

滇西澜沧江构造带中段新生代多期构造变形特征及时代约束*

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Abstract The Lancang tectonic zone is a large strike-slip shearing zone in the southeastern Tibet, which is the boundary between the Lanping-Simao block in the east and the Baoshan-Qiangtang block in the west. There are early oblique stretching lineations with fabrics showing dextral ductile deformation and late horizontal stretching lineations with fabrics showing sinistral brittle-ductile deformation in the middle segment of the Lancang tectonic zone. C-axis LPO patterns of quartz indicate the principal slip systems is prism $\langle a \rangle$ slip in medium temperature (450 ~ 600°C) relating with dextral ductile deformation, and superimposed by basal $\langle a \rangle$ slip in low temperature (300 ~ 550°C) relating with sinistral brittle-ductile deformation. The biotite ^{40}Ar - ^{39}Ar dating is ranging from 15Ma to 17Ma, which represent the change age from dextral ductile to sinistral brittle-ductile deformation according to the closure temperature of the biotite, that is the same as the temperature of the sinistral brittle-ductile deformation. The early dextral oblique strike-slip shearing was related with the extrusion of the Lanping-Simao block during the Early Miocene, and the sinistral horizontal brittle-ductile deformation was formed by the shearing of the southern part of the Baoshan block toward NEE along the NE Nandinghe and Wanding sinistral slip faults during exhumation of the ductile shear zone.

Key words Ductile deformation; ^{40}Ar - ^{39}Ar isotope chronology; Lancang tectonic zone; Western Yunnan

摘要 澜沧江构造带是青藏高原东南缘保山-羌塘地块与兰坪-思茅地块之间的大型走滑剪切带。构造地质学、岩石学和 ^{40}Ar - ^{39}Ar 年代学研究结果表明构造带中段剪切带内部存在早期斜向挤出和晚期水平走滑的两期线理及早期指示右行韧性剪切变形、后期指示左行脆韧性剪切变形的构造指向。糜棱岩中石英晶格优选方位以中温(450 ~ 600°C)柱面 $\langle a \rangle$ 轴底面滑移系为主,叠加中低温(300 ~ 550°C)底面 $\langle a \rangle$ 轴滑移系;剪切带内云母片岩和花岗质糜棱岩中黑云母 ^{40}Ar - ^{39}Ar 坪年龄和等时线年龄都分布于15 ~ 17Ma,反映剪切带隆升过程中脆韧性左行剪切变形阶段的时代。结合前人成果进行分析认为新生代早期保山地块和兰坪-思茅地块向南南东挤出的同时沿澜沧江构造带中段发生大规模右行斜向走滑韧性剪切作用,后期保山地块南部沿北东向畹町和南汀河左行断裂带相对中北段向北东运动,致使隆升到中上构造层次的韧性剪切带发生左行脆韧性变形。

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关键词 韧性变形; ⁴⁰Ar-³⁹Ar 年代学; 澜沧江构造带; 滇西
中图法分类号 P542; P597.3

1 引言

青藏高原东南缘三江构造带是冈瓦纳古陆东北缘块体相继于中晚泥盆世、二叠世和三叠世裂离、向北运移,块体之间的古特提斯洋(澜沧江洋、哀牢山-金沙江洋)、中特提斯洋(班公湖-怒江洋)和新特提斯洋(雅鲁藏布-密支那洋)相继俯冲(Metcalf, 2011, 2013),印支(兰坪-思茅)地块、Sibumasu(保山-羌塘)地块和腾冲-缅甸地块分别于印支期和燕山晚期相互拼合(Şengör, 1984; Dewey et al., 1988; Şengör et al., 1993; 从柏林等, 1993; 莫宣学等, 1993; 钟大赉, 1998; Yin and Harrison, 2000; 李兴振等, 2002; 戚学祥等, 2010a, b, 2012)形成的由多个地块(如腾冲-缅甸、Sibumasu/保山-羌塘和印支)以及块体之间残余缝合带(如高黎贡缝合带、澜沧江缝合带和哀牢山-金沙江缝合带)组成的复杂构造带(Mitchell, 1993; Mitchell et al., 2004; Morley, 2004; Metcalfe, 2006; Acharya, 2007; Hepepe et al., 2007; 杨经绥等, 2012)。在喜山运动期间,印度板块向北俯冲碰

撞导致青藏高原东南缘块体发生大规模旋转和逃逸(Tapponnier and Molnar, 1976; Tapponnier et al., 1982, 1990),块体沿古缝合带发生大规模走滑(丁林, 1991; 钟大赉等, 1991; Leloup et al., 1993, 1995; 钟大赉, 1998; 刘俊来等, 2006, 2007, 2011; 罗照华等, 2006; Morkey, 2007; Searle et al., 2007; 王二七等, 2006; Cao et al., 2010)形成由总体走向近南北,向北收敛、向南撒开的三条大型构造带和三个地块组成的构造格局(图 1a)。前人对东部哀牢山-金沙江构造带和西部高黎贡构造带中的大规模类型走滑剪切作用进行了广泛的研究,认为前者为左行大型韧性走滑剪切带,主要形成于 35 ~ 21Ma (Schärer et al., 1990, 1994; Harrison et al., 1992, 1996; Leloup et al., 1993, 1995, 2001; Zhang and Schärer, 1999; Gilley et al., 2003; Searle et al., 2010; Li et al., 2014),后者为大型右行韧性走滑剪切带,形成于 38 ~ 10Ma(钟大赉等, 1991; 季建清等, 2000; Lee et al., 2003; Wang et al., 2006; Lin et al., 2009; 李再会等, 2012; Zhang et al., 2012b; 王丹丹等, 2013)。但对分隔保山-羌塘地块和兰坪-思茅地

