

阿尔泰造山带富蕴基性麻粒岩折返过程： 来自裂变径迹热年代学的限定^{*}

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Abstract On the basis of previous studies of SHRIMP U-Pb chronology, geochemistry, metamorphic temperature/pressure conditions and tectonic setting, fission-track thermochronology has been conducted in this study to determine the exhumation of the basic granulite. Zircon fission-track shows Triassic and apatite fission-track shows Late Cretaceous to Early Cenozoic ages. Thermal history modeling of the fission-track data, combined with previous studies, indicate that the basic granulite was exhumed to the upper crust under about 7.8 km in Triassic. From Late Cretaceous to Early Cenozoic (around 100~50 Ma), the basic granulite, meta-diabase and the wall rock gneiss were cooled to the blocking temperature of apatite fission-track, which was about 3.5 km below the surface. From approximately 50 to 15 Ma, they stayed at the partly-annealed zone of apatite fission-track, about 1.7 km below the surface. Since about 15 Ma to present, they were uplifted and denuded to the surface.

Key words Altai Orogenic Belt; Fuyun; Basic granulite; Fission-track; Exhumation

摘要 论文在阿尔泰造山带富蕴县乌恰沟基性麻粒岩的锆石 SHRIMP 年代学、地球化学、变质温压条件和形成的大地构造背景研究基础上,利用麻粒岩、围岩片麻岩和侵入到麻粒岩的辉绿岩岩墙的裂变径迹热年代学探讨了麻粒岩从深部折返至地表的过程。裂变径迹年代学研究发现基性麻粒岩的锆石裂变径迹年龄为三叠纪,而麻粒岩、围岩片麻岩和侵入到麻粒岩的辉绿岩岩墙的磷灰石裂变径迹年龄均显示为晚白垩世至新生代早期。对磷灰石裂变径迹测试所得到的径迹长度和单颗粒年龄数据进行热史模拟表明,三叠纪时,基性麻粒岩抬升至约地表以下 7.8 km 的上地壳,温度冷却至锆石裂变径迹的封闭温度;晚白垩世至新生代早期(约 100~50 Ma),麻粒岩、围岩片麻岩和辉绿岩抬升至约地表以下 3.5 km,温度冷却至磷灰石裂变径迹的封闭温度;约 50~15 Ma,三者滞留在约地表以下 1.7 km 的磷灰石部分退火带;约 15 Ma 以来,喜马拉雅运动使得它们被抬升剥蚀至地表。

关键词 阿尔泰造山带;富蕴;基性麻粒岩;裂变径迹;折返

中图法分类号 P588.347;P597.3

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1 引言

阿尔泰造山带是中亚地区一条重要的造山带,是加里东-海西期形成和发展的陆缘增生造山带,由不同时期块体拼贴而成(Sengör *et al.*, 1993; Jahn *et al.*, 2000; Xiao *et al.*, 2003, 2004a, b; Yakubchuk, 2004)。阿尔泰造山带以其复杂的构造演化、岩浆活动、变质作用以及丰富的矿产资源,受到国内外地质学家的广泛重视,并开展了大量的研究工作,取得了很多方面的新认识,如造山带演化(Xiao *et al.*, 2004a, 2008a, b; Buslov *et al.*, 2004);地球化学及同位素年代学(Chen *et al.*, 2002; 胡震琴等, 2006; 童英等, 2006; Briggs *et al.*, 2007; Sun *et al.*, 2008; 刘锋等, 2008; 刘国仁等, 2008; 杨富全等, 2008; 柴凤梅等, 2008; Wang *et al.*, 2009);裂变径迹热年代学(Grave and Van, 2002; Yuan *et al.*, 2006; Grave *et al.*, 2007; Vassallo *et al.*, 2007; Buslov *et al.*, 2008)。

麻粒岩,代表着高温变质和源自下地壳-上地幔的产物,与早期地壳演化、碰撞造山、大陆拉张以及岛弧活动等大陆动力学作用密切相关,为了解地球深部物质组成和结构以及造山带的构造演化提供了窗口(郭敬辉等, 1993; Franco *et al.*, 1995a, b; 钱祥麟, 1996; Thompson *et al.*, 1997; 李江海

