

中国东北中生代构造体制与区域成矿背景： 来自中生代火山岩组合时空变化的制约*

许文良¹ 王枫¹ 裴福萍¹ 孟恩² 唐杰¹ 徐美君¹ 王伟¹

XU WenLiang¹, WANG Feng¹, PEI FuPing¹, MENG En², TANG Jie¹, XU MeiJun¹ and WANG Wei¹

1. 吉林大学地球科学学院, 长春 130061

2. 中国地质科学院地质研究所, 北京 100037

1. College of Earth Sciences, Jilin University, Changchun 130061, China

2. Institute of Geology, Chinese Academy of Geological Sciences, Beijing 100037, China

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Abstract This paper summarizes geochronology and association of Mesozoic volcanic rocks and their spatial and temporal variations, with the aim of constraining evolutionary history, and ore-forming background of the circum Pacific and Mongol-Okhotsk tectonic systems in NE China. Zircon U-Pb dating results indicate that Mesozoic volcanisms in NE China can be subdivided into six stages, i. e., Late Triassic (200 ~ 228Ma), Early-Middle Jurassic (173 ~ 190Ma), Middle-Late Jurassic (158 ~ 166Ma), early Early Cretaceous (138 ~ 145Ma), late Early Cretaceous (106 ~ 133Ma), and Late Cretaceous (88 ~ 97Ma). Late Triassic volcanic rocks in NE China mainly distribute in the eastern Jilin-Heilongjiang provinces and the Lesser Xing'an-Zhangguangcai Ranges. The former consists of A-type rhyolite, the latter is composed of bimodal volcanic rocks, implying that they formed under an extensional environment after the final closure of the Paleo-Asian Ocean. Early-Middle Jurassic volcanic rocks occur in the eastern Jilin-Heilongjiang provinces, the Lesser Xing'an-Zhangguangcai Ranges, and the Erguna district. Those in the eastern Jilin-Heilongjiang provinces and the Erguna district are composed of calc-alkaline volcanic rocks, suggesting that they formed under the subductions of the Paleo-Pacific plate beneath the Eurasian continent and of the Mongol-Okhotsk oceanic plate beneath the Erguna Massif, respectively. However, those in the Lesser Xing'an-Zhangguangcai Ranges are a set of bimodal volcanic rocks, implying that they formed under an extensional environment similar to a back-arc setting of double-direction subduction. Middle-Late Jurassic and early Early Cretaceous volcanic rocks only distribute to the west of the Songliao basin, including the Great Xing'an Range and northern Hebei-western Liaoning provinces. Middle-Late Jurassic volcanic rocks consist of basaltic trachy-andesite, trachy-andesite, and trachyte, whereas early Early Cretaceous volcanic rocks are composed of A-type rhyolite and alkali rhyolite, which formed under an extensional environment related to the collapse or delamination of the thickened lower crust. The late Early Cretaceous volcanic rocks are widespread in NE China, those in the eastern Jilin-Heilongjiang provinces belong chemically to a set of calc-alkaline series, whereas those from the Songliao basin and the Great Xing'an Range are a bimodal volcanic rocks. The former marks the subduction of the Paleo-Pacific plate beneath the Eurasian continent, the latter could form under an extensional environment related to the delamination of the thickened lower crust and/or a back-arc setting. Late Cretaceous volcanic rocks mainly occur in the eastern margin of the Eurasian continent, and consist of calc-alkaline series in the continental margin and alkali basalts in intracontinent, suggesting that they formed under the subduction of the Paleo-Pacific plate. Taken together, we conclude: 1) that the subduction of the circum Pacific tectonic system beneath the Eurasian continent began in the Early Jurassic and took place in three times (Early Jurassic, late Early Cretaceous, and Late Cretaceous) during Mesozoic; the influencing spatial extent of the circum Pacific tectonic system mainly include the Songliao basin and its to east; the active continental margin and Paleo-subduction zone are favorable sites in search of porphyry ore deposits,

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第一作者简介: 许文良,男,1959年生,博士,教授,主要从事火成岩成因与岩石圈动力学研究, E-mail: xuwl@jlu.edu.cn

whereas an extensional regions within intracontinent are favorable for the formation of the epithermal hydrothermal ore deposits in the eastern Jilin-Heilongjiang provinces; 2) that the Mongol-Okhotsk tectonic system experienced Early Mesozoic subduction beneath the Erguna Massif, and Middle Jurassic and early Early Cretaceous thrusting events; the influencing spatial extent of the Mongol-Okhotsk tectonic system include to west of the Songliao basin and northern margin of the North China Craton; the Early Mesozoic subduction of the Mongol-Okhotsk oceanic plate beneath the Erguna Massif is favorable for the formation of porphyry deposits, whereas the extensional environment related to the collapse or delamination of the thickened lower crust in Late Jurassic and Early Cretaceous are favorable for formation of polymetallic deposits.

Key words Northeast China; Mesozoic volcanic rocks; Spatial and temporal variation; Circum-Pacific tectonic system; Mongol-Okhotsk tectonic system; Ore-forming background

摘要 本文系统总结了东北地区中生代火山岩的年代学、岩石组合及其时空分布规律,以便对环太平洋构造体系和蒙古-鄂霍茨克构造体系中生代的演化历史及其东北地区中生代区域成矿背景给出制约。基于火山岩中锆石 U-Pb 定年结果,东北地区中生代火山作用可划分成六期:晚三叠世(200~228Ma)、早-中侏罗世(173~190Ma)、中-晚侏罗世(158~166Ma)、早白垩世早期(138~145Ma)、早白垩世晚期(106~133Ma)和晚白垩世(88~97Ma)。晚三叠世火山作用主要分布在吉黑东部和小兴安岭-张广才岭地区,前者为 A 型流纹岩,后者为双峰式火山岩组合,它们共同揭示了古亚洲洋最终闭合后的伸展环境;早-中侏罗世火山岩主要分布在吉黑东部、小兴安岭-张广才岭和额尔古纳地区,吉黑东部和额尔古纳地区早-中侏罗世钙碱性火山岩的存在分别标志着古太平洋板块和蒙古-鄂霍茨克洋板块俯冲作用的发生,而小兴安岭-张广才岭早-中侏罗世火山岩则以双峰式火成岩组合为特征,反映了双向俯冲的弧后伸展环境;中-晚侏罗世和早白垩世早期火山岩主要分布在松辽盆地以西和冀北-辽西地区,前者为碱性-亚碱性的过渡系列,主要由玄武粗安岩、粗安岩和少量粗面岩组成,后者为 A 型流纹岩或碱性流纹岩组成,这些火山岩形成于加厚陆壳的坍塌或拆沉阶段;早白垩世晚期火山岩广泛分布于东北地区,吉黑东部为钙碱性火山岩组合,而松辽盆地和大兴安岭地区则主要为双峰式火山岩组合,前者标志着古太平洋板块的俯冲,后者与早期加厚陆壳的拆沉和/或类似弧后的伸展环境有关;晚白垩世火山岩主要分布在吉黑东部和陆内,前者为钙碱性火山岩组合,后者为碱性玄武岩,反映了来自东部环太平洋构造体系的俯冲作用。综合上述中生代火山岩的岩石组合及时空分布特征,可以判定:1)环太平洋构造体系对东亚大陆下的俯冲始于早侏罗世,中生代期间存在早侏罗世、早白垩世晚期和晚白垩世三次俯冲事件,其影响的空间范围主要在松辽盆地及其以东地区,陆缘和古俯冲带是寻找斑岩型矿床的有利场所,而陆内的伸展区域主要与浅成低温热液矿床有关;2)蒙古-鄂霍茨克构造体系经历了中生代早期的俯冲事件和中侏罗世及早白垩世早期两次陆内推覆事件,其影响的空间范围主要在松辽盆地以西地区和华北地块北缘,中生代早期的俯冲事件主要与活动陆缘背景下的斑岩型矿床关系密切,而晚侏罗世和早白垩世两次与加厚陆壳拆沉有关的伸展背景有利于多金属矿床的形成。

