# 宁夏泾源石咀子古元古代 A 型花岗岩的形成时代 及其地质意义<sup>\*</sup>

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Abstract The Shizuizi A-type granite is found in the southern part of Helan aulacogen, western margin of the North China Craton and mainly consist of K-feldspar granite porphyry. The granite is characterized by high silicon and rich potassium, including  $Na_2O + K_2O$  average 7. 61%,  $K_2O/Na_2O = 2.17 \sim 7.39$ ,  $Al_2O_3 = 10.59\% \sim 11.84\%$ , and its A/CNK = 0.86  $\sim 1.11(1.01$  on average), and thus is belong to subalkaline metaluminous to weakly prealuminous series, identified as A-type granite. The abundance of  $\Sigma$  REE is high, ranging from 340.4  $\times$  10<sup>-6</sup> to 468.9  $\times$  10<sup>-6</sup>. The granite show a LREE enrichment pattern and obviously intermediate negative Eu anomalies. The trace element geochemistry is characterized evidently by negative anomaly of high strength field elements Ta, Nb, Ti, etc. and positively Rb, Ba, Th, etc., and shows the characteristics of post-orogenic granite. The LA-ICP-MS zircon U-Pb age is 1803  $\pm$  15Ma, which indicates that the Helan aulacogen is formed during Late Paleoprotozoic, and is related to the break-up of the North China Craton from the Columbia supercontinent during the end of Late Paleoprotozoic period.

**Key words** Helan aulacogen; Paleoprotozoic; A-type granite; Break-up of continent

摘 要 对华北克拉通西缘贺兰坳拉谷南段泾源县石咀子花岗斑岩进行了岩石地球化学和年代学研究结果表明,石咀子花岗岩体具有高硅(SiO<sub>2</sub> = 72.28% ~ 76.69%)、富钾特征,Na<sub>2</sub>O + K<sub>2</sub>O 平均值 7.61%,K<sub>2</sub>O/Na<sub>2</sub>O 为 2.17 ~ 7.39,Al<sub>2</sub>O<sub>3</sub> = 10.59% ~ 11.84%,A/CNK 为 0.86 ~ 1.11(平均为 1.01),低钙镁,岩石为高硅准铝质-弱过铝质 A 型花岗岩;稀土元素总量较高,为 340.4 × 10<sup>-6</sup> ~ 468.9 × 10<sup>-6</sup>,轻稀土富集,具有中等的负 Eu 异常,配分曲线呈典型的右倾"海鸥型";高场强元素 Ta、Nb、Ti 具有明显的负异常,大离子亲石元素 Rb、Ba、Th 等相对富集,花岗岩具有造山后岩石地球化学特征。锆石的 LA-ICP-MS U-Pb 年龄为 1803 ± 15Ma,为古元古代晚期,表明贺兰坳拉谷形成于古元古代晚期,其形成与华北克拉通古元古代晚期大陆裂解过程有关。

关键词 贺兰坳拉谷;古元古代;A型花岗岩;大陆裂解

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### 1 引言

尽管对华北克拉通前寒武纪结晶基底的构造格局划分 及演化有不同认识,但大多数学者都承认在古元古晚期-中 元古代早期华北克拉通广泛发育拉张裂解事件,该事件是对 Columbia 超大陆裂解的强烈反映(翟明国和彭澎,2007; Zhao and Zhou,2009;陆松年等,2010)。华北克拉通在中元古代形 成了至少5个拗拉谷(或裂谷),包括燕辽-太行拗拉谷、白云 鄂博裂谷-渣尔泰裂谷、贺兰山拗拉谷、晋陕拗拉谷和熊耳-中 条拗拉谷(李江海等,2000),其中发育有大量不同类型、年龄 峰值为1.8~1.6Ga的火山岩组合,如华北克拉通南缘熊耳-中条拗拉谷发育了以熊耳群为主的火山岩系(赵太平等, 2004,2007;徐勇航等,2007),还发育有同期花岗闪长岩和斑 岩(崔敏利等,2010)及基性岩墙群(侯贵廷等,2010;胡国辉 等,2010;胡俊良等,2007),北部燕辽-太行拗拉谷发育大规 模基性岩墙群和碱性侵入岩(彭澎等,2004;任康绪等,2006; 阎国翰等,2007;杨进辉等,2007;董春艳等,2010)。对贺兰 山拗拉谷、晋陕拗拉谷的存在主要依据石油勘探地震资料和 少量露头沉积学证据:一是这两个裂谷地区发育巨厚的中元 古界地层,其中贺兰拗拉谷南段地层厚度可达 4000~ 5000m,且其沉积相序和岩性组合与燕辽拗拉谷的具有很强 的相似性(华洪和邱树玉,2001;高林志等,2010);二是在地 震剖面上可以看到元古界内发育众多的伸展断层,中元古代 地层厚度展布明显受同沉积期断裂控制,这与燕辽拗拉谷形 态类似(王同和,1995)。但是关于贺兰拗拉谷发育的确切岩石学和年代学证据目前还未见报道。我们在泾源县新民乡石咀子村附近发现了一处元古宙的花岗斑岩,并对其进行了LA-ICP-MS 锆石 U-Pb 定年和地球化学特征分析,从中获得与古元古代构造热事件相当的年龄值,从而为进一步认识华北克拉通西部贺兰拗拉谷的成因等地质问题提供了新的证据。