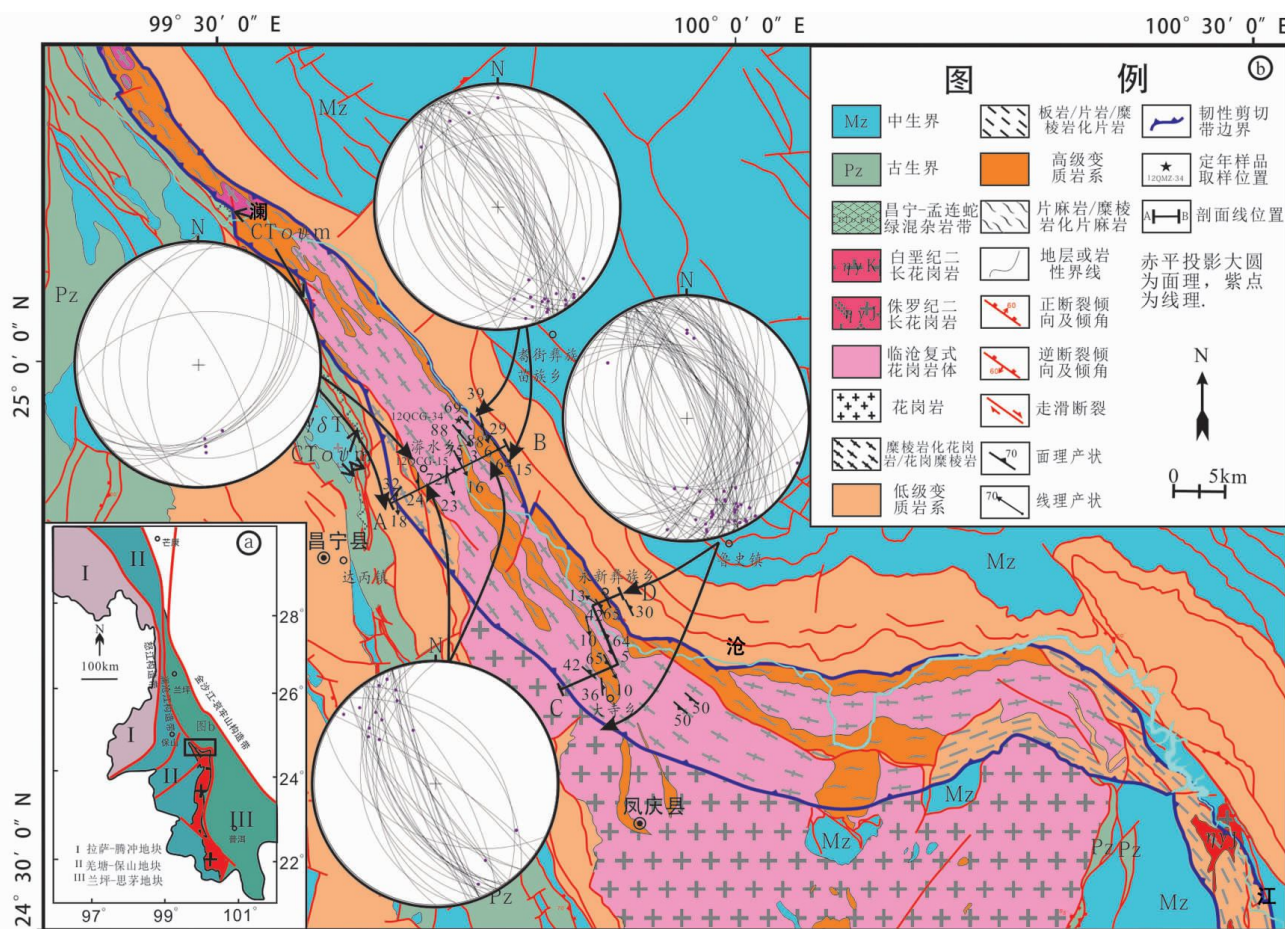


图 1 澜沧江构造带中段地质简图(a,据 Wang and Burchfiel, 1997 修编;b,云南省地质调查院,2011^①修编)
Fig. 1 Simplified geological map of the middle segment of the Lancang tectonic zone (a, modified after Wang and Burchfiel, 1997)

① 云南省地质调查院. 2011. 1 : 250000 凤庆幅地质图

块的边界性澜沧江构造带的研究相对薄弱,且主要集中于北段的碧罗雪山至崇山一带,认为碧罗雪山以北以右行走滑为主,碧罗雪山-崇山段以左行走滑为特征,其走滑时限为 37 ~ 21Ma (陈新跃等, 2006; Wang *et al.*, 2006; Akciz *et al.*, 2008; 张波, 2008; 张波等, 2009, 2011; Zhang *et al.*, 2010, 2012a; 唐渊等, 2013), 澜沧江构造带中段(崇山-临沧岩体段)是澜沧江构造带转折段,兰坪-思茅盆地在此段紧缩呈“蜂腰”状,其构造变形特征、形成时限及其与块体的挤出关系仍未查明。近年来,笔者对中段韧性变形特征进行了野外调研,发现该段同时存在典型的左行和右行走滑剪切标志及两组拉伸线理交切的现象。为此,在野外调研的基础上,通过显微构造学、岩石学和同位素年代学等方法,对该段韧性变形特征,左行和右行走滑作用的时空关系及其形成时代进行研究,并结合前人研究成果,揭示澜沧江构造带韧性变形的形成机制及其与块体旋转挤出的关系,为进一步探讨三江构造带的形成演化提供依据。

2 区域地质概况

澜沧江构造带位于青藏高原东南缘,北与高黎贡和金沙江构造带平行展布,沿碧罗雪山和崇山山脉向南延入临沧岩体,沿临沧岩体东侧直至进入泰国境内,长约 1300km (图 1a)。宽 3 ~ 20km,东部以澜沧江逆冲断层带与兰坪-思茅地块为界,西部以昌宁-凤庆逆冲断裂带分隔于保山地块。澜沧江构造带云南境内总体可分为三部分:北段为碧罗雪山以北向南延伸至崇山山脉,南段自以临沧岩体经景洪向南进入泰国,南北两段之间崇山南部至临沧岩体北部构成构造带的中段,即构造带从近南北走向转向南东东走向后又转向近南北向的过渡地段(图 1)。北段长约 350km,宽 3 ~ 20km,走向近南北,由西侧角闪岩相条带状片麻岩、糜棱岩化片麻岩、花岗质糜棱岩和东侧绿片岩相云母片岩、千枚岩和板岩组成。角闪岩相变质岩糜棱岩化变形明显,糜棱面理和拉伸线理发育,面理总体走向近南北向,倾向、倾角变化较大(倾角一般 $>60^\circ$),两侧倾向相反,呈正花状,线理向南倾伏,倾角较小($<30^\circ$,一般 $<20^\circ$)。指向构造发育,碧罗雪山以北以右旋为主,以南以左旋为主(Akciz *et al.*, 2008; 张波, 2008; 张波等, 2009, 2011; Zhang *et al.*, 2010, 2012; 唐渊等, 2013)。南段由西部的高压蓝片岩带、中部临沧岩体和东部的韧性走滑剪切带构成。其中,东部韧性走滑剪切带宽度变化较大,不同地段糜棱岩化程度不同,多由糜棱岩或初糜棱岩组成。中段长约 100km,宽度 3 ~ 20km,整体走向 NW-SE,于昌宁-凤庆-云县一带转为近 E-W 走向(图 1b),主体由高绿片岩相-角闪岩相片麻岩、混合岩和石榴云母片岩组成。带内岩石糜棱岩化明显,糜棱面理和拉伸线理发育,产状变化较大,总体倾向南西,指向构造揭示其存在典型的右行和左行两种走滑运动性质。澜沧江构造带东侧的兰坪-思茅地块内出露的地层为侏罗系-白垩系陆相沉积(红层),内部褶

皱、逆冲推覆构造发育。构造带西侧保山-羌塘地块内出露的地层为古生界被动陆缘沉积地层和零星分布的中生界沉积地层,大型北东向左行走滑断裂带南汀河断裂带和畹町断裂带将地块内部分为三部分(图 1a),南汀河断裂带北起云县茂兰、向南汇入实皆断裂,未切穿澜沧江构造带(朱玉新等, 1994; 王晋南等, 2006);畹町断裂带向北至湾甸汇入柯街断裂,向南汇入实皆断裂(常祖峰等, 2012; Morley, 2007)。

3 测试方法

3.1 EBSD 组构

EBSD 组构分析由中国地质科学院地质研究所大陆构造与动力学国家重点实验室完成。定向样品沿垂直面理并平行于线理方向切割光薄片,薄片抛光、喷碳后进行 EBSD 测量,测试仪器为日本电子公司(JEOL)制造的 JSM-56101V 型扫描电镜和丹麦 HKL 公司制造的 CHANNEL5 型号 EBSD 仪器。加速电压为 20kV,工作距离为 20mm,将衍射点数进行等面积下半球赤平投影,得到反映晶体 CPO 的各主要晶轴空间分布的极密图。组构图解的坐标轴设置为 X 轴平行于拉伸线理的方向,XY 为面理方向,Z 轴为垂直面理的方向,晶格优选方位极密图统计由 Channel 5 软件完成,详细实验流程参考文献(许志琴等, 2009)。