关键词 东北地区;中生代火山岩;时空变化;环太平洋构造体系;蒙古-鄂霍茨克构造体系;成矿背景

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1 引言

东北地区位于中亚造山带的东段,在古生代期间,该区经历了古亚洲洋构造体系的演化,以多个微陆块(自西向东包括额尔古纳地块、兴安地块、松嫩-张广才岭地块、佳木斯地块和兴凯地块)之间的拼合和古生代晚期-中生代早期古亚洲洋的最终闭合为特征(Sengör and Natalin, 1996; Li, 2006; Tang *et al.*, 2013)(图 1a)。在中生代期间,东北地区不仅经历了环太平洋构造体系的叠加与改造,而且又有蒙古-鄂霍茨克构造体系的叠加与改造(Xu *et al.*, 2009; Wu *et al.*, 2011; 孟恩等, 2011)。近年来,在东北地区区域找矿方面已经取得了重大突破,尤其是大型-超大型斑岩型钼矿的发现,已经使东北成为我国重要的有色及贵金属资源基地(陈衍景等, 2012)。目前,对东北地区中生代区域成矿背景的认识,多数强调环太平洋构造体系的影响(Ge *et al.*, 2007),而忽略了蒙古-鄂霍茨克构造体系对东北地区的影响(图 1b)(孟恩等, 2011)。即使在环太平洋构造体系的研

究中,目前仍有一些关键科学问题没有得到解决,比如,目前对环太平洋构造体系开始的时间还有争论(是晚三叠世?还是早-中侏罗世?)(吉林省地质矿产局, 1988; 黑龙江省地质矿产局, 1993; 赵越等, 1994; 赵春荆等, 1996; 孙德有等, 2005; Wu *et al.*, 2007; 裴福萍等, 2008; Zhou *et al.*, 2009),对环太平洋构造体系中生代的演化历史还不是很清楚,如中生代有几次俯冲事件?环太平洋构造体系影响的空间范围如何?等等。与环太平洋构造体系研究相比,对蒙古-鄂霍茨克构造体系对我国境内的影响研究更少,目前还多数强调该大洋板块的向北俯冲,对其是否存在向南俯冲和中生代的演化历史了解更少。虽然近年来在对东北地区花岗岩的时代、成因与陆壳增生研究中已经取得了突破性进展(Wu *et al.*, 2011),但由于花岗岩成因类型的多样性以及岩浆源区性质的制约,使其在成分极性变异方面难以得到很好地表现,与之相比,火山岩具有更好的成分变化极性,这对揭示大洋板块的俯冲历史更具有重要的指示意义。鉴于此,本文系统总结了东北地区中生代火山作用的年代学以及不同期次火山岩组合的时空变异,以此来揭示东北地区中生代不

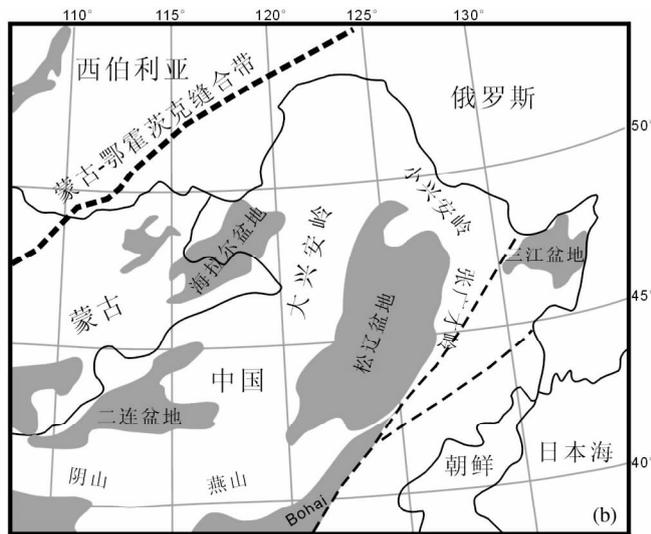
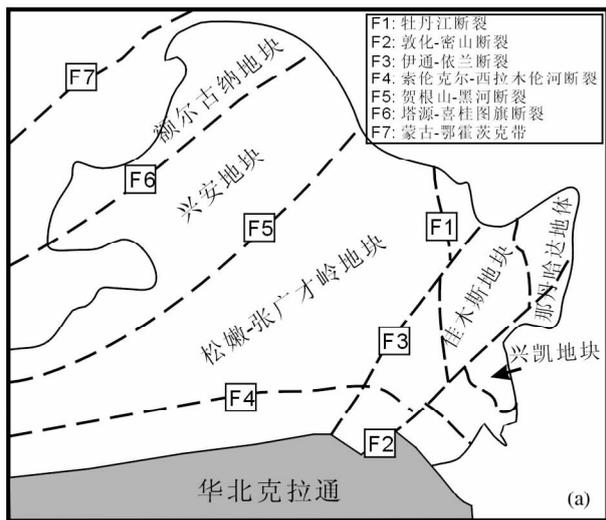


图1 东北地区的构造分区 (a,据 Wu *et al.*, 2011) 和东北的主要地理单元 (b,据 Meng, 2003)

Fig.1 Tectonic subdivisions of northeastern (NE) China (a, after Wu *et al.*, 2011) and main topographical units in NE China (b, after Meng, 2003)

同构造体系的演化历史及其影响的空间范围,这对全面了解和揭示中生代区域成矿背景和指导东北地区的区域找矿均具有重要意义。

2 东北地区古亚洲洋构造体系结束的时限

在讨论东北地区中生代构造体制之前,首先应了解影响该区的古生代构造体制——古亚洲洋构造体系的演化历史和该构造体系的结束时间。目前多数学者认为东北地区多个微陆块的拼合过程主要在加里东期完成,如额尔古纳地块与兴安地块的拼合在早古生代早期已经完成拼合(葛文春等,2007)、松嫩-张广才岭地块与佳木斯地块的拼合于加里东晚期已经拼合(Wang *et al.*, 2012),而兴安地块与松嫩-张广才岭地块的拼合于早石炭世完成(黑龙江省地质矿产局,1993)。这些微陆块作为一个整体(称为阿穆尔地块, Sengör and Natalin, 1996;或佳蒙地块, Wang *et al.*, 2009)与华北地块的拼合过程以及古亚洲洋的最终闭合时间可以从华北地块北缘晚古生代-早中生代的演化历史得到回答。

首先,从华北地块北缘中部晚古生代的岩浆作用历史可以看出,在泥盆纪,该区发育了一套碱性杂岩和基性-超基性杂岩体,前者以A型花岗岩为主,它们共同构成了双峰式火成岩组合,揭示了华北地块北缘伸展构造环境的存在(Zhang *et al.*, 2009)。

其次,华北地块北缘晚石炭世-早二叠世钙碱性火成岩组合的存在,揭示了该区古亚洲洋向南俯冲于华北地块之下的发生(Zhang *et al.*, 2009),吉林中部-延吉地区早二叠世钙碱性火山岩的存在也进一步证明了早二叠世期间古亚洲洋向南俯冲作用的发生(曹花花等,2012)。