### 2 岩体地质概况

研究区处于中宁-海原-关山-宝鸡断裂以东的鄂尔多斯地块边缘区,为元古宙贺兰坳拉谷的南段。对元古宙贺兰坳拉谷的研究主要体现在两个方面,其一是古元古代变质岩的研究,贺兰山北段的孔兹岩系,形成时间是 1.9~2.1Ga(周喜文等,2010; 秋元生等,2010; 校培喜等,2011),与鄂尔多斯地块北缘的集宁岩群和乌拉山岩群十分相似,与吕梁构造运动的时限一致,为古元古代华北克拉通成熟陆壳演化阶段的产物;元古宙的侵入岩报道主要见于贺兰山坳拉谷北段地区,年龄均大于 1.85Ga(耿元生等,2010),为古元古代,而南段未见有元古宙侵入岩报导。宁夏地矿局在进行泾源县幅 1:5万填图中,在泾源县新民乡石咀子附近发现了一处花岗斑岩,未变质,坐标 35°21′04.5″,106°26′22″,采样点花岗岩体目估长约 300m,宽 50m,顶部具有清晰的风化面,上被古近系的清水营组或第四系马兰黄土不整合覆盖(图 1),现场初步定为喜山期侵入岩。岩石具有斑状结构,基质为细粒结构,

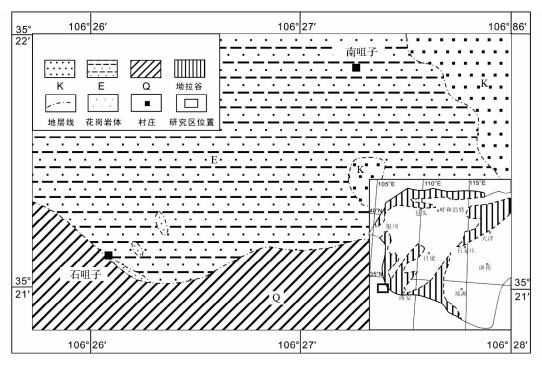


图 1 古元古代泾源花岗岩分布区地质略图(拗拉谷分布据侯贵廷等,2010 改编)

Fig. 1 Simplified geological map of the Jingyuan granite outcrop (aulacogen distribution map modified after Hou et al., 2010)

表 1 泾源花岗岩的主量元素 ( wt% ) 和微量元素 (  $\times$  10  $^{-6}$  ) 分析结果

Table 1 Major element (  $\rm wt\%$  ) and trace element (  $\rm \times\,10^{-6}$  ) compositions of the Shizuizi granites