等, 1998; Schaller *et al.*, 1999; 翟明国和刘文军, 2001; 厉子龙等, 2004a; 陈汉林等, 2006b)。由于麻粒岩记录了如此众多重要的地壳演化和地球动力学信息,因此受到地质学家们的高度关注,成为地球科学的研究热点之一。

从报道阿尔泰造山带富蕴县乌恰沟一带发现基性麻粒岩(Li *et al.*, 2004b)以来,作者对这一下地壳深部变质作用产物相继开展了物质成分、变质作用过程及变质温压条件、锆石 SHRIMP U-Pb 年代学等方面的研究,认为富蕴乌恰沟一带出露的基性麻粒岩的原岩为活动大陆边缘环境下形成的岛弧钙碱性玄武岩,形成时间大概为 268~279 Ma, 其后由于强烈的构造事件,经过麻粒岩相变质作用和角闪岩相变质作用,时间大概为 255 Ma, 从而形成该基性麻粒岩(厉子龙等, 2004a; Li *et al.*, 2004c; Chen *et al.*, 2006a; 陈汉林等, 2006b)。

埋藏于深部的基性麻粒岩何时折返至地表并被剥蚀出露是一个尚未解决且对研究区构造演化具有重要意义的问题。裂变径迹可以记录岩石样品所经历的中低温的热历史演化信息,从而被广泛应用于研究造山带隆升剥露历史。近年来,随着对裂变径迹测试所得出的年龄数据和径迹长度数据的深入研究和分析以及基于磷灰石裂变径迹测试数据的热史模拟研究的不断深入,使得除了裂变径迹表观年龄之外,对岩石样品详细的 T-t 演化研究成为可能(Ketcham *et al.*, 2000)。本文通过对阿尔泰造山带乌恰沟一带出露的基

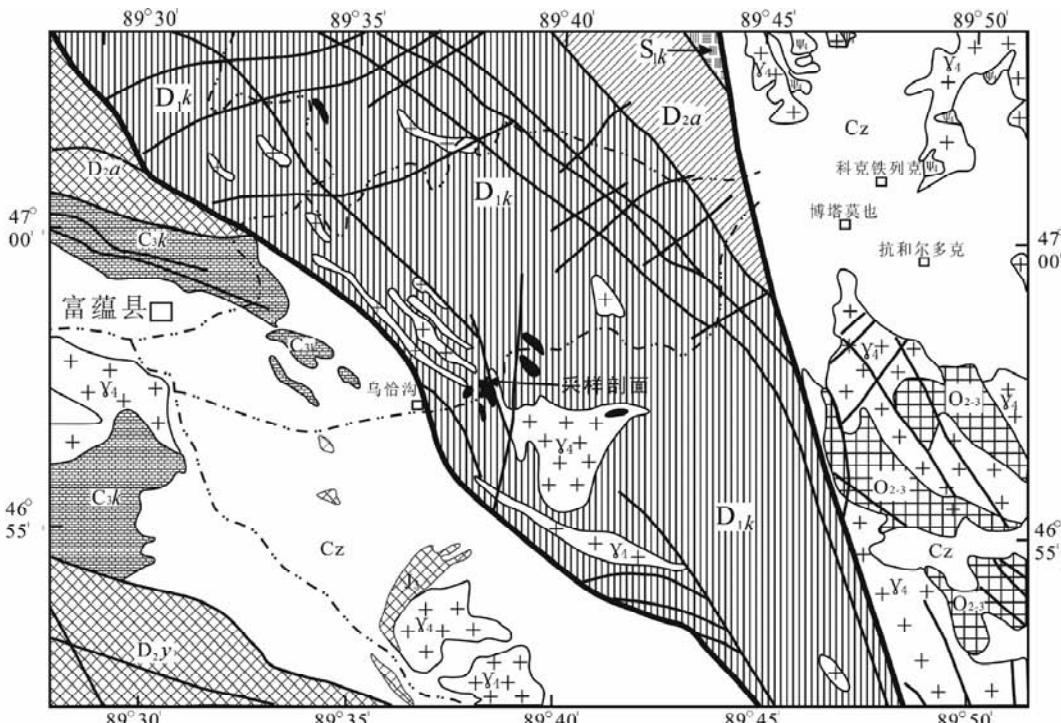


图 1 阿尔泰造山带富蕴地区地质简图及乌恰沟剖面位置(据 Chen *et al.*, 2006a)