第三,进入到晚二叠世,松嫩-张广才岭地块东缘和佳木斯地块西缘以及华北克拉通北缘双峰式火成岩组合的出现,表明该区处于伸展环境(Liu *et al.*, 2010; 王子进等,2013; Yu *et al.*, 2013),这与中二叠世过铝质石榴石二长花岗岩所揭示的碰撞作用之后的伸展环境相适应(Cao *et al.*, 2013)。

第四,至早-中三叠世,兴蒙造山带南缘碰撞型花岗岩的产生标志着古亚洲洋的最终闭合(孙德有等,2004;李锦轶等,2007;Cao *et al.*, 2013)。

综合上述研究成果,可以看出,在东北地区,古亚洲洋的最终闭合时间最晚可能发生在中三叠世,自此以后,该区进入到环太平洋构造体系和蒙古-鄂霍茨克构造体系的演化阶段。

3 东北地区中生代火山作用的时空格架

根据对东北地区中生代火山岩锆石 U-Pb 定年结果的统计,可以将东北地区中生代火山作用划分成六期(图2),晚三叠世(200~228Ma)、早-中侏罗世(173~190Ma)、晚侏罗世(158~166Ma)、早白垩世早期(138~145Ma)、早白垩世晚期(106~133Ma)、晚白垩世(88~97Ma),这六期火山岩的分布如(图3)所示(Xu *et al.*, 2013)。

1)晚三叠世(200~228Ma):该期火山岩主要分布在吉黑东部和小兴安岭-张广才岭地区,前者主要为一套A型流纹岩组合(Xu *et al.*, 2009),后者为一套双峰式火山岩组合(Wang *et al.*, 2011a)。

2)早-中侏罗世(173~190Ma):该期火山岩主要分布在吉黑东部、小兴安岭-张广才岭和额尔古纳地区。吉黑东部

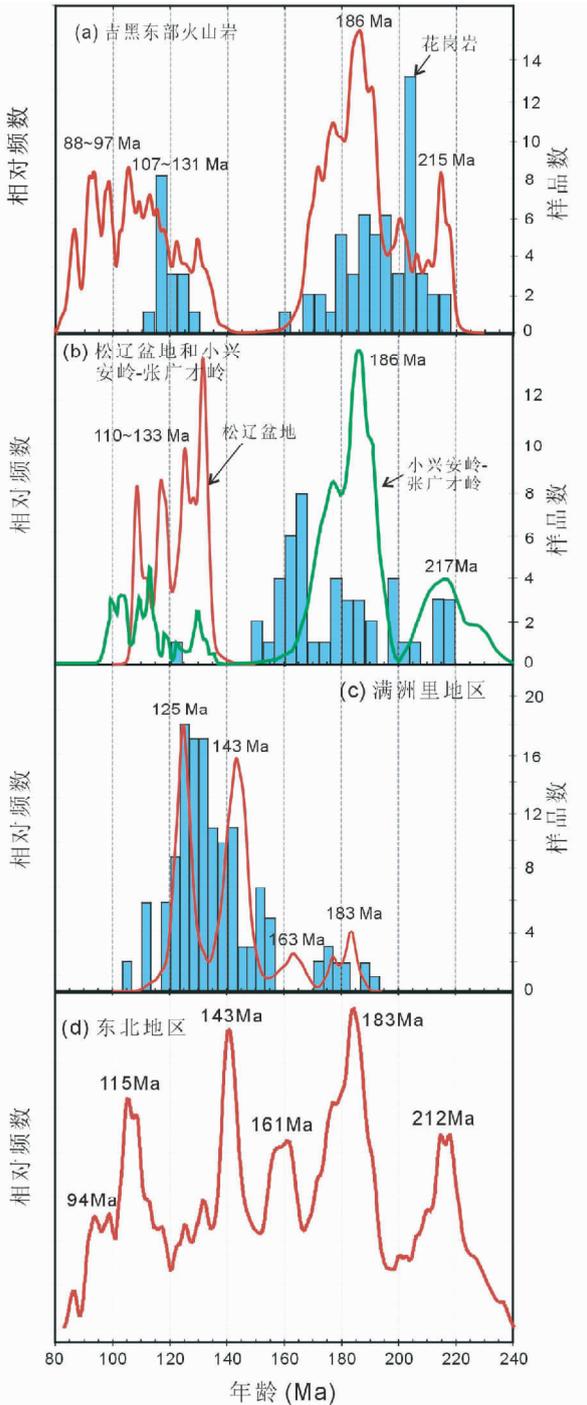


图2 中国东北中生代火山岩锆石 U-Pb 年龄频数图
(a)-吉黑东部区;(b)-松辽盆地和小兴安岭-张广才岭;(c)-满洲里地区;(d)-整个东北地区. 直方图部分为东北地区中生代花岗岩,引自 Zhang *et al.* (2010). LXZR-小兴安岭-张广才岭
Fig. 2 Probability plots of the zircon U-Pb ages for the Mesozoic volcanic rocks in NE China

(a)-the eastern Heilongjiang-Jilin provinces; (b)-the Songliao basin and the Lesser Xing'an-Zhanguangcai Ranges; (c)-the Manzhouli area; (d)-NE China. Age data for Mesozoic granitoids in NE China after Zhang *et al.* (2010). LXZR-the Lesser Xing'an-Zhanguangcai Ranges

(敦-密断裂东南区)和额尔古纳地区早-中侏罗世火山岩主要为钙碱性火山岩系列(许文良等,2008;裴福萍等,2009;Xu *et al.*, 2013),而小兴安岭-张广才岭则为是一套双峰式火成岩组合(唐杰等,2011;Yu *et al.*, 2012;徐美君等,2013)。

3)中-晚侏罗世(158~166Ma):从目前的定年结果来看,该期火山岩只分布在松辽盆地以西地区,主要为一套亚碱性-碱性的过渡类型。岩石组合为:玄武粗安岩-粗安岩-少量粗面岩(Wang *et al.*, 2006; Yang and Li, 2008; Zhang *et al.*, 2008a, 2010; 孟恩等,2011;徐美君等,2011)。

4)早白垩世早期(138~145Ma):该期火山岩的分布范围与晚侏罗世火山岩类似,目前在松辽盆地以东地区尚未发现,主要为一套A型流纹岩,在个别地区为碱性流纹岩(其中存在大量的钠铁闪石)(孟恩等,2011;徐美君等,2011;王建国等,2013)。

5)早白垩世晚期(106~133Ma):该期火山岩不仅在大兴安岭和吉黑东部广泛分布(Wang *et al.*, 2006; Zhang *et al.*, 2008a, 2010; 孟恩等,2011;徐美君等,2011),而且在松辽盆地底部断陷层也广泛产出(裴福萍等,2008;章凤奇等,2009)。

6)晚白垩世(88~97Ma):该期火山岩主要分布在吉黑东部(Ji *et al.*, 2007; 张磊等, 2009; Zhang *et al.*, 2012; 于介江等,2013),在松辽盆地中也发现该期岩浆事件(王璞珺等,2009)。此外,在辽西阜新(许文良等,1999)、辽南曲家屯(王微等,2006)也有该期火山岩的发现。