Sample	jy-01	jy-02	jy-03	jy-04	jy-06	jy-07	jy-08	jy-12
${\rm SiO_2}$	74. 49	73. 42	72. 83	73.83	72.72	72. 28	76.65	76. 69
${\rm TiO_2}$	0.44	0.44	0.51	0.41	0.48	0.46	0.37	0.35
$\mathrm{Al}_2\mathrm{O}_3$	11. 22	11. 15	11.84	11. 29	10.96	10.78	10.59	10. 59
$\mathrm{Fe_2O_3}$	4.08	4. 67	4. 54	4. 26	4. 77	4. 84	3.56	2. 55
MnO	0.02	0.03	0.02	0.02	0.02	0.02	0.01	0.03
MgO	0.59	0.59	0.57	0.56	0.41	0.42	0. 19	0. 16
CaO	0.50	1.33	0.93	0.56	1.75	1.53	0. 17	0.84
$Na_2O$	1.89	1.84	1.90	2. 19	2. 38	1.87	0.95	1. 14
$K_2O$	5.68	5.06	5. 90	5. 38	5. 17	5.77	7.02	6. 77
$\mathrm{P}_2\mathrm{O}_5$	0.08	0.07	0.09	0.07	0.08	0.08	0.06	0.05
LOI	1.02	0.87	0.92	0.92	1.37	1.47	0.38	0.77
La	95.4	73.3	91.7	103.5	76. 3	71.5	79. 5	76. 3
Ce	193.0	139.9	202. 5	182. 8	150.3	134. 3	151.3	148. 2
Pr	22. 3	16. 62	20.77	22. 13	17.60	15. 52	17. 57	16. 94
Nd	90.6	68.6	84. 0	89. 3	71.4	64.8	70.6	68. 6
Sm	14. 94	12.72	14. 63	15.96	12. 95	11.73	13.07	12. 31
Eu	1.93	1.43	1.78	1.89	1.58	1.48	1. 27	1. 21
$\operatorname{Gd}$	14. 6	11.81	13.89	13.99	12.68	11.05	12. 32	11.68
Tb	2. 140	1. 93	2.08	2. 13	2.03	1.82	1.89	1.87
Dy	13.5	12. 14	13. 27	12.38	13. 10	11.33	12. 18	11. 23
Ho	2.66	2. 34	2. 56	2. 35	2. 47	2. 17	2. 25	2. 21
Er	7.44	6.78	8. 03	6.62	6. 98	6.35	6.84	6. 22
Tm	1. 26	1. 17	1.39	1.09	1. 23	1.08	1. 16	1.03
Yb	7.63	7. 07	8. 59	6.56	7. 30	6. 39	7. 19	5. 96
Lu	1.08	0.98	1. 21	0.95	1.07	0.91	0.93	0.82
Ba	1297	1272	1447	1293	1023	1037	1744	1558
$\operatorname{Sr}$	54. 30	125.4	75. 99	53.87	98.96	96.96	50. 26	61. 13
Ga	17. 4	18.88	21. 99	19. 20	16. 40	14. 38	16. 97	14. 94
Pb	5.63	4. 34	5. 45	4. 14	5. 37	4. 98	9. 15	12.05
Rb	124.0	119.7	132. 4	124. 9	100.4	97.61	176. 5	162. 8
Th	15.0	13.9	16. 3	13.4	14. 4	13.3	15.4	14. 2
U	2.31	2. 37	2.66	2. 21	2. 36	2. 39	1.93	2.09
Zr	463	420	4940	409	426	380	407	382
Nb	17. 6	16. 26	19. 57	15.66	17.41	15. 44	14. 98	14. 50
Hf	12. 9	11. 79	13. 65	11.52	12. 33	11.48	12.41	11.78
Ta	0.99	0.97	1.09	0.95	0.99	0.85	0.91	0.86
Y	59. 0	51.7	51. 1	49. 4	55. 1	49.8	54. 3	48. 9

斑晶成分主要为石英、钾长石、酸性斜长石,大小1~15mm不等,石英具有港湾状熔蚀,约占10%,长石呈板状或柱状,以钾长石为主,占8%,长石斑晶具有少量的高岭石、绢云母化现象,基质由细粒的钾长石、酸性斜长石、石英等矿物的集合体组成,蚀变中等到强,微观定名为花岗斑岩。