O₂₋₃-中-晚奥陶世哈巴河群; S₁k-志留纪库鲁姆提群; D₁k-早泥盆世康布铁堡组; D₂y-中泥盆世蕴都喀拉组; D₂a-中泥盆世阿勒泰组; C₃k-晚石炭世喀喇额尔齐斯组; J₃-晚侏罗世石树沟群; Cz-新生界; Υ₄-华力西期基性、超基性岩; γ₄-华力西期花岗质岩石

Fig. 1 Simplified geological map and the location of the Wuqiaogou section (after Chen *et al.*, 2006a)

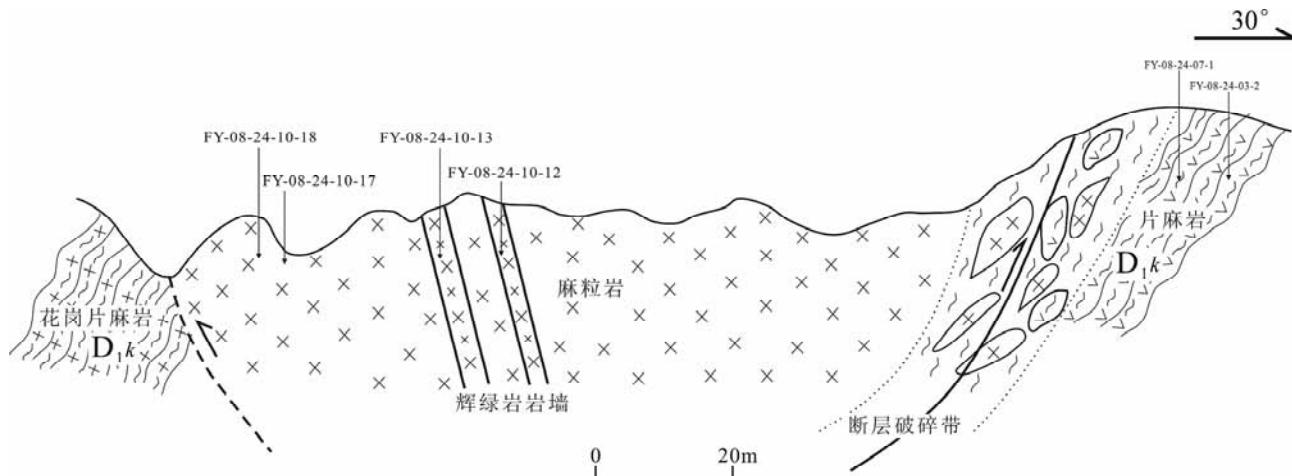


图2 富蕴乌恰沟剖面实测剖面图及裂变径迹样品位置(据Chen et al., 2006a)

Fig. 2 Cross-section and FT sampling locations of Wuqiaogou section in Fuyun (after Chen et al., 2006a)

性麻粒岩、围岩片麻岩和侵入到麻粒岩的辉绿岩岩墙开展裂变径迹研究,探讨该麻粒岩从深部折返至地表的过程。

2 地质概况

本文讨论的阿尔泰基性麻粒岩发现于新疆阿勒泰地区富蕴县城东的乌恰沟一带(图1)。野外调查查明该处麻粒岩出露共约117m,呈透镜状产于早泥盆世康布铁堡组。北东侧麻粒岩与康布铁堡组断层接触(图2),往西南方向,与康布铁堡组呈突变接触关系,推测为断层接触。研究表明,康布铁堡组火山岩成岩年龄为 407.3 ± 9.2 Ma,属于早泥盆世,其变质时间可能为359.5~339.3 Ma(张进红等,2000)。麻粒岩体内见后期辉绿岩脉穿插现象。麻粒岩可进一步分为粗粒和细粒麻粒岩,二者呈渐变过渡。