4 东北地区中生代环太平洋构造体系的演化历史与区域成矿

4.1 古太平洋构造体系俯冲于欧亚大陆之下开始时间的确定

在东北乃至整个中国东部环太平洋构造体系中生代的演化历史并不是很清楚。首先,古太平洋板块俯冲开始的时间一直是个争论的问题,目前主要有两种观点,一种认为是晚三叠世(彭玉鲸和陈跃军,2007),而另一种认为是早-中侏罗世(Xu *et al.*, 2009; Wu *et al.*, 2011)。吉黑东部晚三叠世A型流纹岩的存在(Xu *et al.*, 2009)和张广才岭晚三叠世双峰式火山岩的存在(Wang *et al.*, 2011a),与张广才岭晚三叠世碱长花岗岩一起(Wu *et al.*, 2011),表明该区晚三叠世期间处于一种伸展环境,并不代表古太平洋构造体系俯冲作用的开始。

钙碱性火山岩组合的存在可作为判定古俯冲作用发生的有效标志。近年来,吉黑东部(东宁-汪清-珲春一线)早-中侏罗世(173~190Ma)钙碱性火山岩组合的发现,为判定古太平洋构造体系开始俯冲的时间提供了可信的证据(许文良等,2008;裴福萍等,2009),而小兴安岭-张广才岭地区同时代双峰式火成岩组合的存在(唐杰等,2011;Yu *et al.*, 2012;徐美君等,2013),揭示了自陆缘向陆内火山岩成分极性的变

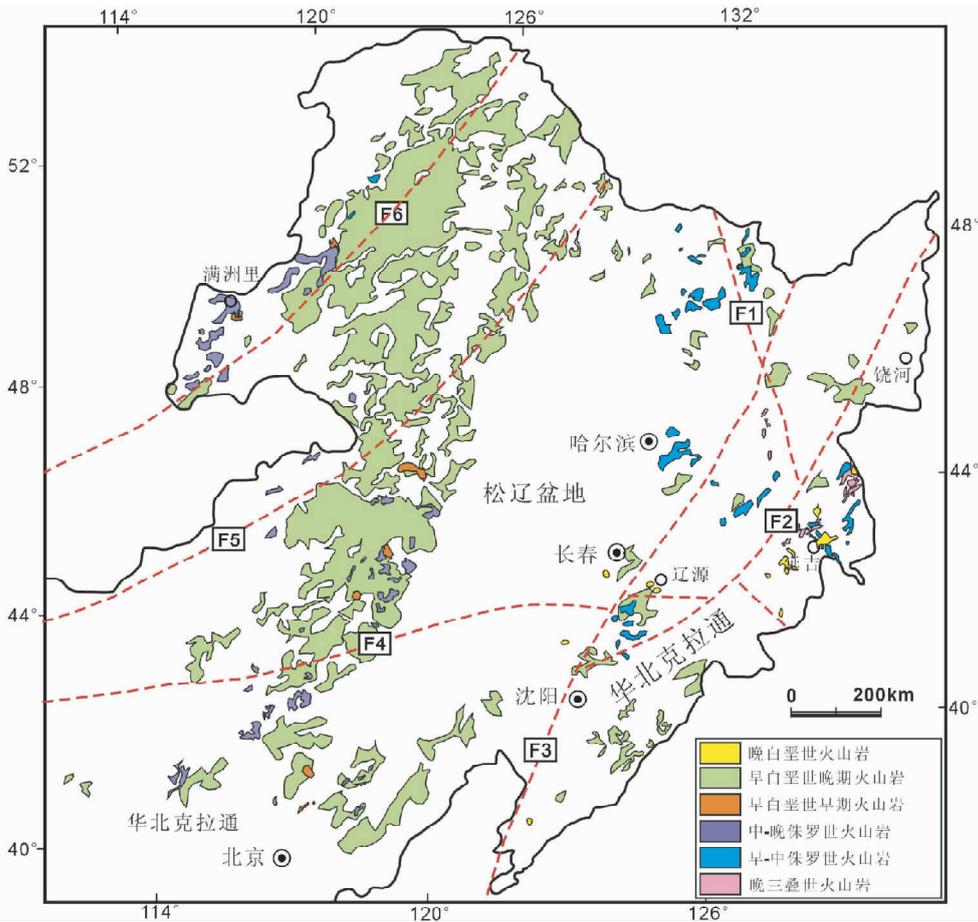


图3 东北地区中生代火山岩的分布图

Fig. 3 Distribution map of Mesozoic volcanic rocks in NE China

化——即揭示了来自东部俯冲作用的开始(许文良等, 2012),这一俯冲作用也导致了位于松嫩-张广才岭地块与佳木斯地块之间黑龙江杂岩的构造就位(李锦轶等, 1999; Wu *et al.*, 2007; Zhou *et al.*, 2010)。

4.2 中侏罗世晚期-早白垩世早期:古太平洋板块俯冲的间歇期

首先,从已有的火山岩定年结果来看,目前在吉黑东部尚未发现 140 ~ 165Ma 的火山岩,而这一阶段的火山岩主要分布在松辽盆地以西地区,这暗示此阶段的岩浆事件与古太平洋板块的俯冲作用无关。

其次,位于东北东部的那丹哈达地体蛇绿混杂岩中存在中侏罗世晚期(166Ma 左右)的火山岩和辉长岩(赵海玲等, 1996;程瑞玉等,2006),表明该期欧亚大陆陆缘东侧是一个处于拉张环境的海盆,而地体拼贴时间发生在晚侏罗世-早白垩世之间,蛇绿混杂岩已被早白垩世花岗岩所穿切就证明了这一点(程瑞玉等,2006)。

综上所述,在中侏罗世晚期-早白垩世早期阶段,古太平洋板块对欧亚大陆下并没有发生俯冲作用。

4.3 早白垩世晚期和晚白垩世:古太平洋板块对欧亚大陆下的两次俯冲作用

早白垩世晚期(106 ~ 133Ma)在东北地区,该期火山作用广泛而强烈,然而,在东北地区的不同区域该期火山岩的成分具有明显变异。在吉黑东部(包括华北地块东北缘)代表性地层有:自北向南包括皮克山组、裴德组、东山组、金沟岭组、泉水村组、二股砬子组,除出露于佳木斯地块上的部分样品具有相对较高的碱性组分外,该期火山岩主要由一套钙碱性火山岩所组成(Yu *et al.*, 2009),代表了古太平洋板块(Izanagi)俯冲作用的存在。而同时代的火山岩,在松辽盆地和大兴安岭地区主要表现为双峰式火山岩组合(葛文春等, 1999;Wang *et al.*, 2006; Zhang *et al.*, 2008a, 2010; 裴福萍等,2008;孟恩等,2011),相应时代的花岗岩主要为碱性或碱长花岗岩(Wu *et al.*, 2011)。自陆缘向陆内,火山岩中的碱性组分具有增高的成分极性变化,表明了来自东部板块俯冲作用的发生。

晚白垩世(88 ~ 97Ma)晚白垩世火山岩主要分布在东北的东部,在陆缘区(东宁-珲春一线)代表性火山岩地层有绥芬河组(Ji *et al.*, 2007)和屯田营组(Xu *et al.*, 2013),它们主要