3.2 $^{40}\text{Ar}/^{39}\text{Ar}$ 年代学

黑云母 $^{40}\text{Ar}/^{39}\text{Ar}$ 定年由 中国地质科学院地质研究所同位素热年代学实验室完成。样品破碎后经重液浮选和磁选,选出黑云母初级样品,然后在双目镜下经手工仔细挑选,样品纯度在 99% 以上。选纯的黑云母单矿物(纯度 $>99\%$)用超声波清洗。清洗后的样品被封进石英瓶中送核反应堆中接受中子照射。照射工作是在中国原子能科学研究所的“游泳池堆”中进行的,使用 B4 孔道,中子流密度约为 $2.65 \times 10^{13} \text{ n} \cdot \text{cm}^{-2} \text{ S}^{-1}$ 。照射总时间为 1444min,积分中子通量为 $2.30 \times 10^{18} \text{ n} \cdot \text{cm}^{-2}$;同期接受中子照射的还有用做监控样的标准样:ZBH-25 黑云母标样,其标准年龄为 $132.7 \pm 1.2 \text{ Ma}$,K 含量为 7.6%。

样品的阶段升温加热使用石墨炉,每一个阶段加热 30min,净化 30min。质谱分析是在多接收稀有气体质谱仪 Helix MC 上进行的,每个峰值均采集 20 组数据。所有的数据在回归到时间零点值后再进行质量歧视校正、大气氩校正、空白校正和干扰元素同位素校正。中子照射过程中所产生的干扰同位素校正系数通过分析照射过的 K_2SO_4 和 CaF_2 来获得,其值为: $(^{36}\text{Ar}/^{37}\text{Ar}_0)_{\text{Ca}} = 0.0002389$, $(^{40}\text{Ar}/^{39}\text{Ar})_{\text{K}} = 0.004782$, $(^{39}\text{Ar}/^{37}\text{Ar}_0)_{\text{Ca}} = 0.000806$ 。 ^{37}Ar 经过放射性衰变校正; ^{40}K 衰变常数 $\lambda = 5.543 \times 10^{-10} \text{ y}^{-1}$ (Steiger and Jager, 1977);用 ISOPLLOT 程序计算坪年龄及正、反等时线(Ludwig,

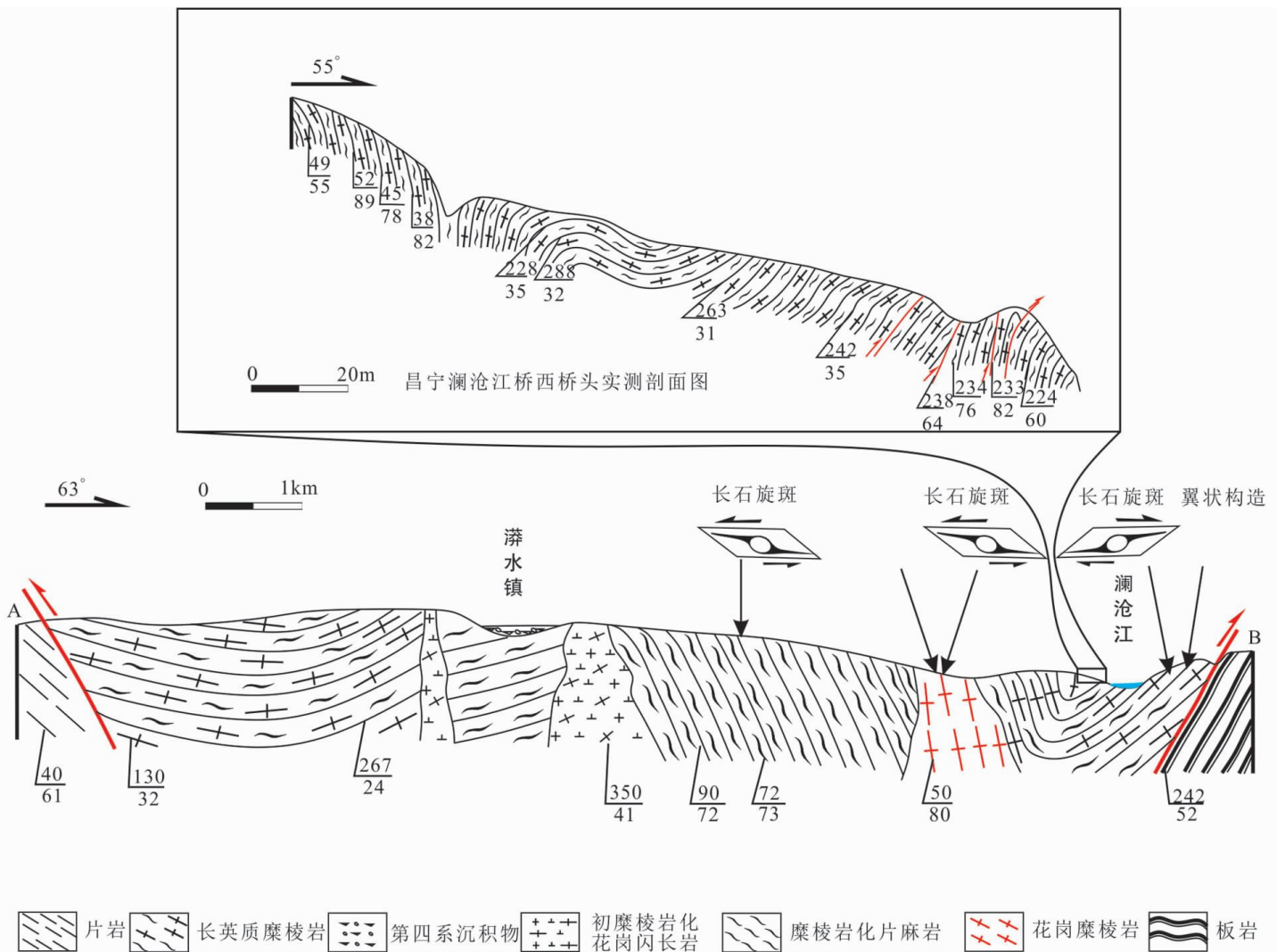


图2 澜沧江构造带中段昌宁-耆街剖面构造剖面图

Fig. 2 Changning-Goujie structural section in the middle segment of the Lancang tectonic zone

2001)。坪年龄误差以 2σ 给出。详细实验流程参考文献(陈文等, 2006; 张彦等, 2006)。

4 构造变形特征

4.1 韧性剪切变形

韧性剪切带内的岩石都发生不同程度的韧性变形, 其中的石英具有明显的波状消光、亚颗粒和动态重结晶现象, 长石多发育机械双晶、核幔结构和应力出溶作用(应力蠕英结构、应力条纹结构)、不对称旋转碎斑、多米诺构造, 以及长英质条带定向排列构成糜棱面理和拉伸线理, 局部形成 S-C 组构等。构造带内的岩石都发生不同程度的糜棱岩化, 形成超糜棱岩、糜棱岩和初糜棱岩, 其中, 糜棱岩构成剪切带的主体, 超糜棱岩主要分布于剪切带西南部, 初糜棱岩不均匀分布。矿物组合主要有: 石榴石 + 黑云母 + 斜长石 + 石英、角闪石 + 黑云母 + 长石、黑云母/白云母 + 斜长石 + 钾长石 + 石英。糜棱面理产状变化较大, 但整体上具有西部倾向南东, 倾角 30° 左右, 中部走向近南北, 近直立, 东部则变为倾向西, 倾角 25° 左右的变化趋势, 且在局部存在“背形”或“向