3 样品采集与实验方法

本论文共采集富蕴县乌恰沟附近新鲜样品六块,其中基性麻粒岩样品两块,分别为FY-08-24-10-17和FY-08-24-10-18;麻粒岩中侵入的辉绿岩样品两块,分别为FY-08-24-10-12和FY-08-24-10-13;麻粒岩周围围岩康布铁堡组片麻岩样品两块,分别为FY-08-24-03-2和FY-08-24-07-1。每个样品重大于2kg。六块样品中,样品FY-08-24-03-2、FY-08-24-07-1、FY-08-24-10-12、FY-08-24-10-13和FY-08-24-10-18五块样品进行磷灰石裂变径迹测试,样品FY-08-24-10-17进行锆石裂变径迹测试。详细的锆石与磷灰石矿物挑选、样品处理、实验流程、试剂选择及计算公式等见袁万明等(2004a, b)和袁万明等(2005),样品测试由中国科学院高能物理研究所袁万明完成,Zeta值为 385 ± 12 。

4 测试结果及分析

五个磷灰石样品的裂变径迹中值年龄为55~76 Ma,每个样品的池年龄与中值年龄基本一致(表1)。样品的裂变径迹年龄远小于成岩/变质年龄,且所有样品的成岩/变质温度皆远大于磷灰石裂变径迹的重置温度,这表明样品在成岩/变质过程中,裂变径迹被重置。五个磷灰石样品中,四个样品的 $P(x^2)$ 检验值远大于5,说明这四个样品的单颗粒磷灰石年龄比较集中,这也从年龄分布直方图和雷达图中得到体现。另一个样品(FY-08-24-10-18)虽然 $P(x^2)$ 检验值小于5,但从单颗粒磷灰石年龄分布直方图与雷达图上看(图3),年龄比较集中,所有单颗粒磷灰石年龄均小于100 Ma,远比样品磷灰石裂变径迹重置的时间年轻,且中值年龄与池年龄十分接近,所以,该样品也不存在未退火完全而存在早期裂变径迹影响的问题,该样品的裂变径迹也反映的是样品变质之后的热事件,而更年轻的热事件的影响较小,故该样品的年龄与其它样品并无太大差别。

五个磷灰石裂变径迹样品的裂变径迹平均径迹长度 $12.1 \sim 13.1 \mu\text{m}$,裂变径迹平均径迹长度较短,反映出样品在经历过度重置之后,在部分退火带(约60~90°C)中滞留时间较长,使得裂变径迹缩短。从裂变径迹长度分布来看,主要呈宽的单峰分布,既有较长的径迹存在,同时也有短径迹存在,这反映出样品中既存在较新生成的径迹,同时又有较老的径迹经部分退火之后长度缩短的现象。长度分布多峰现象不明显,可能预示着样品在径迹重置之后,并没有经历多次的加热事件。

一个锆石裂变径迹样品的裂变径迹中值年龄为 218 ± 14 Ma,池年龄为 214 ± 11 Ma,两者在标准差范围内一致(表2)。 $P(x^2)$ 检验值小于5,且从年龄分布直方图和雷达图来

表 1 阿尔泰造山带富蕴乌恰沟剖面磷灰石裂变径迹测试结果

Table 1 Apatite fission-track results of the samples from Wuqiagou section in Fuyun, Altai Orogenic Belt

样品号 (颗粒数)	岩性	$\rho_s (10^5 \cdot \text{cm}^2)$ (Ns)	$\rho_i (10^5 \cdot \text{cm}^2)$ (Ni)	$\rho_d (10^5 \cdot \text{cm}^2)$ (N)	$P(x^2)$ 检验(%)	中值年龄 $\pm 1\sigma (\text{Ma})$	池年龄 $\pm 1\sigma (\text{Ma})$	平均径迹长度 ± 误差(μm)
FY-08-24-03-2 (23)	片麻岩	4.742 (1004)	2.905 (615)	1.776 (3144)	58	56 ± 4	57 ± 4	13.0 ± 1.7 (92)
FY-08-24-07-1 (28)	片麻岩	4.571 (678)	2.481 (368)	2.050 (3144)	100	72 ± 5	72 ± 5	13.1 ± 2.0 (213)
FY-08-24-10-12 (28)	辉绿岩	3.718 (443)	2.149 (256)	1.670 (3144)	88	55 ± 5	55 ± 5	12.5 ± 1.8 (95)
FY-08-24-10-13 (28)	辉绿岩	5.610 (986)	2.304 (405)	1.628 (3144)	81	76 ± 5	76 ± 5	12.1 ± 1.9 (121)
FY-08-24-10-18 (28)	麻粒岩	12.540 (4440)	6.829 (2418)	2.029 (3144)	0	70 ± 4	71 ± 3	12.2 ± 1.8 (88)