# 3 岩石化学特征

主微量元素分析均在中国科学院地质与地球物理研究 所进行,测试流程见陈志广等(2010)。从表 1 可以看出,样

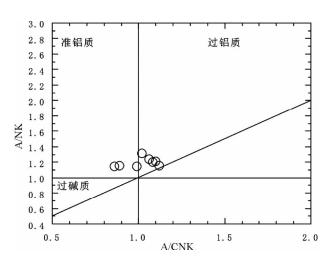


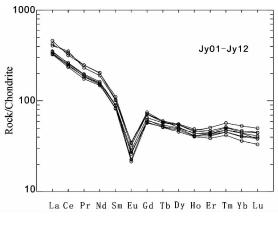
图 2 泾源花岗岩 A/NK-A/CNK 图解

Fig. 2  $\,$  A/NK-A/CNK diagram of the Jingyuan granites

品具有高硅(SiO<sub>2</sub> = 72.28% ~ 76.69%)特征,相对富铝, Al<sub>2</sub>O<sub>3</sub> = 10.59% ~ 11.84%, 岩石的 Na<sub>2</sub>O 和 K<sub>2</sub>O 含量高, Na,O+K,O 平均值 7.61%,且 K,O/Na,O 为 2.17~7.39,富 钾是其最显著的特点之一,在SiO2-K2O图解中落入钾玄岩 系列,在TAS图解中落入亚碱性花岗岩区域;样品的MgO、 TiO,和P,O,含量相对低(MgO=0.16%~0.59%,TiO,= 0.35% ~ 0.51%, P<sub>2</sub>O<sub>5</sub> = 0.05% ~ 0.09%)。样品的 MgO、 TiO<sub>2</sub>、P<sub>2</sub>O<sub>5</sub>、Sr、Zr等均与SiO<sub>2</sub>呈负相关关系,表明随着岩浆 演化可能存在斜长石、Fe-Ti 氧化物和磷灰石等的分离结晶。 岩石铝饱和指数 A/CNK 值(0.86~1.11),多数小于 1.1,岩 石的碱铝指数 AKI 介于 0.76~0.87, 在 A/CNK-A/NK 关系 图上,样品投点在亚碱性准铝质-弱过铝质区(图2),CIPW 标准矿物中多有刚玉(C)分子出现,属铝质花岗岩;碱度率 指数(AR)为3.47~6.71,平均4.65,在AR-SiO2 碱度关系图 中样品投点全部落在碱性区,但标准矿物中未见霞石(Ne)、 白榴石(Lc)、霓石(Ac)等碱性标准矿物,反映出该岩体岩石 的碱质并不过剩,总体显示了富硅钾低钙镁弱碱性花岗岩 特征。

### 4 稀土元素和微量元素特征

花岗岩样品稀土元素总量( $\Sigma$ REE)较高(表 1),为 340.4×10<sup>-6</sup>~468.9×10<sup>-6</sup>,远高于 I 型和 S 型的稀土总量 114.7×10<sup>-6</sup>和 173.1×10<sup>-6</sup>(吴锁平等,2007),轻重稀土元素比值(LREE/HREE)为 7.04~9.02,(La/Yb)<sub>N</sub>为 7.44~11.32,相对富集轻稀土。δEu 为 0.11~0.51,Eu 亏损显著;Ce 含量较高,均值达 163×10<sup>-6</sup>。(La/Sm)<sub>N</sub>和(Gd/Yb)<sub>N</sub>值分别为 3.72~4.19和1.34~1.62,轻重稀土元素分馏明显,所有样品均表现出轻稀土比重稀土分馏更显著,稀土元素球粒陨石标准化分布型式呈典型的 A 型花岗岩"海鸥型"样式(图 3)。



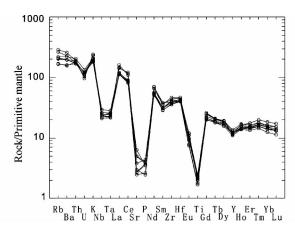


图 3 泾源花岗岩稀土元素配分曲线和微量元素配蛛网图(标准化值据 Sun and McDonough, 1989)

Fig. 3 Chondrite-normalized REE distribution patterns and primitive mantle-normalized spidergrams of Jingyuan granite samples (normalization values after Sun and McDonough, 1989)

微量元素中大离子亲石元素 Rb(97.6×10<sup>-6</sup>~176.5×  $10^{-6}$ ) Ba(  $1022 \times 10^{-6} \sim 1743 \times 10^{-6}$ ) Th(  $13.32 \times 10^{-6} \sim$ 16.33×10<sup>-6</sup>)、U(1.93×10<sup>-6</sup>~2.66×10<sup>-6</sup>)等含量高(表 1),明显富集 Rb、Ba、Th、K 等大离子亲石元素(LILE),在原 始地幔的标准化蛛网图中(图3)亏损高场强元素,Nb、Ta、Ti (HFSE)、P、Sr、Eu 十分明显负异常,相对富集 Zr、Hf, Zr 达 380×10<sup>-6</sup>以上,这些特征与扬子地块北缘华山观元古代环 斑花岗岩(张丽娟等,2011)、吕梁地区晚元古代晚期造山后 花岗岩(耿元生等,2004)、柴北缘元古界环斑花岗岩(胡能高 等,2007)的特征接近。Sr、Eu 强烈负异常表明岩浆发生了 明显的长石分离结晶,P和Ti亏损说明岩浆经历了磷灰石以 及钛铁矿等矿物的分离结晶作用, Nb、Ta 的亏损可能与岩浆 源区岩石中陆壳组分的参与有关(Kalsbeek at al., 2001)。 (Rb/Nb)<sub>N</sub> 比值为 6.5~13.2, 明显高于大陆壳的(Rb/Nb)<sub>N</sub> 比值(2.2~4.7),暗示着陆壳物质的贡献较大,导致 Rb 的含 量相对增大。相容组分 Cr、Co、Ni 含量较低,分别为 2.97 ×  $10^{-6} \sim 14.69 \times 10^{-6}$ ,  $1.78 \times 10^{-6} \sim 3.66 \times 10^{-6}$ ,  $2.54 \times 10^{-6}$ ~4.06×10<sup>-6</sup>, Mg<sup>#</sup>相对也低(9.64~22.44, 平均16.72), 这 显示出岩浆地壳来源而非地幔源的特征。