形”式的变化(图 2、图 3)。拉伸线理相对稳定, 向南南东倾伏, 倾伏角 $10^\circ \sim 25^\circ$ 。局部同一面理上发育两组斜交的线理 ($120^\circ \sim 157^\circ \angle 6^\circ \sim 15^\circ$ 和 $308^\circ \sim 337^\circ \angle 29^\circ \sim 39^\circ$), 且倾伏角较大的线理被近水平线理切割(图 4a), 反映两期走滑剪切变形特征。

韧性剪切带内指向构造发育。露头尺度上和显微尺度下, 在平行拉伸线理和垂直面理的 XZ 面上, 长英质条带被拉断旋转形成的不对称旋转透镜体和眼球状旋转碎斑、长石及石英团块构成的“ δ ”或“ σ ”旋转碎斑(图 4)、云母鱼、石英被拉长斜列与云母等片状矿物一起构成 S-C 组构、以及长石沿解理面发生滑动形成的“多米诺”构造(图 5)显示出明显的走滑性质。其中, 这些构造指向标志出构造带中-西部以左行走滑为主, 东部以右行走滑为主的特点(图 4)。

4.2 脆韧性变形

该阶段构造变形继承了早期糜棱面理, 局部存在糜棱面理被拉断、错开, 并被同构造石英充填形成的翼状构造 (Passchier, 2001; Mukherjee, 2014)(图 4f), 长英质脉体被错断而形成的“书斜”构造(图 4e)和 S-C 组构, 以及标志右

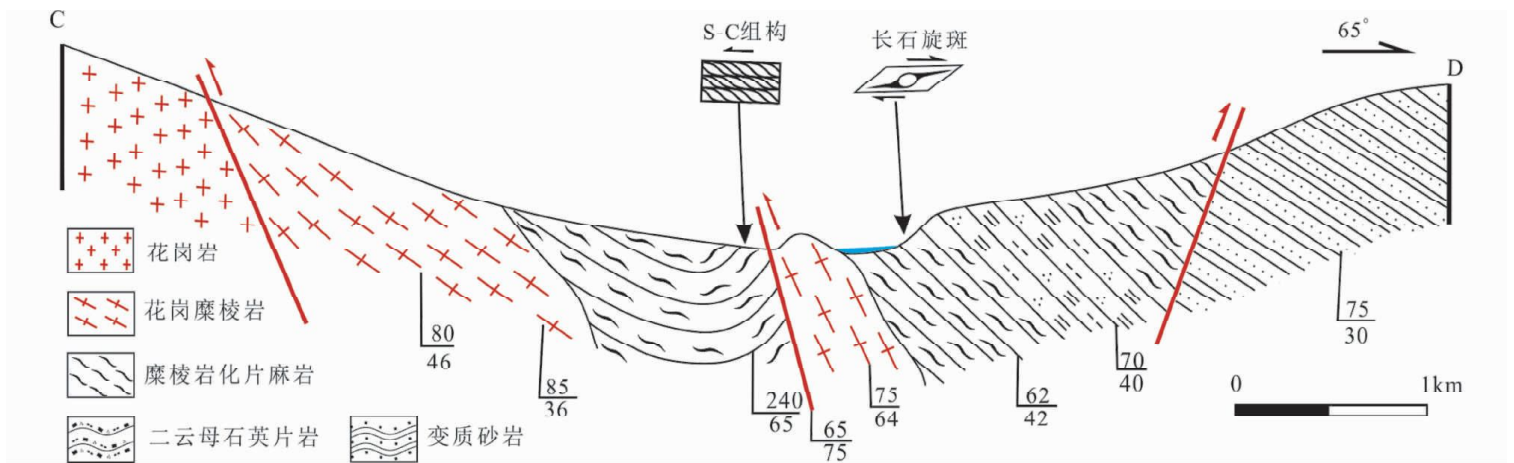


图3 澜沧江构造带中段凤庆-鲁史剖面构造剖面图

Fig. 3 Fengqing-Lushi structural section in the middle segment of the Lancang tectonic zone

行剪切作用的云母鱼和钾长石被后期左行剪切作用错开(图5e)等脆韧性变形特征,并在整个剪切带中都展示出左行剪切的运动机制,反映其形成于韧性剪切带隆升到中上部构造层次后发生的构造变形。

总体来看,澜沧江构造带中段韧性剪切带整体上具有正花状构造特征,即糜棱面理在西部呈低角度倾向北东东,东部低角度倾向南西西,中部面理近直立。韧性剪切带早期以右行韧性走滑剪切变形,后期以左行脆韧性变形走滑剪切变形为特征。

5 石英组构特征

石英晶格优选方位是查明韧性剪切变形走滑方向和变形条件的重要方法。为此,笔者分别对剪切带北部的昌宁-耆街剖面和南部的凤庆-鲁史剖面中的糜棱岩样品进行EBSD石英C轴组构分析。其中12QMZ-1-1、13QCG-13和13QCG-25分别采自昌宁-耆街剖面西翼、中部和东翼的长英质糜棱岩;样品12QFL-8、12QFL-36和12QFL-53分别采自凤庆-鲁史剖面西南翼的花岗质糜棱岩、中部的长英质糜棱岩和东北翼的长英质糜棱岩。测试结果如图6所示。

从图6可以看出,所有样品石英C轴晶格优选方位极密均存在平行于Y轴和分布于Z轴附近的2个极密,前者为柱面<a>滑移系,反映形成于中温(450~600℃)的变形环境(Schmid and Casey, 1986),后者为叠加于前期的中低温菱面<a>滑移系和低温底面<a>滑移系的单斜单环带组构类型,变形温度约300~550℃(Takeshita, 1996)。昌宁-耆街剖面中、东部和凤庆-鲁史剖面样品石英C轴晶格优选方位图中温变形显示右行剪切指向,后期中低温变形显示左行特征,昌宁-耆街剖面西翼则显示两期左行特征,与剖面整体呈向形构造相符。样品石英C轴晶格优选方位图结果与野外露头 and 镜下观察剪切标志相符。

6 年代学研究

6.1 样品描述

黑云母⁴⁰Ar-³⁹Ar定年样品采自于的昌宁-耆街剖面中黑云母片岩(12QMZ-15, 24°54'12"N, 99°43'38"E)和花岗糜棱岩(12QMZ-34, 24°56'53"N, 99°43'58"E)(图1b)。12QMZ-15黑云母片岩主要矿物组成为黑云母(45%),白云母(5%)、石英(40%)、长石(10%)。其中黑云母呈他形片状定向分布,未发生任何蚀变,石英呈条带状,间隔分布于云母富集带间,定向排列的黑云母、多晶条带状石英构成面理,含少量保留斑晶形态的细粒化长石,其边部零星见蠕英结构,反映岩石经历过韧性变形和动态重结晶作用。12QMZ-34花岗质糜棱岩主要矿物组成为长石(55%)、石英(35%)、黑云母(9%),副矿物主要有锆石、磷灰石、榍石及磁铁矿等不透明矿物(1%),长石残斑粒度变化较大,约1×3mm~3×5mm,长石机械双晶、细粒化作用显著,局部发育蠕英结构。云母呈残缕状,未发生任何蚀变,与细粒化的长石、动态重结晶石英颗粒组成基质。