注: ρ_s 为矿物中自发裂变径迹密度, Ns 为自发裂变径迹数; ρ_i 为云母外探测器记录的矿物中诱发裂变径迹密度, Ni 为诱发裂变径迹数; ρ_d 为中子注量监测器标准铀玻璃组件的诱发裂变径迹密度, N 为所测径迹数; $P(x^2)$ 为 Chi-sq 检验概率

表 2 阿尔泰造山带富蕴乌恰沟剖面锆石裂变径迹测试结果

Table 2 Zircon fission-track results of the sample from Wuqiagou section in Fuyun, Altai Orogenic Belt

样品号 (颗粒数)	岩性	$\rho_s (10^5 \cdot \text{cm}^2)$ (Ns)	$\rho_i (10^5 \cdot \text{cm}^2)$ (Ni)	$\rho_d (10^5 \cdot \text{cm}^2)$ (N)	$P(x^2)$ 检验(%)	中值年龄 $\pm 1\sigma (\text{Ma})$	池年龄 $\pm 1\sigma (\text{Ma})$
FY-08-24-10-17 (12)	麻粒岩	105.851 (2414)	81.910 (1868)	25.367 (27782)	0.5	218 ± 14	214 ± 11

注: ρ_s 为矿物中自发裂变径迹密度, Ns 为自发裂变径迹数; ρ_i 为云母外探测器记录的矿物中诱发裂变径迹密度, Ni 为诱发裂变径迹数; ρ_d 为中子注量监测器标准铀玻璃组件的诱发裂变径迹密度, N 为所测径迹数; $P(x^2)$ 为 Chi-sq 检验概率

看(图3),单颗粒锆石裂变径迹年龄比较分散,从300 Ma至150 Ma均有年龄值存在,不但存在经过变质作用而重置的径迹,而且存在混入的未经过重置的径迹。但是单颗粒锆石的径迹年龄主体集中在250~200 Ma,这与袁万明等(2005)报道的阿尔泰造山带的一期构造事件的锆石裂变径迹年龄可以进行对比。

5 磷灰石裂变径迹结果的热史模拟

随着近年来对磷灰石裂变径迹年龄及长度数据研究和分析的不断深入,使得可以利用磷灰石裂变径迹测试所获得的年龄和长度数据进行热史模拟,从而揭示样品所经历和记录的T-t演化。本文对磷灰石裂变径迹测试所获得的单颗粒裂变径迹年龄数据和径迹长度数据进行热史模拟,模拟所采用的模型参见Ketcham *et al.* (2000)。

模拟过程中,现今温度被设定为10~25℃,样品所经历的最高温度设定为100~130℃。由于磷灰石裂变径迹在高于130℃时被重置,所以不能反映完全退火之前的热史信息,故本研究未对样品在130℃之前的热史进行模拟。从沉积地层来看,研究区自晚古生代以来沉积地层仅局限在山间的局部地区,且厚度一共仅几百米,所以本文模拟过程中认为这些采集于造山带内部的样品未经历沉积埋藏加热的过程,而

仅表现为稳定抬升剥蚀冷却。模拟结果列于图4,样品FY-08-24-07-1没有获得满意的模拟结果而未列出。

模拟结果的好坏程度由K-S Test和Age GOF两个参数来衡量,一般两个参数均大于0.05就认为模拟结果可以接受,所给的模型可以得出测试结果;当两个参数都大于0.50时,认为模拟结果好,测试结果支持所给的热史模型。本文模拟结果的两个评价参数均远大于0.50,可以认为磷灰石裂变径迹的测试结果支持热史模拟所给模型。