# 5 锆石 U-Pb 同位素年龄

锆石原位 U-Pb 同位素年龄分析在西北大学大陆动力学 国家重点实验室的 LA-ICP-MS 仪器上用标准测定程序进行, 详细的实验原理和流程及仪器参数见袁洪林等(2003)。岩 石锆石 U-Pb 年龄测试结果见(表2),用 LAM-ICPMS Common Lead Correction (ver3. 15)对其进行了普通铅校正,年龄计算 及谐和图采用 Isoplot (ver3. 0)完成。

泾源花岗岩体中锆石为无色透明或浅黄色,结晶较好, 呈典型的长柱状晶形。锆石的阴极发光图像均具有清晰的 震荡环带结构,显示岩浆成因锆石特征。锆石的 Th/U 比值

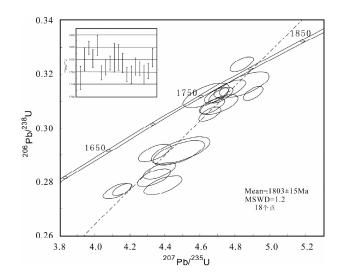


图 4 泾源花岗岩的锆石 U-Pb 谐和图

Fig. 4 U-Pb concordia diagram of zircons from Jingyuan granite

变化范围在  $0.26 \sim 0.96$  之间,均大于 0.1,也指示其岩浆成因。锆石的 LA-ICP-MS U-Pb 定年结果见表 2。在 24 次分析结果中,有 18 个测点构成条不一致曲线,与谐和线的上交点为  $1804 \pm 26$  Ma,其 $^{207}$  Pb/ $^{206}$  Pb 加权平均年龄为  $1803 \pm 15$  Ma  $(2\sigma,MSWD=1.2)$ (图 4),两个年龄在误差范围内一致,以加权平均年龄  $1803 \pm 15$  Ma 代表泾源花岗岩体的结晶年龄,为古元古代晚期岩浆活动之产物。

# 5 岩石成因与构造环境探讨

#### 6.1 岩石成因

岩石的 Ga 含量较高,平均  $17.5 \times 10^{-6}$ ,  $10000 \times Ga/Al$  值为  $2.52 \sim 3.51$ ,除一个数据外其余均大于 A 型花岗岩的下限值 2.6,明显高于 I 型和 S 型花岗岩的平均值(分别为 2.1 和 2.28)(Whalen *et al.*, 1987),在 10000Ga/Al-Ce 和 10000Ga/