6.2 定年结果

2个样品⁴⁰Ar-³⁹Ar年龄分析的数据见表1,坪年龄和等时线年龄见图7。12QMZ-15样品中黑云母840~1060℃的7个加热步给出了近一致的表现年龄,对应的坪年龄为17.8±0.2Ma(MSWD=0.83),析出³⁹Ar量占总量的66.5%;等时线年龄为17.7±0.3Ma(MSWD=1.4)。12QMZ-34样品中黑云母840~1170℃的10个加热步给出了近一致的表现年龄,对应的坪年龄15.0±0.2Ma(MSWD=0.59),析出³⁹Ar量占总量的85.3%;等时线年龄为14.7±0.4Ma(MSWD=1.04),两样品初始⁴⁰Ar/³⁶Ar值分别为296.7±5.5和300.4±7.8,与现今大气氩的标准值(295.5±5)误差范围内一致,反映样品受污染的程度较小,且等时线年龄与坪年龄在误差范围内

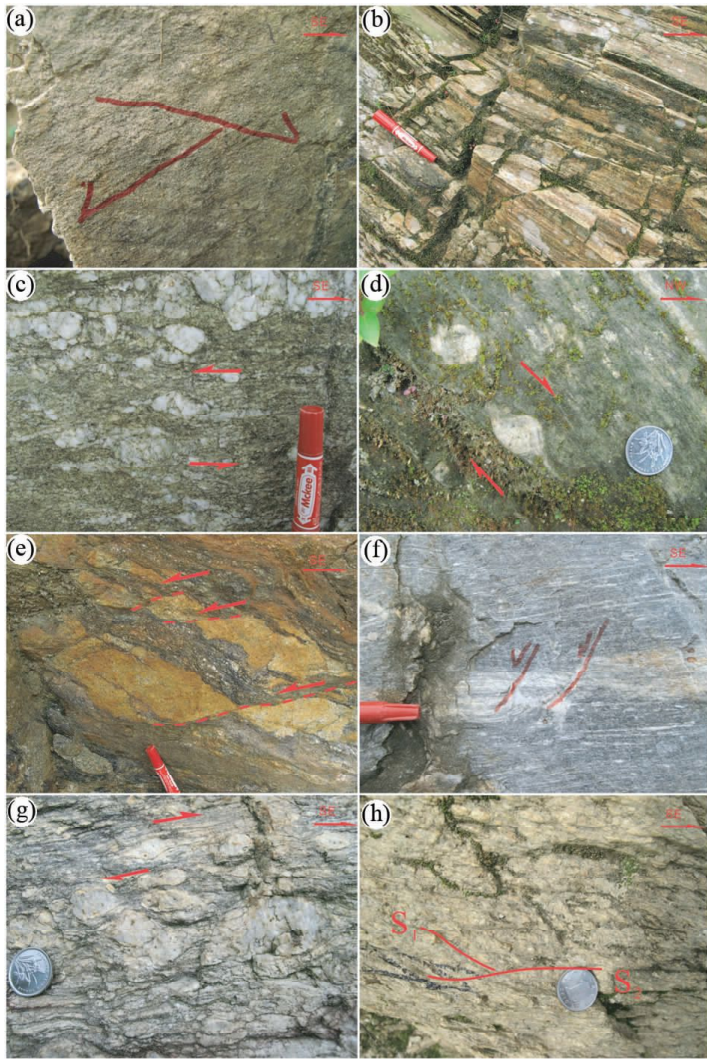


图4 澜沧江构造带中段韧性变形特征

(I) 昌宁-耆街剖面:(a)剪切带东翼花岗质糜棱岩而理面上发育的两期线理;(b)剪切带西翼超糜棱岩及糜棱线理;(c)剪切带西翼花岗质糜棱岩发育 σ 型旋斑,指示左行(XZ面);(d)剪切带东翼花岗质糜棱岩发育 σ 型旋斑, σ 型旋斑指示右行(XZ面);(e)剪切带西翼云母片岩中的浅色花岗岩脉被剪切错开,指示后期脆韧性左行剪切(XZ面);(f)剪切带东翼长英质糜棱岩发育的翼状构造指示左行脆韧性变形(XZ面)。(II) 凤庆-鲁史剖面:(g)剪切带东翼花岗质糜棱岩发育 σ 型旋斑, σ 型旋斑指示右行(XZ面);(h)剪切带东翼花岗质糜棱岩发育两期面理(XZ面)

Fig. 4 Ductile deformation characterization for the middle segment of the Lancang tectonic zone

(I) Changning-Goujie section: (a) two generation mineral stretching lineations on the foliation in granitic mylonite, the east side of the shear zone; (b) ultramylonite in the west side of the shear zone; (c) σ -type porphyroblast in granitic mylonite, indicating sinistral shearing (XZ), west side of the shear zone; (d) σ -type porphyroblast in granitic mylonite, indicating dextral shearing (XZ), east side of the shear zone; (e) dislocation structure in mica schist, indicating brittle-ductile sinistral shearing in later stage (XZ), west side of the shear zone; (f) flanking structures in felsic mylonite, indicating brittle-ductile sinistral shearing (XZ), east side of the shear zone. (II) Fengqing-Lushi section: (g) σ -type porphyroblast in granitic mylonite, indicating dextral shearing (XZ), east side of the shear zone; (h) two generation foliations in granitic mylonite (XZ), east side of the shear zone