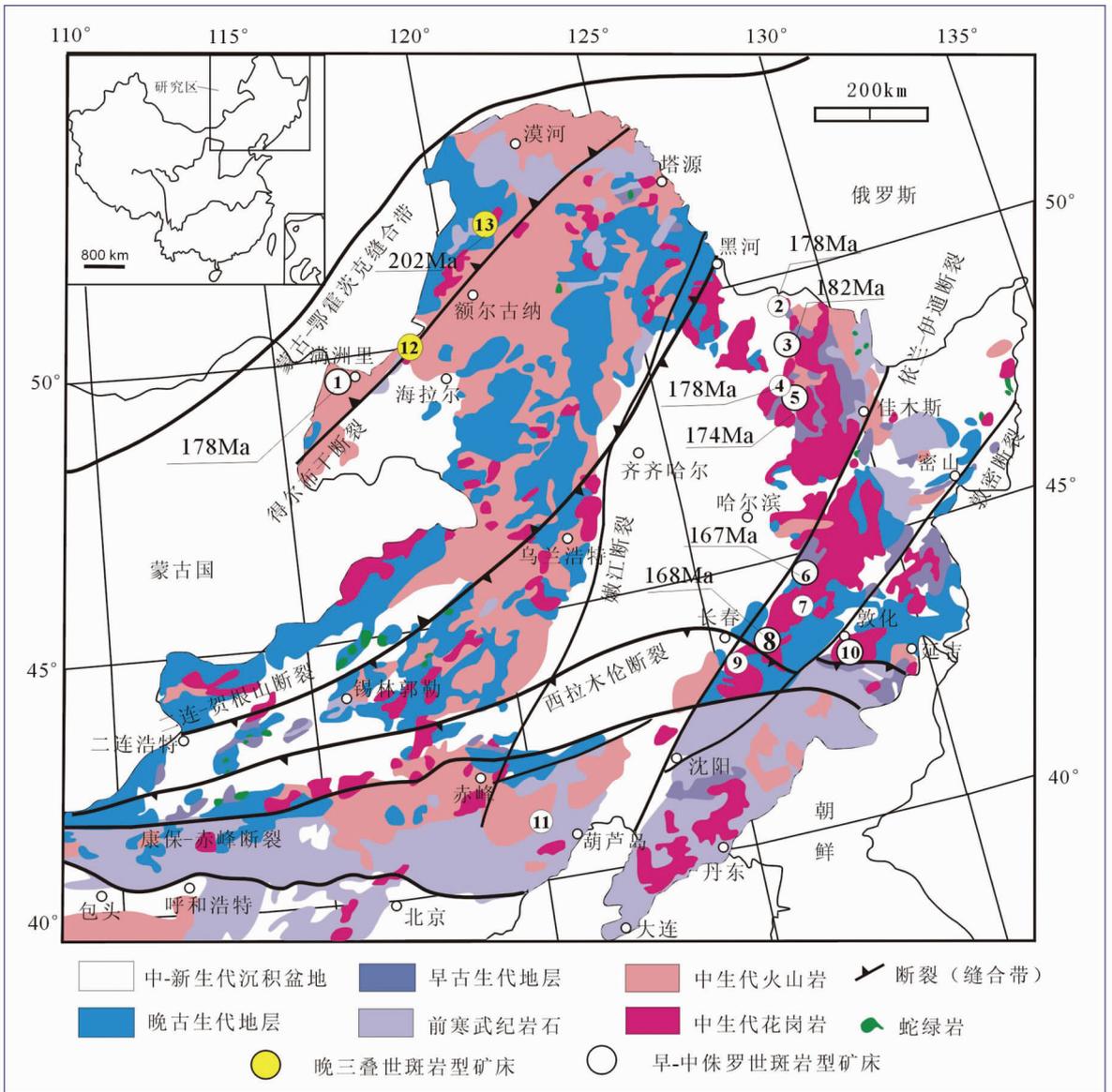


图4 东北地区早中生代斑岩型矿床分布图(据白令安等,2012;孙景贵等,2012;陈衍景等,2012 修改)

矿床名称:1-乌努格吐山;2-翠宏山;3-霍吉河;4-翠岭;5-鹿鸣;6-福安堡;7-季德屯;8-大黑山;9-后倒木;10-刘生店;11-杨家杖子;12-八大关;13-太平川

Fig. 4 Distribution map of Early Mesozoic porphyry-type ore deposits in NE China (modified after Bai *et al.*, 2012; Sun *et al.*, 2012; Chen *et al.*, 2012)

由一套钙碱性火山岩所组成,而在陆内辽西的阜新、辽南的曲家屯,同时代的火山岩主要为一套碱性玄武岩(许文良等,1999;Yan *et al.*, 2003;王微等,2006),在佳木斯地块上,同时代的侵入岩为一套碱性辉长岩(张磊等,2009)。自陆缘向陆内火山岩碱性组分的增高,表明来自东部俯冲作用的发生。

4.4 古太平洋板块中生代对欧亚大陆下的俯冲作用与区域成矿

从吉黑东部中生代火山作用的时代与性质可以看出,在中生代期间,古太平洋板块向欧亚大陆下的俯冲作用主要发

生在三个阶段:即早-中侏罗世、早白垩世晚期和晚白垩世。这三个阶段,欧亚大陆的东北部陆缘处于活动大陆边缘的构造背景,是我国东北地区东部寻找斑岩型铜钼矿床、浅成热液金铜矿床的有利场所。近年来,吉黑东部早-中侏罗世斑岩型钼矿正是古太平洋板块俯冲之下的产物(图4)(Ge *et al.*, 2007;孙景贵等,2012;陈衍景等,2012)。与早-中侏罗世相比,在吉黑东部,早白垩世晚期和晚白垩世岩浆作用的强度明显减弱,尤其是后者,在靠近陆缘的外侧(如东宁-琿春一线),主要以寻找富金斑岩型铜矿床、高硫化型金矿床为主(Han *et al.*, 2013),而进入到陆缘的内侧,主要表现为类

