21.4×10^{-6} ~ 50.3×10^{-6} 、 6.3×10^{-6} ~ 14.2×10^{-6} 、 6.63×10^{-6} ~ 7.78×10^{-6} 、 4.52×10^{-6} ~ 5.53×10^{-6} 和 37.6×10^{-6} ~ 42.6×10^{-6} , 与地幔值的对比结果指示岩石未与地幔发生反应(Atherton and Petford, 1993; 邓晋福等, 1999). 钨异常不明显($\delta\text{Eu} = 0.96 \sim 1.03$), 贫 Yb ($1.35 \times 10^{-6} \sim 1.53 \times 10^{-6}$) 和 Y ($10.3 \times 10^{-6} \sim 11.5 \times 10^{-6}$), 表明岩浆源区石榴石稳定存在(张旗等, 2006), 结合 Sr 含量 ($304 \times 10^{-6} \sim 370 \times 10^{-6}$) 笔者认为源区可能不稳定存在斜长石. 此外, 大量前人研究资料显示东北地区侏罗纪期间可能存在地壳加厚的现象(刘燊等, 2009; 张超, 2014; 杨凤超等, 2015), 因此笔者推测岩石应为加厚下地壳部分熔融产物.

锆石 Hf 同位素也是示踪岩浆源区的重要手段(Griffin *et al.*, 2000; Bouvier *et al.*, 2008; 李碧乐等, 2016). 研究区锆石的 $\epsilon_{\text{Hf}}(t)$ 值为 $-20.2 \sim -24.1$, 图 8 表明岩石为壳源(Yang *et al.*, 2006; 吴福元等, 2007a), 此外锆石 Hf 同位素成分变化范围较小, 说明岩石的源区单一. 结合二阶段模式年龄 $t_{\text{DM2}}(\text{Hf})$ 为 2 498~2 740 Ma 及区域动力学演化背景(李雪梅, 2009), 笔者认为八道沟花岗岩属新太古代下地壳部分熔融的产物.

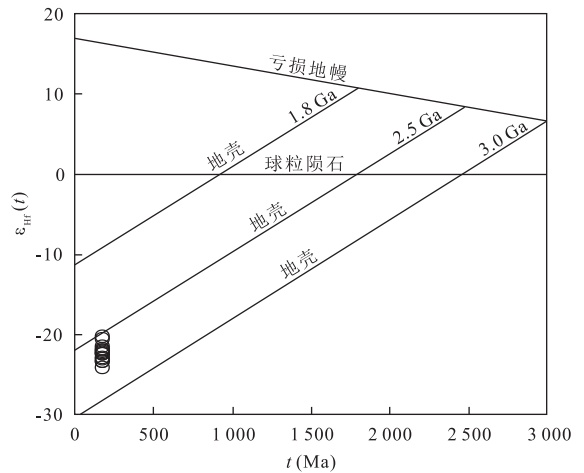


图 8 八道沟二长花岗岩 Hf 同位素特征

Fig.8 Hf isotopic compositions of Badaogou monzonitic granites
据 Yang *et al.* (2006)

5.2 构造背景及成因

八道沟花岗岩富硅、碱, 铝含量中等, 贫 MgO 、 CaO 、 $\text{Fe}_2\text{O}_3^{\text{T}}$, 轻稀土元素明显富集, 钨异常不明显, A/CNK 值均小于 1.1 (0.92~1.02), A/NK 值均大于 1.0 (1.35~1.38), 属准铝质高钾钙碱性系列, 显示为 I 型花岗岩特征. 此外, 岩石样品锆石 U-Pb 测年结果显示其加权平均年龄为 176 ± 1 Ma, MSWD=0.92, 为早侏罗世产物; 而对研究区附近出露的花岗岩岩体前人已进行过研究, 如孙德有等(2005)利用 CHIME 定年法测定梨树沟 (173 ± 10 Ma)、草山 (175 ± 6 Ma) 岩体年龄, 秦亚等(2013)利用锆石 U-Pb 测年法测得荒沟山地区老秃顶子 (179 ± 1 Ma) 及草山岩体 (181 ± 1 Ma) 年龄, 结果显示其均与本次研究样品属同一时代的岩浆产物. 大量研究表明, 沿华北板块东部至张广才岭地区分布着一期呈东北向展布岩性相似的 I 型侏罗纪花岗岩(图 1), 其年龄从东到西呈由老变新的趋势(张炯飞和祝洪臣, 2000; Wu *et al.*, 2005; 隋振民, 2007; 杨进辉等, 2007; 裴福平, 2008; 张娟, 2011; 张超, 2014); 结合八道沟花岗岩及周围岩体地球化学数据及年代学分析结果, 笔者判断其均为该期侏罗纪花岗岩的一部分.