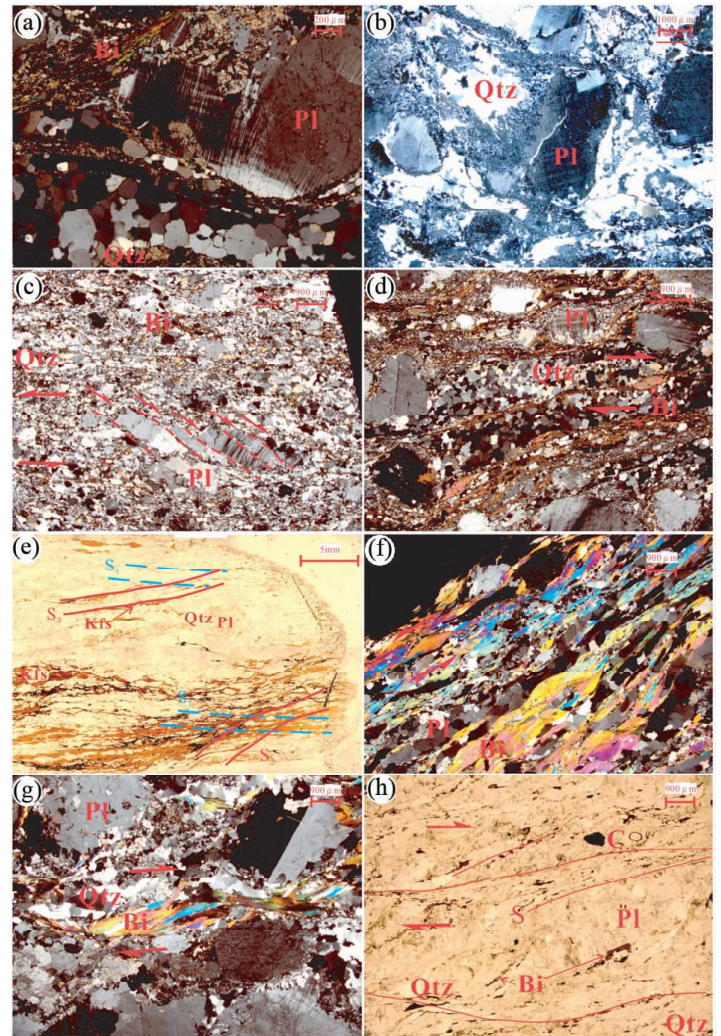


图5 澜沧江构造带中段韧性剪切带显微构造变形特征

(I) 凤庆-鲁史剖面:(a)剪切带西翼花岗质糜棱岩中长石的机械双晶(正交偏光);(b)剪切带西翼花岗质糜棱岩中的核幔构造(正交偏光);(c)剪切带西翼花岗质糜棱岩中长石书斜构造,反映后期脆韧性左行走滑剪切(正交偏光);(d)剪切带东翼花岗质糜棱岩中的云母鱼旋斑指示右行(正交偏光)。(II) 昌宁-耆街剖面:(e)剪切带西翼黑云母片岩中发育两期面理(单偏光);(f)剪切带西翼黑云母片岩中的云母鱼指示左行(正交偏光);(g)剪切带东翼花岗质糜棱岩中云母鱼指示右行(正交偏光);(h)剪切带东翼花岗质糜棱岩中指示右行的S-C组构(单偏光). Qtz-石英;Bi-黑云母;Kfs-钾长石;Pl-斜长石

Fig. 5 Microstructural characteristics in the middle segment of the Lancang tectonic zone

(I) Fengqing-Lushi section: (a) plagioclase deformation twins in granitic mylonite (CPL), west side of the shear zone; (b) core and mantle structure in granitic mylonite (CPL), west side of the shear zone; (c) domino structure in granitic mylonite, indicating sinistral shearing (CPL), west side of the shear zone; (d) mica fish in granitic mylonite, indicating dextral shearing (CPL), west side of the shear zone. (II) Changning-Goujie cross section: (e) two generation foliations in biotite schist, west side of the shear zone (PPL); (f) mica fish in biotite schist, indicating sinistral shearing (CPL), east side of the shear zone; (g) mica fish in granitic mylonite, indicating dextral shearing (CPL), east side of the shear zone; (h) S/C fabric in granitic mylonite, indicating dextral shearing (PPL), east side of the shear zone. Qtz-quartz; Bi-biotite; Kfs-K-feldspar; Pl-plagioclase

表1 澜沧江构造带中段韧性剪切带内黑云母 $^{40}\text{Ar}/^{39}\text{Ar}$ 阶段升温测年数据Table 1 $^{40}\text{Ar}/^{39}\text{Ar}$ analytical data of biotite from shear zone in the middle segment of the Lancang tectonic zone

序号	T ($^{\circ}\text{C}$)	累积 ^{39}Ar (%)	年龄 (Ma)	误差 (1 σ)	Ca/K	序号	T ($^{\circ}\text{C}$)	累积 ^{39}Ar (%)	年龄 (Ma)	误差 (1 σ)	Ca/K
12QMZ-15-3 黑云母 $J=0.003742$						12QMZ-34-3 黑云母 $J=0.003803$					
1	800	2.22	15.47	0.84	0.2161	1	760	0.91	3.90	2.80	0.1199
2	840	12.08	17.87	0.27	0.0000	2	800	3.87	9.77	0.96	0.2193
3	880	31.57	17.75	0.20	0.0313	3	840	13.09	14.28	0.32	0.0670
4	920	47.55	17.74	0.21	0.0257	4	870	27.09	14.87	0.23	0.0125
5	950	54.76	18.13	0.27	0.0509	5	910	38.89	14.99	0.22	0.0394
6	980	59.56	17.68	0.30	0.1055	6	960	45.58	15.13	0.28	0.0000
7	1020	64.66	17.84	0.43	0.0639	7	1020	56.69	15.34	0.24	0.0559
8	1060	68.74	17.00	0.45	0.1778	8	1070	78.07	14.86	0.17	0.0402
9	1100	74.45	18.89	0.42	0.0266	9	1120	95.31	15.10	0.22	0.0543
10	1150	86.66	18.6	0.28	0.0000	10	1170	98.38	15.40	0.91	0.1943
11	1400	100.00	17.21	0.23	0.1141	11	1400	100.00	7.70	1.70	0.5945

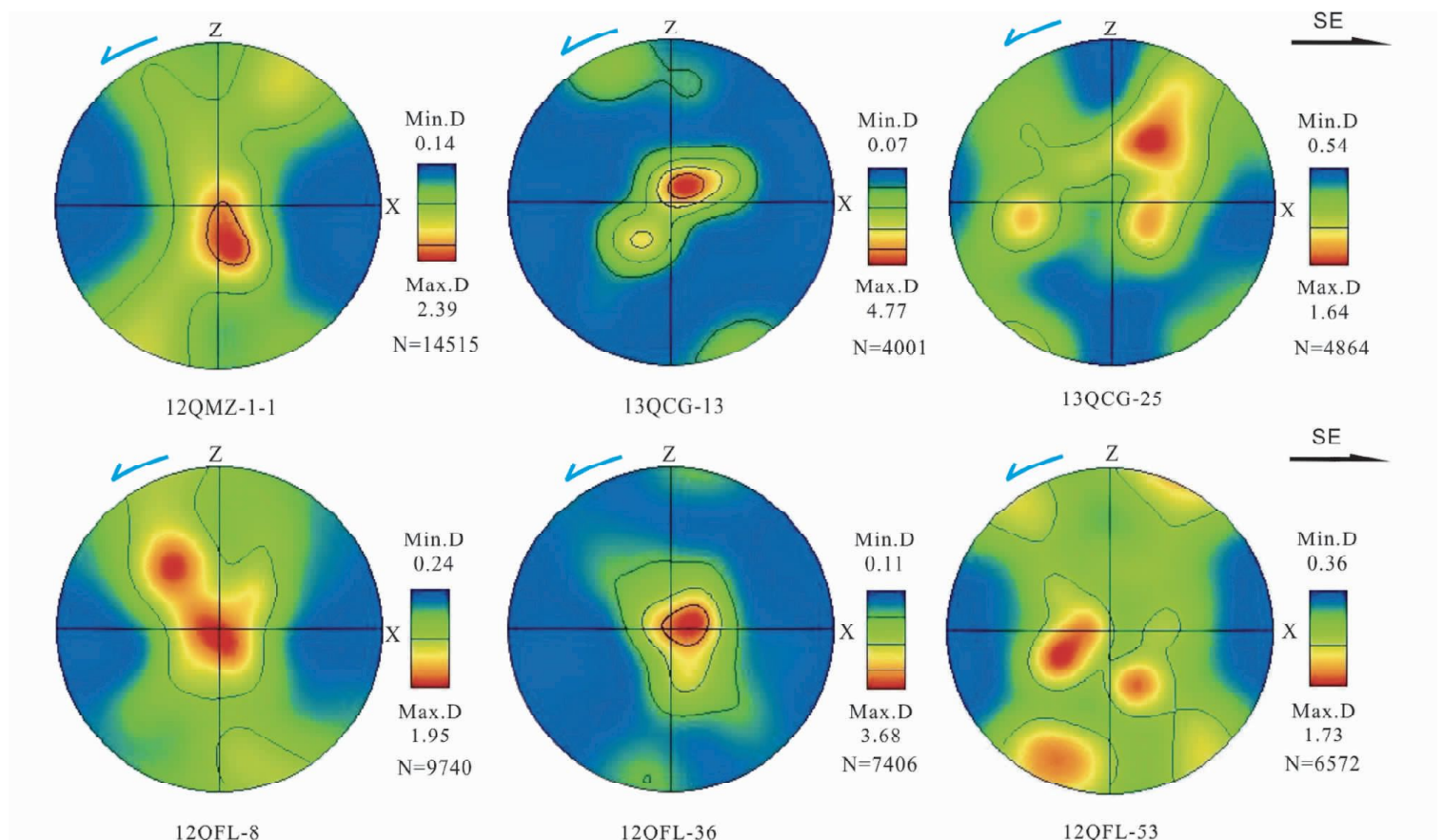


图6 石英c轴晶轴优选方位图(箭头指示后期叠加左行剪切运动方向)