### 表 2 泾源花岗岩锆石 LA-ICPMS U-Pb 定年数据

Table 2 The LA-ICP-MS zircon U-Pb isotopic data of the Shizuizi granite

	元素含量(×10 <sup>-6</sup> )				视年龄(Ma)				
测点号	IJ	Th	U/Th	$^{207}\mathrm{Pb}/^{206}\mathrm{Pb}$	$^{207}\mathrm{Pb}/^{235}\mathrm{U}$	$^{206}\mathrm{Pb}/^{238}\mathrm{U}$	$^{207}\mathrm{Pb}/^{206}\mathrm{Pb}$	$^{207}\mathrm{Pb}/^{235}\mathrm{U}$	$^{206}{\rm Pb}/^{238}{\rm U}$
	U	111		±1 $\sigma$	$\pm 1\sigma$	±1 $\sigma$	$\pm 1\sigma$	$\pm 1\sigma$	$\pm 1\sigma$
01	53. 8	40. 6	1. 32	0. 1161 ± 0. 0024	4. 5836 ± 0. 0882	0. 2864 ± 0. 0027	1897 ± 37	1746 ± 16	1623 ± 14
02	61.6	53.8	1. 15	$0.1077 \pm 0.0022$	$4.6343 \pm 0.0874$	$0.3122 \pm 0.0029$	$1760 \pm 37$	$1756 \pm 16$	$1751 \pm 14$
04	145. 6	125. 1	1. 16	$0.1146 \pm 0.0025$	$4.8177 \pm 0.0977$	$0.3050 \pm 0.0031$	$1873 \pm 39$	$1788 \pm 17$	$1716 \pm 16$
05	58.8	45.6	1. 29	$0.1246 \pm 0.0032$	4. 8457 ± 0. 1199	$0.2821 \pm 0.0035$	$2023 \pm 45$	$1793 \pm 21$	$1602 \pm 17$
06	70. 9	59. 3	1. 20	0. $1113 \pm 0.0023$	$4.8523 \pm 0.0931$	$0.3163 \pm 0.0031$	$1820 \pm 37$	1794 ± 16	$1772 \pm 15$
07	203.0	183.8	1. 10	$0.1136 \pm 0.0013$	$4.9024 \pm 0.0369$	$0.3130 \pm 0.0016$	$1858 \pm 21$	$1803 \pm 6$	$1755 \pm 8$
08	91. 2	85. 6	1.07	$0.1114 \pm 0.0018$	$4.3677 \pm 0.0599$	$0.2842 \pm 0.0021$	$1823 \pm 29$	$1706 \pm 11$	$1613 \pm 10$
09	75.7	62. 1	1. 22	$0.1141 \pm 0.0021$	$4.3879 \pm 0.0717$	$0.2788 \pm 0.0023$	$1866 \pm 33$	$1710 \pm 14$	$1585 \pm 12$
10	398. 7	371.4	1.07	$0.1087 \pm 0.0014$	4. $1541 \pm 0.0357$	$0.2772 \pm 0.0015$	$1778 \pm 23$	$1665 \pm 7$	$1577 \pm 8$
11	131.2	117. 9	1. 11	$0.1098 \pm 0.0015$	$4.6674 \pm 0.0480$	$0.3084 \pm 0.0019$	$1795 \pm 25$	$1761 \pm 9$	$1733 \pm 9$
12	227. 2	218. 1	1.04	$0.1106 \pm 0.0014$	$4.6548 \pm 0.0414$	$0.3053 \pm 0.0017$	$1809 \pm 23$	$1759 \pm 7$	$1718 \pm 9$
13	42.0	24. 5	1.71	$0.1116 \pm 0.0029$	$4.4898 \pm 0.1103$	$0.2918 \pm 0.0034$	$1826 \pm 46$	$1729 \pm 20$	$1651 \pm 17$
14	57. 5	42. 4	1.36	$0.1116 \pm 0.0026$	$4.4790 \pm 0.0964$	$0.2910 \pm 0.0030$	$1826 \pm 41$	$1727 \pm 18$	$1647 \pm 15$
15	78. 7	58. 0	1.36	$0.1106 \pm 0.0019$	$4.7091 \pm 0.0706$	$0.3089 \pm 0.0024$	$1809 \pm 31$	$1769 \pm 13$	$1735 \pm 12$
16	128.8	107. 1	1. 20	$0.1083 \pm 0.0014$	$4.8374 \pm 0.0443$	$0.3239 \pm 0.0018$	$1772 \pm 23$	$1791 \pm 8$	$1809 \pm 9$
17	154. 4	142.7	1.08	$0.1082 \pm 0.0017$	$4.1294 \pm 0.0533$	$0.2767 \pm 0.0019$	$1770 \pm 28$	$1660 \pm 11$	$1575 \pm 10$
18	88. 0	70.6	1. 25	$0.1099 \pm 0.0017$	$4.7691 \pm 0.0600$	$0.3147 \pm 0.0022$	$1798 \pm 28$	$1780 \pm 11$	$1764 \pm 11$
19	345. 2	89. 0	3.88	$0.1214 \pm 0.0015$	$5.2297 \pm 0.0459$	$0.3125 \pm 0.0018$	$1977 \pm 22$	$1858 \pm 7$	$1753 \pm 9$
20	453.0	162. 8	2.78	$0.1218 \pm 0.0016$	$5.6406 \pm 0.0568$	$0.3360 \pm 0.0021$	$1982 \pm 24$	$1922 \pm 9$	$1867 \pm 10$
21	194. 7	178. 6	1.09	$0.1096 \pm 0.0014$	$4.7178 \pm 0.0450$	$0.3123 \pm 0.0018$	$1792 \pm 24$	$1770 \pm 8$	$1752 \pm 9$
22	73.6	59.6	1. 23	$0.1207 \pm 0.0035$	$4.0854 \pm 0.1131$	$0.2455 \pm 0.0032$	$1967 \pm 51$	$1651 \pm 23$	$1415 \pm 17$
23	131.3	106. 9	1. 23	$0.1082 \pm 0.0018$	4. $3418 \pm 0.0609$	$0.2910 \pm 0.0022$	$1770 \pm 30$	$1701 \pm 12$	$1646 \pm 11$
24	122. 5	91.3	1.34	$0.1090 \pm 0.0014$	$4.6932 \pm 0.0457$	$0.3123 \pm 0.0019$	$1783 \pm 24$	$1766 \pm 8$	$1752 \pm 9$
25	117.7	108. 3	1.09	$0.1116 \pm 0.0019$	$4.3368 \pm 0.0625$	$0.2818 \pm 0.0021$	$1826 \pm 30$	$1700 \pm 12$	$1600 \pm 11$