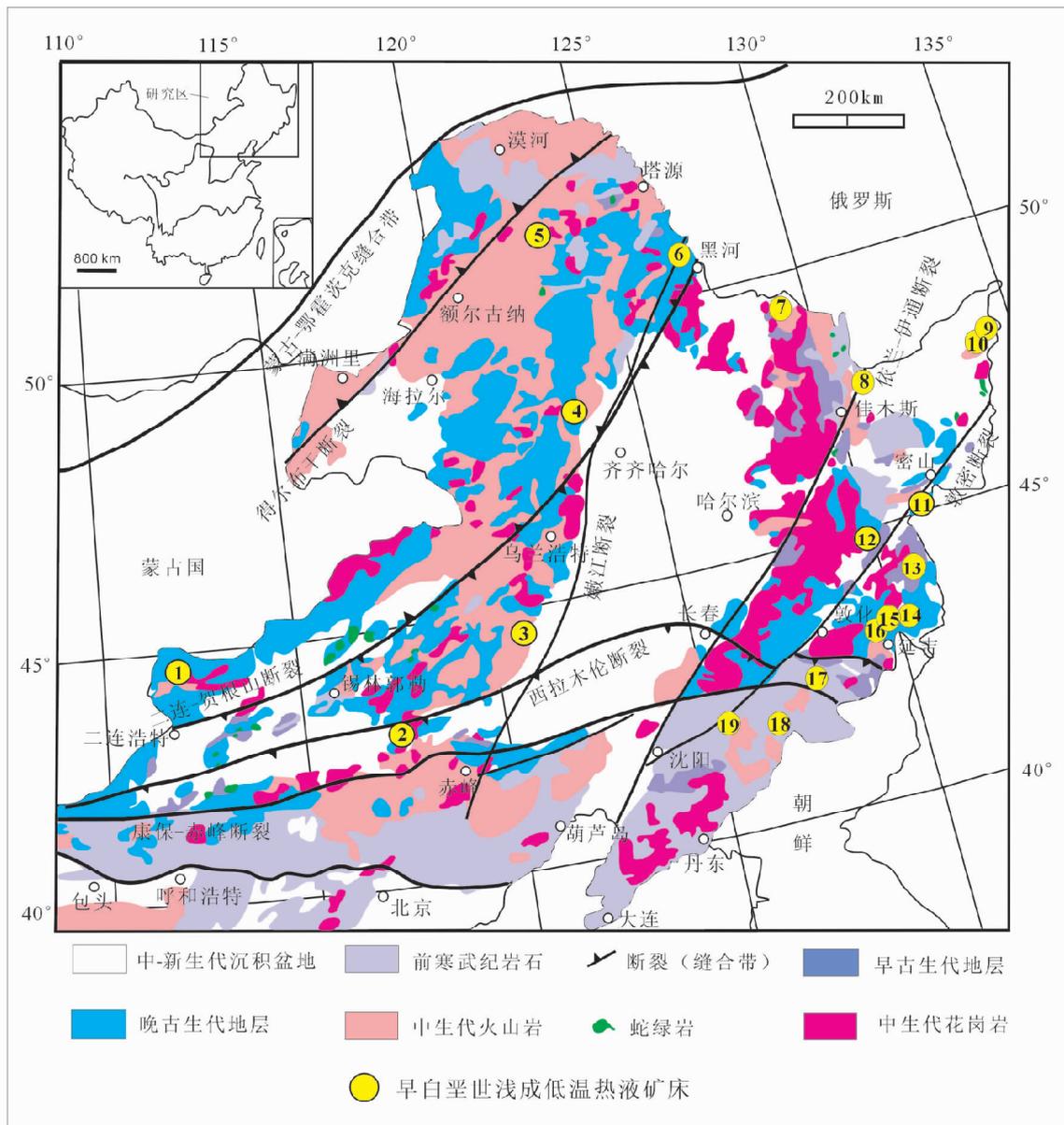


图5 东北地区早白垩世浅成热液矿床分布图(据白令安等,2012;孙景贵等,2012;Han *et al.*, 2013; 陈衍景等,2012 修改)
 矿床名称:1-乌兰德勒;2-小东沟;3-敖仓花;4-太平沟;5-兴阿;6-三道湾子;7-东安;8-平顶山;9-四平山;10-大王砬子;11-金场沟;12-砍楼沟;13-九沸沟;14-小西南岔;15-杜荒岭;16-闹枝;17-天宝山;18-板庙子;19-香炉碗子

Fig. 5 Distribution map of Early Cretaceous epithermal deposits in NE China (modified after Bai *et al.*, 2012; Sun *et al.*, 2012; Han *et al.*, 2013; Chen *et al.*, 2012)

似弧后的伸展环境,与之相对应的矿产主要以浅成低温热液型金矿床为主(图5)(孙景贵等,2009)。

5 蒙古-鄂霍茨克构造体系中生代的演化历史与区域成矿

5.1 蒙古-鄂霍茨克缝合带演化历史概述

蒙古-鄂霍茨克缝合带从蒙古的杭爱山脉一直延伸到鄂霍茨克海的乌达海湾,长约3000km,宽约300km,古生代和早

中生代时期位于西伯利亚板块、华北板块和古太平洋板块之间。它是东亚北部一条具有较长地质历史的造山带,它在东亚大陆形成演化的历史上占有极为重要的位置(李锦轶等,2009)。由于该造山带主要位于俄罗斯和蒙古境内,对其演化历史的研究,尤其是该构造带对我国境内影响的研究较少。目前主要认识是,蒙古-鄂霍茨克洋在晚古生代末期较为宽广(Cogné *et al.*, 2005),局部已经存在俯冲(Tomurtogoo *et al.*, 2005; Mazukabzov *et al.*, 2010),并延续至三叠纪;由于西伯利亚板块相对于中蒙地块的旋转,造成了鄂霍茨克板

块从西向东的剪刀式关闭,西部晚三叠世开始闭合、向东部在晚侏罗世闭合 (Zonenshain *et al.*, 1990; Zorin, 1999; Sorokin *et al.*, 2004),也有人认为东段的碰撞可以持续到晚侏罗世-早白垩世 (Kravchinsky *et al.*, 2002)。就蒙古-鄂霍茨克洋的俯冲方向而言,多数学者认为该大洋板块向北俯冲 (Zorin, 1999),是否存在向南俯冲一直存在争议 (Wu *et al.*, 2011)。近年来,虽然对额尔古纳地块中早中生代花岗岩的成因和斑岩型矿床的成因一直强调该构造带的影响 (秦克章等, 1999; 陈志广等, 2010; Wu *et al.*, 2011; 王伟等, 2012; 余宏全等, 2012),但是,对蒙古-鄂霍茨克缝合带的演化历史及其对我国境内影响的时空范围并不清楚。大兴安岭及燕辽地区中生代火山岩的岩石组合及其时空分布,结合与吉黑东部中生代火山岩组合与时空分布特征的对比,可以对蒙古-鄂霍茨克缝合带中生代的演化历史给出限定。

5.2 三叠纪:蒙古-鄂霍茨克大洋板块向南俯冲作用的确定

在已有的研究中,中生代早期蒙古-鄂霍茨克洋的演化主要强调向北俯冲作用,而近年来在蒙古国和我国境内额尔古纳地块早中生代花岗岩和成矿作用的研究,揭示该大洋板块中生代早期存在向南俯冲的构造演化历史。如在蒙古国额尔登特发育了岛弧环境的侵入杂岩,形成特大型斑岩铜-钼矿床,成矿成岩时代为240Ma左右 (江思宏等, 2010);在中蒙古形成的 Hangayn 岩基(255 ~ 230Ma),属安第斯型大陆边缘弧环境 (Tomurtogoo *et al.*, 2005; Orolmaa *et al.*, 2008);此外,在我国境内额尔古纳地块中的太平川斑岩型铜钼矿床形成于晚三叠世 (锆石 U-Pb 年龄为 202 ± 6 Ma),该矿床形成于鄂霍茨克洋板块向额尔古纳地块之下俯冲的活动陆缘弧环境 (陈志广等, 2010); Wu *et al.* (2011) 对额尔古纳地块北段早中生代花岗岩的研究也认为,它们的形成也与蒙古-鄂霍茨克带的演化有关。综合上述研究成果,可以判定,在中生代早期,我国境内已经受到了蒙古-鄂霍茨克缝合带的影响,并主要表现为蒙古-鄂霍茨克大洋板块向额尔古纳地块和中蒙古地块之下的俯冲作用。位于内蒙古自治区林西中三叠世高镁 adakitic 安山岩的形成应是北部鄂霍茨克洋板块向额尔古纳-中蒙古地块之下俯冲作用的远程效应,即在伸展环境下早期俯冲板片熔融的熔体与地幔橄榄岩反应的结果 (Liu *et al.*, 2012)。

5.3 早侏罗世(180 ~ 185Ma):蒙古-鄂霍茨克大洋板块对额尔古纳地块下的俯冲作用

Zorin (1999) 的研究认为早侏罗世期间蒙古-鄂霍茨克大洋板块主要表现为向北的俯冲作用,还有部分学者认为此时在西部已经发生蒙古-鄂霍茨克洋的闭合 (Metelkin *et al.*, 2007)。那么,该期是否存在蒙古-鄂霍茨克大洋板块对额尔古纳地块之下的俯冲作用? 额尔古纳地块上该期的岩浆作用性质对此给予了回答。首先,在额尔古纳地块上的额尔古纳-根河地区确定了一套早侏罗世玄武岩-玄武安山岩钙碱性