所得出的四个模拟结果显示出一致的三阶段热史演化过程,第一个阶段为约100~50 Ma,表现为样品自进入重置温度之后的冷却,使样品进入部分退火带;第二阶段为约50~15 Ma,表现为样品长时间稳定滞留在部分退火带;第三阶段为约15 Ma~现今,表现为样品快速冷却至地表温度。

第一阶段(约100~50 Ma),样品温度约从120℃冷却至约60℃,冷却速率为1.2℃/Ma,以30℃/km的地温梯度来计算,样品约抬升2 km,平均抬升速率40 m/Ma。第二阶段(约50~15 Ma),样品长期滞留在部分退火带,这与裂变径迹长度较短的测试结果相对应,径迹在生成之后由于长期滞留部分退火带,而使径迹缩短。第三阶段(约15 Ma~现今),样品温度从约60℃冷却至现今地表温度(约15℃),冷却速率为3℃/Ma,以30℃/km的地温梯度来计,样品约抬升1.5 km,抬升速率100 m/Ma。

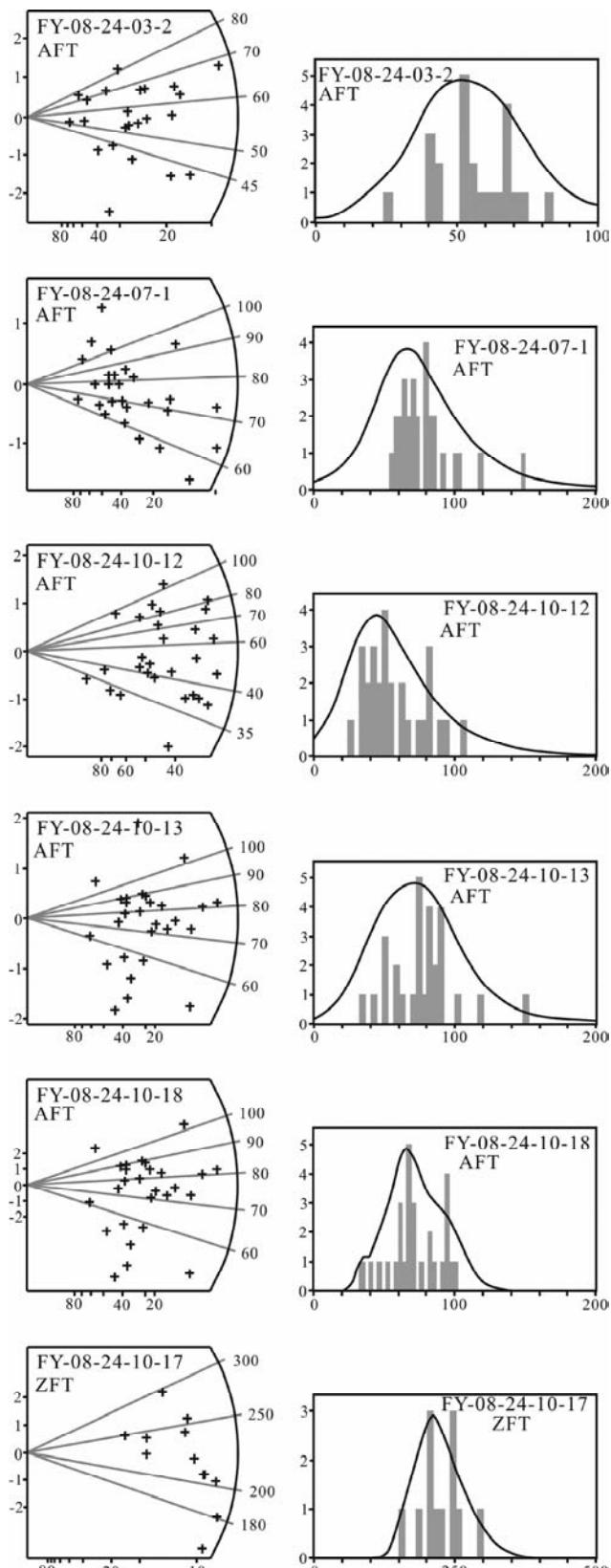


图3 裂变径迹测试单颗粒年龄分布雷达图与直方图

Fig. 3 Radial plots and distributions of single grain ages for the FT samples

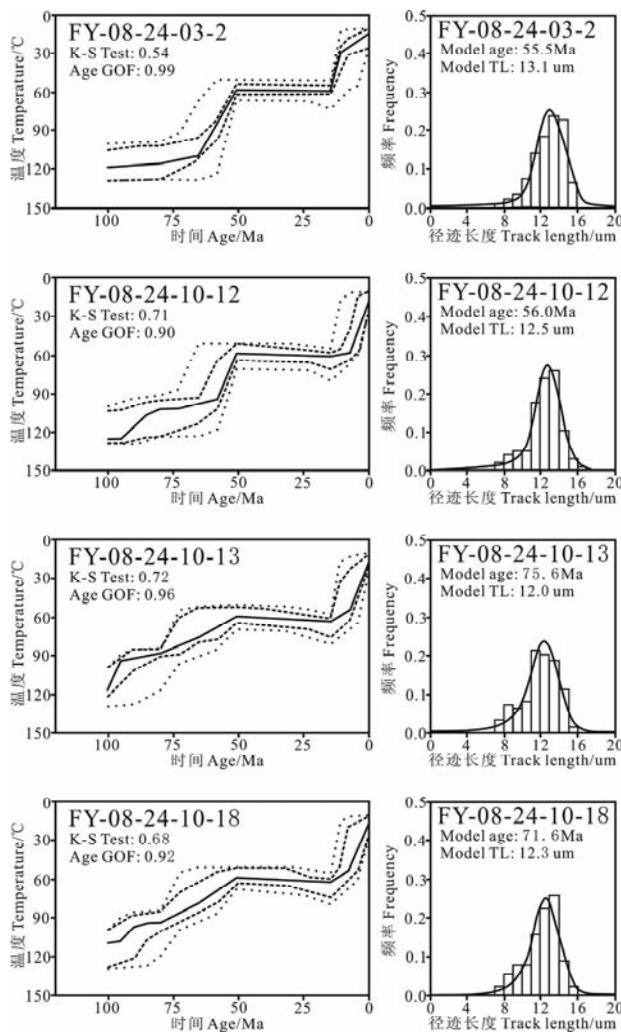


图4 磷灰石裂变径迹数据热史模拟结果

Fig. 4 Thermal history modeling of the AFT data

6 基性麻粒岩的折返过程