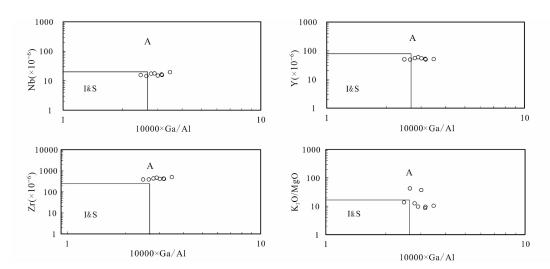


图 5 A 型花岗岩判别图解(据 Whalen et al., 1987)

Fig. 5 Discriminant diagram of A-type granitoid (after Whalen et al., 1987)

Al-Zr 等图解上,都位于 A 型花岗岩区域(图 5)。HFSE 元素含量高,元素组合 Zr + Nb + Ce + Y =  $580 \times 10^{-6} \sim 767 \times 10^{-6}$ 高于 A 型花岗岩下限值( $350 \times 10^{-6}$ )花岗岩。泾源花岗岩的 Y/Nb 比值为 2. 61 ~ 3. 62,大于 1. 2,具 A<sub>2</sub> 型花岗岩特征(贾小辉等, 2009), Y-Nb-Ce 和 Y-Nb-3Ga 花岗岩类型判别图解

上(图 6),都表现为  $A_2$  型花岗岩的特征。泾源花岗岩 Rb/Sr 比值范围在  $0.95 \sim 2.66$ ,平均为 1.94,远高于原始地幔平均值(0.03),说明岩浆的演化程度高。La/Nb 为  $4.38 \sim 6.61$ ,远远大于 1,与地幔来源的岩浆差异明显(DePaolo and Daley, 2000),Th/U 比值( $6.56 \sim 7.95$ )平均 6.37,接近下地壳值

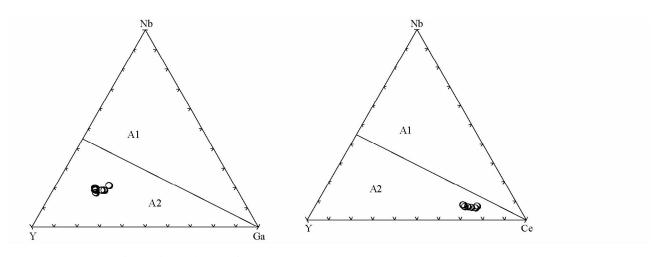


图 6 A<sub>1</sub> 和 A<sub>2</sub> 型花岗岩亚类的三角形判别图解(据 Eby, 1992)

Fig. 6 Discrimination diagrams of the Jingyuan A-type granitoid (after Eby, 1992)

(≈6)(Rudnick and Gao,2003),在 Nb/Y-Rb/Y 图解中样品点集中于下地壳平均线附近,La/Sm-La 图解显示泾源花岗岩由原岩的部分熔融形成,没有经历明显的分离结晶作用,暗示着该花岗岩可能主要由陆壳的部分熔融所形成。由于还缺乏详细的 Sr-Nd 同位素资料,我们目前还不能对泾源花岗岩体的源区以及成因模式进行进一步的限定。