火山岩组合 (Zhang *et al.*, 2008a; Xu *et al.*, 2013),该套火山岩组合反映了活动陆缘的构造背景;其次,在满洲里地区与乌奴格吐山斑岩型铜钼矿床密切相关的花岗岩和具有矿化的蚀变花岗岩的形成时代也为早侏罗世 (180 ~ 183Ma, 王伟等, 2012),而反映成矿时代的 Re-Os 同位素年龄为 178Ma (余宏全等, 2012)。综合上述特征,可以判定,在早侏罗世期间,额尔古纳地块西缘处于活动陆缘的构造背景。此时,小兴安岭-张广才岭存在同时代的双峰式火成岩,由此可以排除额尔古纳地块上发育的早侏罗世钙碱性火山岩并不是古太平洋板块俯冲作用的产物,而应是蒙古-鄂霍茨克大洋板块向额尔古纳地块下俯冲作用的结果。

5.4 中侏罗世(165 ~ 170Ma):蒙古-鄂霍茨克洋的闭合与区域陆壳加厚事件

从已有的研究成果中可知,中生代期间,蒙古-鄂霍茨克洋的闭合自西向东具有剪刀式闭合的特征,东部闭合的时间可能持续到晚侏罗世-早白垩世 (Kravchinsky *et al.*, 2002; Sorokin *et al.*, 2010)。对于额尔古纳地块而言,其西北侧蒙古-鄂霍茨克洋的闭合时间应发生在中侏罗世,证据有:1) 在位于兴安地块与松嫩-张广才岭地块之间黑河-孙吴地区发现了具有“S”型花岗岩地球化学属性的白云母花岗岩,其形成时代为 168 ± 2 Ma,我们认为该白云母花岗岩形成于陆壳加厚的背景 (作者未发表资料),这与在该区发现的二云母花岗岩的形成时代相吻合 (Wu *et al.*, 2011);2) 在冀北-辽西地区,广泛存在一个区域性的地层不整合——即在海房沟组之下存在的自北向南的逆冲构造,表明了陆壳加厚过程的存在 (赵越等, 1994, 2004; Zhang *et al.*, 2008b),其发生的时间可以从海房沟组之上的蓝旗组和髫髻山组火山岩的形成时代 (155 ~ 158Ma) 得到限制 (Gao *et al.*, 2004; 刘健等, 2006; Yang and Li, 2008);3) 在满洲里-额尔古纳地区形成了一套具有碱性-亚碱性过渡属性的火山岩组合,以塔木兰沟组玄武粗安岩-粗安岩-粗面岩为代表,而在冀北辽西地区则以蓝旗组和髫髻山组为代表,前者形成于 158 ~ 166Ma 之间 (其峰期时代为 162Ma) (Wang *et al.*, 2006; Zhang *et al.*, 2008a; 孟恩等, 2011; 徐美君等, 2013; Xu *et al.*, 2013),后者形成于 155 ~ 158Ma 之间,它们都形成于与加厚陆壳垮塌阶段相对应的伸展环境 (Xu *et al.*, 2013);4) 该期火山岩 (155 ~ 166Ma) 只分布在松辽盆地以西地区,反映其形成应与蒙古-鄂霍茨克缝合带的演化有关,而与环太平洋构造体系无关,而该期火山岩自北向南时代变新的迁移趋势 (Zhang *et al.*, 2008a, 2010) 再次证明它们的形成应是蒙古-鄂霍茨克缝合带演化的产物。

综合上述特征,可以得出在中侏罗世期间,大兴安岭西坡-冀北-辽西地区存在一次重要的陆壳加厚过程,这一过程恰恰是蒙古-鄂霍茨克缝合带闭合的时间,而广泛发育的中-晚侏罗世火山岩应是这次加厚陆壳坍塌或拆沉阶段的产物 (孟恩等, 2011; Xu *et al.*, 2013)。

5.5 早白垩世早期(138~145Ma):区域伸展事件与A型流纹岩的形成

早白垩世早期火山岩的分布与中-晚侏罗世火山岩的分布类似——只分布在松辽盆地以西地区(包括大兴安岭西坡和冀北-辽西地区),代表性地层包括大兴安岭北段的吉祥峰组 and 南段的玛尼吐组、冀北-辽西的张家口组,前者主要形成于138~145Ma之间(峰期年龄为142Ma, Wang *et al.*, 2006; Zhang *et al.*, 2008a; Ying *et al.*, 2010; 孟恩等, 2011; 徐美君等, 2011; 王建国等, 2013),后者主要形成于135~136Ma之间(牛宝贵等, 2003; 张宏等, 2005)。自北向南,该期火山岩形成的时间具有变新的趋势(Xu *et al.*, 2013)。在地球化学属性上,该期火山岩具有A型流纹岩的特征,在大兴安岭南段该期火山岩(142Ma)呈不整合覆盖在满克头鄂博组(162Ma)之上,并为富含钠铁闪石的碱性流纹岩(王建国等, 2013)。上述岩石组合揭示该区在早白垩世早期处于伸展环境。从冀北-辽西该期火山岩的分布可知,它们不整合产于土城子组之上,而土城子组之上发育的区域性不整合代表了另一次自北向南的推覆事件(Zhang *et al.*, 2008b),该期火山岩就是在这次推覆事件之后形成的,这表明该期火山岩的形成应与这次陆壳加厚时间的坍塌或拆沉阶段相对应——即伸展的构造背景(韦忠良等, 2008)。

5.6 早白垩世晚期(110~131Ma):区域伸展与双峰式火山岩形成

早白垩世晚期火山岩在东北地区广泛分布,尤其是大兴安岭地区,在其北部以上库力组流纹岩和伊列克得组玄武岩为代表,在南部以白音高老组流纹岩和梅勒图组玄武岩为代表,它们形成于114~131Ma之间(峰期年龄为125Ma;林强等, 1999; 葛文春等, 2001; Wang *et al.*, 2006; Zhang *et al.*, 2008a, 2010; Ying *et al.*, 2010; 孟恩等, 2011; 徐美君等, 2011)。在冀北-辽西地区,该期火山岩以义县组为代表,形成于120~126Ma(韦忠良等, 2008)。该期火山事件表现为典型的双峰式火山岩组合,揭示了区域性伸展环境的存在,这也得到了早白垩世晚期广泛发育的变质核杂岩、A型花岗岩和裂谷盆地的支持(Davis *et al.*, 2001; Wu *et al.*, 2002; Wang *et al.*, 2011b)。该期火山事件既可与古太平洋板块俯冲于欧亚大陆之下的弧后伸展环境相联系,又可与蒙古-鄂霍茨克缝合带闭合后加厚陆壳的拆沉过程相联系,对于大兴安岭地区早白垩世双峰式火山岩组合的形成应与后者关系更为密切(Xu *et al.*, 2013)。

5.7 蒙古-鄂霍茨克缝合带中生代的构造演化与区域成矿

从我国境内的大兴安岭北段、南段和冀北-辽西地区中生代岩浆作用的时空格架以及岩石组合特征可以看出,蒙古-鄂霍茨克缝合带中生代期间对我国的影响是强烈的,其大致可以划分成以下几个阶段:

1)早中生代(240~180Ma) 蒙古-鄂霍茨克大洋板块向

额尔古纳地块之下的俯冲阶段:依据该区花岗岩的岩浆作用期次与性质及其成矿特征(秦克章等, 1998; 陈志广等, 2010; Wu *et al.*, 2011; 王伟等, 2012; 余宏全等, 2012),大致可以划分成三次俯冲事件(245Ma ±、200Ma ± 和 185Ma ±),与之相对应的成矿作用主要表现为活动陆缘背景下的斑岩型(铜)钼和铜铅锌多金属矿床的形成(图4)。

2)中生代晚期(160~120Ma) 该阶段主要表现为两次陆壳加厚过程(170Ma左右和145Ma左右)和与之相对应的陆壳伸展环境(162Ma、142~120Ma),该期成矿作用主要表现为伸展背景下的浅成热液型金矿、多金属矿床和稀有、稀土金属矿床的形成(图5)(白令安等, 2012; 王建国等, 2013)。

6 古缝合带的再活化与区域成矿

东北地区处于中亚造山带的东段,该区古生代期间主要表现为多个微陆块之间的拼合过程和多岛弧的弧陆拼合过程(Sengör and Natalin, 1996; Li, 2006)。在古生代晚期-早中生代早期,该区经历了古亚洲洋的最终闭合,完成了中亚造山带的演化历程,之后进入到了蒙古-鄂霍茨克构造体系和环太平洋构造体系的叠加演化阶段,这些晚期的构造体系对古亚洲洋构造体系时期形成的古俯冲带(缝合带)有何影响?对区域成矿作用有何指示意义?这可从东北地区近年来不同类型矿床的发现得到回答。

6.1 环太平洋构造体系对古俯冲带(或缝合带)的影响与成矿

基于吉黑东部中生代火山岩的研究,可知环太平洋构造体系在欧亚大陆东北缘中生代的演化主要表现为三次重要的俯冲事件:即早-中侏罗世(173~190Ma)、早白垩世晚期(106~133Ma)和晚白垩世(88~97Ma)。而位于东北地区东部的古俯冲带主要有松嫩-张广才岭地块与佳木斯地块之间南北向展布的加里东期缝合带(Wang *et al.*, 2012)和位于华北地块北缘东西向展布的晚古生代缝合带(Zhang *et al.*, 2009; 曹花花等, 2012)。在古太平洋板块俯冲于欧亚大陆之下的过程中,这些古老的俯冲带再次复活,形成了具有活动陆缘构造背景以钼为主的斑岩型矿床,如小兴安岭-张广才岭早侏罗世斑岩型矿床(霍吉河钼矿, 184Ma, 陈静等, 2012; 鹿鸣斑岩型钼矿, 176Ma, 杨言辰等, 2012; 翠岭钼矿, 178Ma, 杨言辰等, 2012)的形成、华北地块北缘早-中侏罗世斑岩型矿床(大黑山斑岩型钼矿, 168Ma, 王成辉等, 2009; 福安堡斑岩型钼矿, 167Ma, 李立兴等, 2009)的形成都应古太平洋板块俯冲作用下的古缝合带活化有关(图4)。从目前对东北地区成矿年龄的统计可以看出,这些古俯冲带上的成矿作用主要发生在早-中侏罗世,而早白垩世晚期(106~133Ma)和晚白垩世(88~97Ma)两次俯冲作用对古俯冲带的影响较小,东部陆缘内一侧主要表现为类似弧后的伸展环境,而对应的成矿作用主要以富金斑岩型铜矿和浅成低温热

液型金矿为主(图5)(Han *et al.*, 2013)。

6.2 蒙古-鄂霍茨克构造体系对古俯冲带(或缝合带)的影响与成矿

6.2.1 中生代早期俯冲阶段

从中生代火山岩的空间展布可以判定,蒙古-鄂霍茨克构造体系对对我国的影响主要在松辽盆地以西地区(包括冀北-辽西地区)。中生代早期,蒙古-鄂霍茨克大洋板块俯冲于额尔古纳地块之下,其影响空间范围可以达到华北地块的北缘,其东西向展布的早侏罗世南大岭组玄武岩(赵越等, 2006)应是这次俯冲作用产生的远程效应,这主要表现为对华北地块北缘晚古生代缝合带的影响。此外,华北地块北缘东西向展布的早侏罗世以钼为主的斑岩型矿床的形成(如新台门钼矿, 183Ma, 张遵忠等, 2009; 蓝家沟钼矿, 187Ma, 黄典豪等, 1996; 杨家杖子钼矿, 187Ma, 黄典豪等, 1996)也可能与蒙古-鄂霍茨克大洋板块向额尔古纳地块下俯冲的远程效应有关(图4),位于华北陆块之下的古亚洲洋板块的复活可能是导致这些斑岩型矿床形成的主要动力学背景。

6.2.2 晚中生代蒙古-鄂霍茨克洋的闭合碰撞与伸展阶段

在中侏罗世(海房沟组之下的区域不整合)和早白垩世早期(土城子组之上的区域不整合)两期构造推覆事件(Zhang *et al.*, 2008b)对古俯冲带的影响主要表现为陆壳加厚过程的发生和古俯冲带的再活化,前者主要表现为adakitic火山岩的形成(Yang and Li, 2008),后者主要表现在华北地块北缘中-晚侏罗世砂卡岩型矿床的形成(如肖家营子钼矿, 170Ma, 代军治等, 2008; 姚家沟钼矿, 169Ma, 方俊钦等, 2012)。而区域成矿作用主要发生在加厚陆壳坍塌或拆沉阶段伸展背景下,以浅成低温热液金矿床为主,同时在古俯冲大洋板片存在的部位,具有稀有-稀土矿床的形成,比如与吉祥峰组(或玛尼吐组)A型流纹岩或碱性花岗岩有关的稀土矿产的形成,应是已经发生细碧-角斑岩化的古俯冲大洋板片,在伸展环境下部分熔融的产物(王建国等, 2013)。

7 结论

基于对东北地区中生代火山岩的年代学与岩石组合的时空变化规律研究,得出如下结论:

(1)东北地区中生代火山作用可划分成六期:晚三叠世(200~228Ma)、早-中侏罗世(173~190Ma)、晚侏罗世(158~166Ma)、早白垩世早期(138~145Ma)、早白垩世晚期(106~133Ma)和晚白垩世(88~97Ma)。

(2)环太平洋构造体系对欧亚大陆下的俯冲作用始于早侏罗世,中生代期间存在三次(早侏罗世、早白垩世晚期和晚白垩世)俯冲事件,其成矿作用主要表现为俯冲阶段的斑岩型矿床和陆内伸展背景下的浅成低温热液矿床,同时在早侏罗世俯冲阶段,对古俯冲带的改造与活化也是寻找斑岩型矿床的主要区域。

(3)蒙古-鄂霍茨克构造体系中生代的演化经历了中生代早期(245~180Ma)的俯冲作用和中生代晚期两次(~170Ma和~145Ma)重要的陆壳加厚过程和与之相对应的两次区域性的伸展事件(早白垩世早期:~142Ma和晚期:~125Ma)。在额尔古纳地块和兴安地块的北段是寻找与俯冲时间有关斑岩型矿床的主要区域,而在华北地块北缘以及大兴安岭南段是寻找伸展背景下浅成低温热液矿床和多金属矿床的重点区域。

(4)环太平洋构造体系中生代对东北亚大陆影响的空间范围主要在松辽盆地及以东地区;而蒙古-鄂霍茨克构造体系影响的空间范围主要在松辽盆地以西(可能包括松辽盆地的西部斜坡区)以及华北地块北缘。

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