作者对富蕴县乌恰沟出露的基性麻粒岩的SHRIMP锆石U-Pb定年,在锆石核部得到268~279 Ma的年龄,而在锆石的变质边部得到255 Ma的年龄(陈汉林等,2006b)。地球化学测试数据显示,基性麻粒岩的原岩为形成于岛弧环境的钙碱性玄武岩(Chen et al., 2006a),麻粒岩形成的温压条件为750~780 °C,6~7 kbar(厉子龙等,2004a; Li et al., 2004c; Chen et al., 2006a),推测钙碱性玄武岩被俯冲至深地壳经受麻粒岩相变质作用。按照地表温度15 °C,地温梯度30 °C/km计算,则基性麻粒岩形成于地表以下约24.5~25.5 km处。

到三叠纪,剧烈的印支造山运动使得本区经历了一次显著的抬升事件,这期抬升运动被麻粒岩中的锆石裂变径迹所记录,出现 218 ± 14 Ma的中值年龄。区域上,三叠纪时期造山运动也被其它地区的锆石裂变径迹所记录。袁万明等

(2005)在阿尔泰造山带的锆石裂变径迹研究中揭示了一期189~243 Ma的造山抬升作用,本文的锆石裂变径迹结果可以与这一结果进行对比。Xiao et al. (2008b)认为晚二叠世古亚洲洋消失并在西伯利亚克拉通与塔里木地块之间形成一个复杂的造山拼贴体。三叠纪时,该造山拼贴体的活动可能使乌恰沟剖面中所露出的基性麻粒岩、其间侵入的辉绿岩岩墙以及围岩康布铁堡组作为冲断片整体经历了共同的抬升。按照锆石裂变径迹开始计时的温度约为250°C、地表温度15°C、地温梯度30°C/km计算,则基性麻粒岩、辉绿岩岩墙以及围岩康布铁堡组在晚三叠世整体抬升至地表以下约7.8 km,已经抬升至上地壳。