关于古太平洋板块俯冲开始时间一直为研究区内争议的热点问题, 主要有以下两种观点, 一种认为是晚三叠世(彭玉鲸和陈跃军, 2007), 另一种则认为是早一中侏罗世(Xu *et al.*, 2009; Wu *et al.*, 2011). 有学者研究认为华北东部并不存在与太平洋板块边界相平行的三叠纪花岗岩带(Li and Li, 2007), 而东北地区晚三叠世火山岩呈面型展布, 两者均无法用

太平洋板块俯冲模式来解释;此外,大洋板块俯冲最直接的标志是岩浆作用和陆缘增生杂岩及其沉积作用,而东北亚东部大陆边缘未发现晚三叠世钙碱性火山岩及其同期增生岩(唐杰等,2016),区域内发育的晚三叠世海陆交互相沉积地层显示为被动陆缘沉积组合(Zhang *et al.*, 2015),由此笔者推断晚三叠世并不存在古太平洋板块的俯冲作用。而小兴安岭—张广才岭地区发育一系列早—中侏罗世双峰式火山岩组合(唐杰等,2011; Yu *et al.*, 2012; 徐美君等,2013),吉黑东部—朝鲜半岛北端发育同时代类似的陆缘钙碱性火山岩组合(许文良等,2008;唐杰等,2016),结合该期花岗岩整体宏观趋势及呈带状展布的特点,笔者推断其形成与古太平洋板块在早—中侏罗世处于俯冲背景下有关。而刘万臻(2014)对福安堡钼矿进行的研究表明其属于细网脉状钼矿床,赋矿围岩属 I 型花岗岩(形成时间分别为 179 ± 2 Ma, 172 ± 1 Ma),成矿时代为中侏罗世,指示该时期的成矿作用与环太平洋体系有关。

吴福元等(2007b)对辽东半岛中生代花岗岩进行锆石饱和温度计算发现,侏罗纪花岗岩形成温度约为 750°C ,而白垩纪花岗岩形成温度为 $799 \sim 865^\circ\text{C}$ 。前者属低温花岗岩,其形成可能与流体加入有关,并很可能反映一种俯冲的构造背景;而后者属高温花岗岩,其热能的产生可能与俯冲作用后期岩石圈拆沉导致软流圈上涌带来大量热能有关。笔者利用锆石 Ti 地质温度计算公式(Watson *et al.*, 2006)对八道沟花岗岩锆石 18 个有效点的 Ti 含量进行计算,得到其最低温度平均值为 760°C ,最高温度平均值为 773°C ,属低温花岗岩。Xia *et al.*(2013)利用傅里叶变换红外光谱(FTIR)对华北板块费县玄武岩单斜辉石斑晶的结构水含量进行分析,结果显示其初始熔体水含量高于洋中脊玄武岩、洋岛玄武岩和弧后盆地玄武岩(Dixon *et al.*, 1988, 1997; Michael, 1988, 1995; Sobolev and Chaussidon, 1996),与岛弧玄武岩水含量相当(Hochstaedter *et al.*, 1990; Danyushevsky *et al.*, 1993),说明中生代玄武岩地幔源区可能富水。综合上述分析,笔者推测研究区花岗岩形成可能有流体加入。洋壳在深俯冲过程中俯冲板片脱水导致地幔部分熔融,是火山弧、弧后及板内岩浆作用的源区(Maruyama *et al.*, 2007; Zhao *et al.*, 2007)。随着俯冲深度的增加,温度和压力逐渐升高,导致俯冲板片及其所携带的沉积物发生变质、脱水形成俯冲带流体,该流体在一定深度范围内直接交代地幔使其熔点降低并发生部分熔

融。研究区八道沟花岗岩的形成则与早—中侏罗世古太平洋板块的俯冲作用导致俯冲带流体交代地幔楔、使其部分熔融形成玄武质岩浆并发生大规模的底侵作用、大量的热能和流体导致下地壳物质部分熔融有关。

6 结论

(1)八道沟二长花岗岩 LA-ICP-MS 锆石 U-Pb 测年结果显示,岩石加权平均年龄为 176 ± 1 Ma, MSWD=0.92,属早侏罗世产物。

(2)八道沟二长花岗岩富硅、碱,贫 MgO、CaO、 $\text{Fe}_2\text{O}_3^{\text{T}}$,轻稀土元素明显富集,铕异常不明显;锆石 $\epsilon_{\text{HF}}(t)$ 值为 $-20.2 \sim -24.1$, t_{DM2} 为 $2498 \sim 2740$ Ma,表明其源于新太古代下地壳部分熔融。

(3)八道沟二长花岗岩形成于古太平洋板块向欧亚大陆俯冲作用过程中,属岩浆弧环境下的产物。

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