Fig. 6 C-axis LPO patterns of quartz (arrows indicate the later stage sinistral shear direction)

一致,年龄真实可信。

7 讨论

7.1 走滑作用时代

近年来,前人对澜沧江构造带北段(即碧罗雪山-崇山一带)韧性剪切变形特征及其形成时代进行了研究。通过对韧性剪切带中糜棱岩、同构造花岗岩、变形岩脉及未变形岩脉

等进行独居石/锆石 U-Pb 和黑云母/白云母 $^{40}\text{Ar}-^{39}\text{Ar}$ 同位素年代学测试(Akciz *et al.*, 2008; Zhang *et al.*, 2010, 2012a; 唐渊等, 2013),结果表明同构造花岗岩中独居石/锆石、变形花岗岩脉中锆石变质边 U-Pb 年龄为 37 ~ 21Ma,糜棱岩、同构造花岗岩、变形岩脉中黑云母/白云母 $^{40}\text{Ar}-^{39}\text{Ar}$ 分别为 13 ~ 19Ma、14 ~ 20Ma,说明韧性剪切作用形成于 37 ~ 21Ma。黑云母和白云母封闭温度(350 $^{\circ}\text{C}$ 和 400 $^{\circ}\text{C}$)的特点说明该年龄反映深变质变形带隆升的时代。前人对构造带中段糜棱

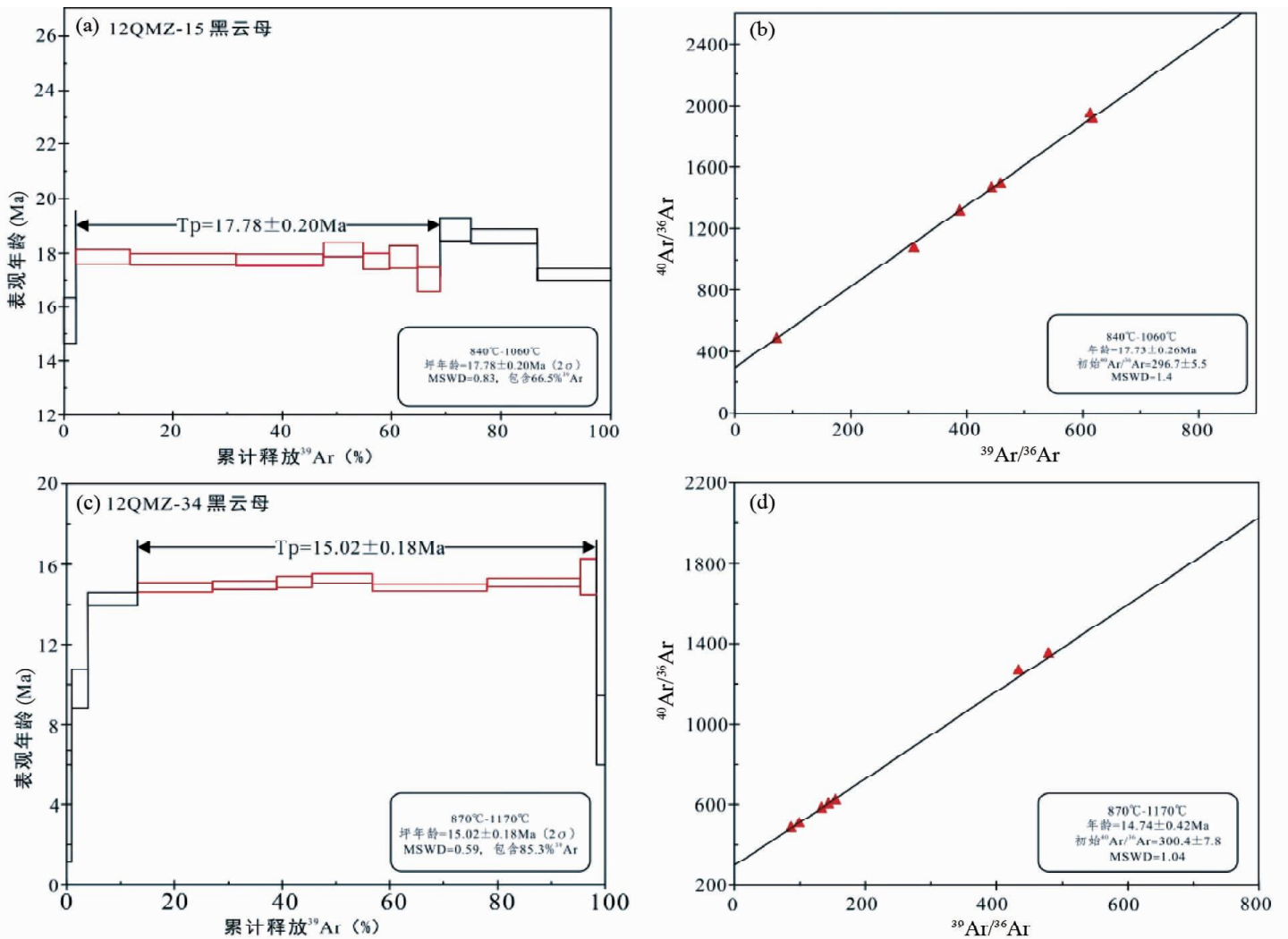


图7 澜沧江构造带中段韧性剪切带内黑云母 $^{40}\text{Ar}/^{39}\text{Ar}$ 坪年龄(a、c)和等时线年龄(b、d)

Fig. 7 $^{40}\text{Ar}/^{39}\text{Ar}$ plateau age spectra (a, c) and isochron ages (b, d) of biotite from shear zone in middle segment of the Lancang tectonic zone

岩中的黑云母进行了 $^{40}\text{Ar}-^{39}\text{Ar}$ 同位素测试,获得27~28Ma(陈新跃等,2006;Wang *et al.*,2006)和17Ma(李述靖和单业华,1995)两组同位素数据。野外和镜下观察表明构造带中段存在早期韧性和后期左行脆韧性两个阶段的变形,前者形成于中高温条件($450 \sim 600^\circ\text{C}$),后者形成于中低温($350 \sim 550^\circ\text{C}$)条件,本文获得的黑云母 $^{40}\text{Ar}-^{39}\text{Ar}$ (17~15Ma)与前人获得两组年龄数据中的17Ma年龄数据一致,是对后期左行脆韧性变形期时代上限、韧性剪切变形时代下限的限定,而27~28Ma $^{40}\text{Ar}-^{39}\text{Ar}$ 同位素年龄位于北段与韧性剪切变形有关花岗岩锆石U-Pb年龄区间,代表韧性剪切变形形成和隆升时代。