燕山运动晚期,本区经历了一次陆内挤压抬升作用,这次事件被基性麻粒岩、辉绿岩岩墙以及围岩康布铁堡组片麻岩的磷灰石裂变径迹所记录。区域上,该期构造运动也被其它地区的裂变径迹所反映(刘顺生等,2002;袁万明等,2004a,b;Yuan et al., 2006)。该期运动在本区表现为基性麻粒岩、辉绿岩岩墙以及围岩康布铁堡组片麻岩整体作为冲断片在晚白垩世至新生代早期(约100~50 Ma)的逆冲抬升,使之进入磷灰石裂变径迹的退火带。按照磷灰石裂变径迹重置温度约120°C、地表温度15°C、地温梯度30°C/km计算,此时麻粒岩被抬升至地表以下约3.5 km。

在新生代很长一段时间内(约50~15 Ma),基性麻粒岩、辉绿岩岩墙以及围岩康布铁堡组片麻岩滞留在磷灰石裂变径迹部分退火带中,使得先前生成的裂变径迹长度缩短。按照热史模拟所得出的古地温约65°C、地表温度15°C、地温梯度30°C/km计算,此时这些岩石滞留于地表以下约1.7 km。

此后,从中新世中期(约15 Ma)开始,受喜马拉雅运动影响,基性麻粒岩、辉绿岩岩墙及围岩康布铁堡组片麻岩经历

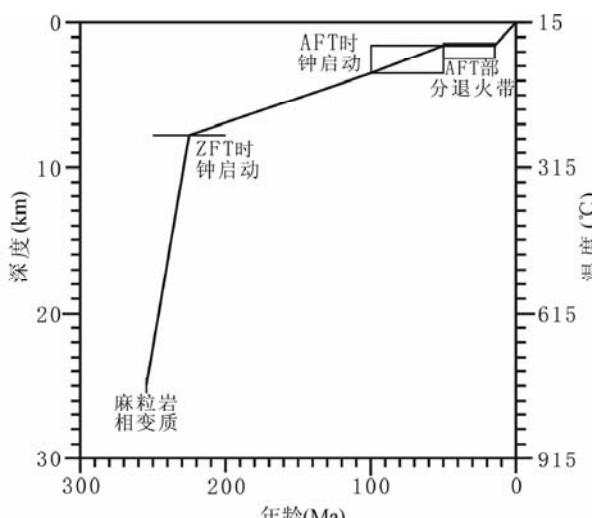


图5 富蕴乌恰沟剖面麻粒岩折返过程(温度-时间与深度-时间演化)

Fig. 5 Exhumation process (T - t and D - t evolution) of the basic granulite of Wuqiagou section in Fuyun

了最新一期的抬升,并被剥蚀至地表。相同的隆升年代也被Grave et al. (2007)所报道。

综合上述认识,本文初步提出如图5所示的阿尔泰造山带富蕴县乌恰沟基性麻粒岩的温度-时间(T - t)和深度-时间(D - t)演化模式。

7 结论

基于阿尔泰造山带富蕴县乌恰沟出露的基性麻粒岩、辉绿岩岩墙及围岩康布铁堡组片麻岩的裂变径迹分析,得出以下初步认识:

(1)三叠纪时,基性麻粒岩、辉绿岩岩墙以及围岩片麻岩共同组成逆冲断片被抬升至约地表以下7.8 km的上地壳,温度冷却至锆石裂变径迹的封闭温度;

(2)晚白垩世至新生代早期(约100~50 Ma),基性麻粒岩、辉绿岩岩墙和围岩片麻岩被抬升至约地表以下3.5 km,温度冷却至磷灰石裂变径迹的封闭温度;

(3)约50~15 Ma,基性麻粒岩、辉绿岩岩墙和围岩片麻岩滞留在约地表以下约1.7 km,温度为磷灰石部分退火带;

(4)约15 Ma以来,喜马拉雅运动使得它们被抬升剥蚀至地表。

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