#### 6.2 构造动力学背景及其对贺兰坳拉谷的制约

目前为止学者对 A 型花岗岩的成因的观点还不很一致, 但大多数学者认为 A 型花岗岩形成于造山后或非造山的岩 石圈伸展构造环境(Eby, 1992; Whalen et al., 1996; Forster et al., 1999; Kusky and Li, 2003; Kalsbeek F, 2001; 靳松等, 2010;杨崇辉等,2011),A型花岗岩的形成与地壳的拉张作 用有关,岩浆可能起源于经历了陆-陆碰撞或岛弧岩浆作用 的陆壳或板下地壳,为造山作用结束后不久即开始的拉张。 在SiO<sub>2</sub>与Al<sub>2</sub>O<sub>3</sub>构造环境判别图上投点到RRG+CEUG范 围内,表明与伸展有关的环境;在Rb-Y和Nb-Y环境判别图 上,样品点都落入板内或后碰撞范围内,在 R<sub>1</sub>-R<sub>2</sub> 构造环境 图解上上投点全部位于造山后区域(图略),这个与河北赞皇 地区许亭元古代型花岗岩(杨崇辉等,2011)、嵩山古元古代 A型花岗岩(Zhao and Zhou, 2009)、官昌圈椅埫古元古代 A 型花岗岩(熊庆等,2008)以及吕梁云中山古元古代花岗岩 (耿元生等,2004)特征非常相似,总体表现了造山期后花岗 岩特点。

在华北克拉通的北缘以麻粒岩相、孔兹岩系及造山型花岗岩为标志的造山带也形成近于同一时期(郭敬辉等,2002, Zhao et al.,2005;赵国春等,2009;钟长汀等,2007),华北南部同样也发育该时期的造山事件,北秦岭西段陇山杂岩、胡店和太白山片麻状花岗岩(何艳红等,2005;王洪亮等,2007;何世平等,2007;第五春荣等,2010;时毓等,2011)等1.9~1.8Ga的透入性变质事件同样指示了古元古代晚期的造山

事件;华北克拉通西部贺兰山、阿拉善地区广泛发育的孔兹 岩系与伴生花岗岩(耿元生等,2009;周喜文等,2010;校培喜 等,2011),指示了鄂尔多斯地块与阴山地块、阿拉善地块古 元古代经历了强烈的造山过程,充分表明了华北克拉通周缘 古元古代造山事件具有大面积区域性特征(翟明国等, 2007)。对于华北克拉通古元古代晚期造山后伸展事件与坳 拉谷的形成前人已有大量的论述,华北克拉通北缘的斜长岩 环斑花岗岩非造山岩浆组合(杨进辉等,2005)、京晋冀辽地 区古元古代晚期花岗岩组合(耿元生等,2004;杨进辉等, 2007)、华北克拉通中部镁铁质岩墙群(胡俊良等,2007)、华 北克拉通南缘偏碱性的熊耳群火山岩系以及碱性花岗岩组 合(王团华等,2008;包志伟等,2009;崔敏利等,2010)都表明 古元古代晚期地球动力学背景为区域性伸展环境,形成了华 北克拉通的多个坳拉谷,并认为是与哥伦比亚超大陆裂解有 关的全球构造在中国的集中体现。目前华北克拉通该期事 件的岩石学证据与论述多集中于中部与南北两侧,但贺兰坳 拉谷前人没有提出相应的岩石学证据。前已论述,贺兰坳拉 谷发育区无论南北均已发现古元古代造山之岩石证据,在古 元古代晚期在整个克拉通地壳-岩石圈减薄及软流圈地慢上 涌有关的背景下,贺兰-六盘山地区也难以独善其身,应该发 育有伸展构造背景下可以区域性对比的沉积序列和相应的 岩浆岩组合,泾源地区古元古代晚期 A 型花岗岩的发现为贺 兰坳拉谷的形成提供了有力的岩石学和年代学证据。

# 7 结论

- (1)精确的锆石 LA-ICP-MS U-Pb 年龄表明华北克拉通 西缘贺兰坳拉谷南段泾源花岗岩体年龄为 1803 ± 15Ma,属 古元古代而非喜山期。
- (2)地球化学特征研究表明泾源花岗岩为 A<sub>2</sub> 型花岗岩, 为华北克拉通古元古代区域性伸展背景下地壳部分熔融之

产物,为元古宙贺兰坳拉谷的成因提供了充分的岩石学证据。

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