7.2 构造演化

青藏高原东南缘三江构造域是经历多期次构造运动改造形成的块体拼合、挤出和大规模走滑作用构造变形域,尤其是新生代印度板块向北俯冲碰撞导致青藏高原东南缘块体——印支地块和Sibumasu地块发生大规模旋转和逃逸(Tapponnier and Molnar, 1976; Tapponnier *et al.*, 1982,

1990),块体沿古缝合带发生大规模走滑作用(丁林,1991;钟大赉等,1991;钟大赉,1998;Leloup *et al.*,1993,1995;刘俊来等,2006,2007,2011;罗照华等,2006;Morley,2007;Searle *et al.*,2007;王二七等,2006;Cao *et al.*,2010),形成东部以哀牢山-金沙江大规模左行走滑剪切带、西部以高黎贡右行韧性剪切带和中部澜沧江韧性剪切带为界的构造格局(图1a)。澜沧江韧性走滑剪切带的形成时代与前二者一致,反映其形成与兰坪-思茅地块和保山地块之间的相对运动有关。即在挤出过程中,兰坪-思茅地块向南挤出速度快于保山地块时,澜沧江韧性剪切带发生右行走滑剪切作用,形成一系列反映右行走滑剪切的指向构造,反之则形成左行走滑剪切变形。澜沧江构造带中段中部面理糜棱产状近直立,东、西两侧糜棱面理产状和构造指向都截然相反,展示出典型的“花”状构造,以及糜棱面理中保留有早期倾伏角较大的拉伸线理,揭示中下地壳形成的韧性剪切带在斜向挤出隆升过程中,韧性变形岩石以中部近直立糜棱面理糜棱岩为界,分别向东、西两侧斜向逆冲形成的,剪切带两侧边缘部位在中部物质向外挤出、并向东、西两侧逆冲过程中致使糜棱面理(S_1)发生弯曲变形,形成两侧分别为倾向相

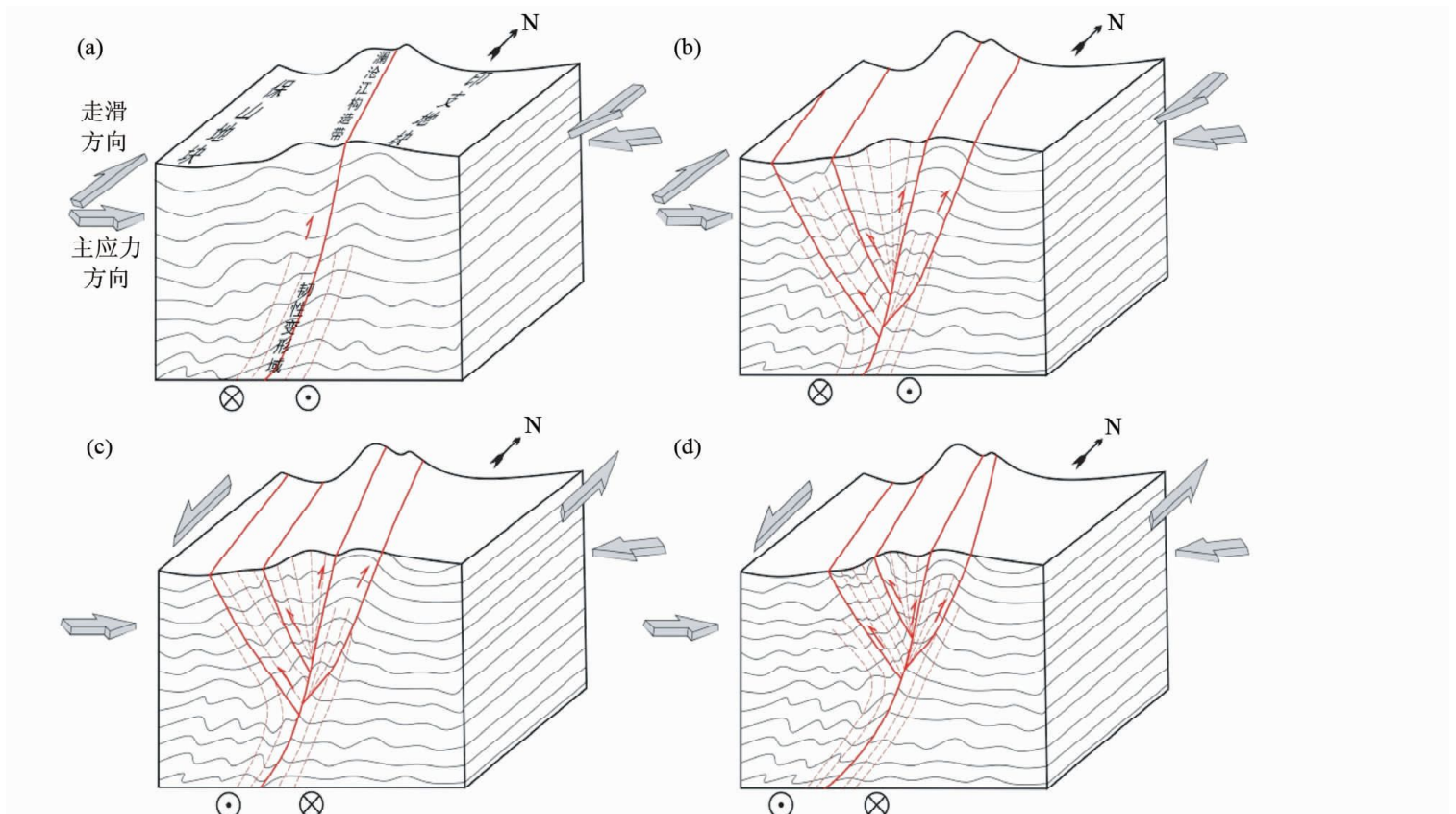


图8 澜沧江构造带中段构造变形过程示意图

Fig. 8 Structural deformation process evolution in the middle segment of the Lancang tectonic zone

反的两条逆冲断层带控制的正花状构造的雏形(图 8a, b)。在斜向挤出作用下韧性剪切带隆升到上部构造层次后,保山地块内北东走向的畹町和南汀河两条构造带发生左行走滑作用(Eroğlu *et al.*, 2013),致使南汀河构造带南部的物质(包括临沧岩体)向北东挤压,形成澜沧江构造带中段向北东弯曲的弧形,南段(临沧岩体)直接逆冲于兰坪-思茅盆地之上的构造特征(杨振德, 1996; Wang and Burchfiel, 1997; 段建中, 1999; 段建中和谭筱红, 2000),在此过程中,韧性剪切带东部物质相对向北运移,在韧性剪切带内叠加左行脆韧性变形构造(图 8c, d)。

8 结论

(1) 澜沧江构造带中段经历了两阶段变形事件:早期南西向北东斜冲右行韧性变形过程形成宏观韧性变形标志东部糜棱面理低角度倾向西、西部低角度倾向东、中部近直立,两侧分别逆冲于未发生变形变质的中生代沉积岩之上的花状构造;后期构造带整体经历近水平左行脆韧性剪切变形作用。

(2) 黑云母 $^{40}\text{Ar}-^{39}\text{Ar}$ 定年结果及前人同位素年代学数据表明澜沧江构造带中段早期韧性走滑剪切作用变形作用不晚于 28Ma, 止于 17 ~ 15Ma, 脆韧性变形起始于 17 ~ 15Ma。

致谢 EBSD 测试工作由陈方远副研究员完成, $^{40}\text{Ar}-^{39}\text{Ar}$

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